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# **N<sub>2</sub>O emissions from eutrophic lakes: sources and significance**

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## Abstract

For decades, high emissions of the greenhouse gas and ozone-depleting pollutant nitrous oxide ( $\text{N}_2\text{O}$ ) have been repeatedly reported from eutrophic environments. Acknowledging this fact, the Intergovernmental Panel on Climate Change (IPCC) recently increased the emission factor (EF) used to compute indirect  $\text{N}_2\text{O}$  emissions from wastewater discharge into eutrophic and nutrient-impacted aquatic environments from 0.005 to 0.019 kg  $\text{N}_2\text{O}$ -N emitted per kg of N received. However, because the IPCC still considers that bacterial nitrification and denitrification are the only significant  $\text{N}_2\text{O}$  biological mechanisms, it computes  $\text{N}_2\text{O}$  emissions by assuming a linear relationship between the amount of  $\text{N}_2\text{O}$  emitted from an aquatic environment and the amount of nitrogen (N) reaching this environment. This bacteria-centric assumption may be challenged by the ability of microalgae to synthesize significant amount of  $\text{N}_2\text{O}$ . Indeed, as microalgae blooms can be triggered by N and/or phosphorus (P) pollution, the ability of these organisms to produce  $\text{N}_2\text{O}$  may mean that  $\text{N}_2\text{O}$  emissions from eutrophic environments cannot only be correlated to N inputs. Thus, we hypothesize that microalgae significantly contribute to  $\text{N}_2\text{O}$  emissions in eutrophic aquatic environments and that P inputs should also be considered to accurately estimate  $\text{N}_2\text{O}$  emissions from eutrophic aquatic environments.

This thesis initially sought to determine if algae-rich eutrophic ecosystems indeed generate significant  $\text{N}_2\text{O}$  emissions and if microalgae indeed contribute to these emissions. To reduce the scope of the research, emphasis was given to the study of eutrophic lakes as a 'worse-case environment'. Our methodology was based around field monitoring and laboratory assays testing artificial and natural microalgae-based microcosms.

Preliminary data gathered from the eutrophic Lake Horowhenua (Levin, New Zealand) show that the lake is a source of  $\text{N}_2\text{O}$  ( $0.4 - 8.7 \text{ g N-N}_2\text{O}\cdot\text{ha}^{-1}\cdot\text{yr}^{-1}$ ,  $n = 29$  sampling events). However, no relationship between algal biomass concentration and  $\text{N}_2\text{O}$  production could be evidenced. In parallel, laboratory batch assays were performed to investigate physiological conditions influencing microalgal  $\text{N}_2\text{O}$  synthesis and the potential enzymes involved. Preliminary data showed that wild type *Chlamydomonas reinhardtii* produced  $\text{N}_2\text{O}$  up to  $31.5 \pm 10.7 \mu\text{mole N}_2\text{O}\cdot\text{g DW}^{-1}$  over 24 hours in autotrophic conditions and when supplied 10 mM  $\text{NO}_2^-$ . Inhibiting the electron flow coming from PSII in the same conditions inhibited  $\text{N}_2\text{O}$  production, suggesting that the electron transport chain is involved in  $\text{N}_2\text{O}$  synthesis. These results informed us on the  $\text{N}_2\text{O}$  synthesis pathways that should be investigated during future field work and allowed us to identify candidate strains to create a mutant unable to synthesize  $\text{N}_2\text{O}$  for future microcosms study.

Because accurate data is paramount to robust policy, this research is the first step to establish the significance of a new  $\text{N}_2\text{O}$  source that is not necessarily linked to N pollution and could trigger a paradigm shift in how  $\text{N}_2\text{O}$  emissions are estimated in greenhouse gas inventories. This may in turn affect how  $\text{N}_2\text{O}$  sources are prioritized for mitigation strategies and, if deemed necessary, will also provide the methodologies and knowledge needed for efficient monitoring of eutrophic aquatic bodies, including marine environments.

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## 1. Introduction

The Nature Education knowledge project defines eutrophication as “excessive plant and algal growth due to the increased availability of one or more limiting growth factors needed for photosynthesis (Schindler 2006), such as sunlight, carbon dioxide, and nutrient fertilizers” (Chislock et al., 2013). The New Zealand Ministry for the Environment is more specific and states that “eutrophication is the result of nutrients built up in a body of water, frequently due to run-off from land, which causes dense growth of aquatic plants or algae” (*Assessment of the Eutrophication Susceptibility of New Zealand’s Estuaries*, 2019). Indeed, the artificial input of nitrogen (N) and phosphorus (P) in water bodies has considerably increased the occurrence of eutrophication to an extent that it is now considered a global environmental issue and the overgrowth of algae and aquatic plants causes deterioration of water taste and odour, loss of transparency, and even toxicity in some cases (Xu et al., 2010; Aubriot et al., 2020; Huang et al., 2020). For example, the water quality of Rotorua Lakes in New Zealand has declined since the 1950s due to the intensive use of fertilizers and urbanisation (Morgenstern et al., 2015), even leading to the closure of Lake Rotoiti for 2 months during summer 2003 because of the toxicity of a large microalgal bloom (Hamilton, 2003). The New Zealand Ministry for the Environment has estimated that around half of New Zealand monitored lakes are eutrophic or worse (Ministry for the Environment & Stats NZ, 2020). Globally, the figure is similar with 30 – 40% of all lakes being eutrophic (Yang et al., 2008).

Another potential and poorly known consequence of eutrophication is the emission of nitrous oxide ( $\text{N}_2\text{O}$ ), a potent greenhouse gas (GHG) and ozone-depleting atmospheric pollutant (Ravishankara et al., 2009). Several authors have indeed estimated that global  $\text{N}_2\text{O}$  emissions from eutrophic lakes could represent 52 to 211 Tg  $\text{CO}_2\text{-year}^{-1}$  (DelSontro et al., 2018; Plouviez et al., 2019a; Plouviez & Guieysse, 2020), which represent 3 to 12% of all direct anthropogenic  $\text{N}_2\text{O}$  emissions from agriculture, the largest contributor to anthropogenic  $\text{N}_2\text{O}$  emissions globally (Tian et al., 2020). However, and while significant  $\text{N}_2\text{O}$  emissions have been repeatedly reported in microalgae-rich aquatic ecosystems (e.g. eutrophic lakes), we still do not know the magnitude of these emissions and how much of these emissions are directly caused by microalgae (Ciais et al., 2013; Plouviez et al., 2019a). This is particularly relevant when we consider that the ability of microalgae to produce  $\text{N}_2\text{O}$  has now been unambiguously demonstrated in the laboratory (Burlacot et al., 2020; Guieysse et al., 2013; Plouviez et al., 2019a; Weathers, 1984; Weathers & Niedzielski, 1986).

Acknowledging that high  $\text{N}_2\text{O}$  emissions have indeed been often recorded from “nutrient-impacted waters such as eutrophic lakes”, the Intergovernmental Panel on Climate Change (IPCC) recently increased the value of the Tier 3 emission factor used to compute  $\text{N}_2\text{O}$  emissions from “waters

receiving treated or untreated effluent” (see section 2.1 for explanation) from 0.05 to 0.019 kg N<sub>2</sub>O-N emitted per kg of N received (Bartram et al., 2019). However, it is clear from its narrative that the IPCC still assumes these N<sub>2</sub>O emissions are only occurring via bacterial nitrification and denitrification. This ‘bacteria-centric’ view implies that current N<sub>2</sub>O field monitoring and budgeting methodologies may not be accurate in eutrophic environments. This also means that mitigation strategies based solely on N control may not be efficient because microalgae growth can be triggered by P, meaning a dual control of nutrients might be required to reduce N<sub>2</sub>O emissions from eutrophic aquatic environments (Conley et al., 2009; Paerl et al., 2016). To identify knowledge gaps and propose a tangible research methodology, the following review is divided in 6 sections:

- Section 2.1 describes how N<sub>2</sub>O emissions are currently computed from aquatic environments and how these estimations compare to field data. This section identifies the knowledge gaps in the current assumptions and methodology used to estimate N<sub>2</sub>O emissions.
- Section 2.2 describes the currently known microbial N<sub>2</sub>O synthesis pathways, with emphasis on microalgae. This section provides the knowledge needed to understand the assumptions behind current N<sub>2</sub>O budgeting methodologies, and the uncertainty these assumptions are causing.
- Section 2.3 briefly describes the diversity of freshwater lakes and the tools used to evaluate their water quality. This section provides the basic knowledge to measure lakes’ trophic level.
- Section 2.4 describes the eutrophication process of lakes with emphasis on the nutrients involved and the proposed mitigation strategies. This section outlines the assumptions used in current mitigation strategies and help understand their potential inadequacy in microalgae-rich ecosystems.
- Section 2.5 describes fundamentals in algae cell biology with emphasis on algal N and P metabolisms. This section evidences the importance of understanding how N and P impact algal growth and, by extension, algal N<sub>2</sub>O synthesis.
- Section 2.6 outlines the ‘research gaps’ identified from the literature review, formulates our research hypothesis and concludes the review.

The literature review (Section 2) is followed by the description of the proposed research strategy and methodology (Sections 3 to 5). Finally, preliminary results are presented and discussed (Section 6).

## 2. Literature review

### 2.1. Anthropogenic N<sub>2</sub>O emissions from inland and coastal waters

N<sub>2</sub>O is an ozone-depleting pollutant (Ravishankara et al., 2009) and greenhouse gas (GHG) with a global warming potential 273 times higher than CO<sub>2</sub> on a 100-year time scale (Masson-Delmotte et al., 2021). Anthropogenic N<sub>2</sub>O emissions have increased by 23% over the past 40 years, reaching 7,300 kt N-N<sub>2</sub>O-year<sup>-1</sup> between 2007 and 2016. Observed atmospheric N<sub>2</sub>O concentrations are even starting to surpass all predictions by the IPCC (Tian et al., 2020).

**Box 1:** Intergovernmental Panel on Climate Change (IPCC).

The IPCC was founded in 1988 by the United Nations Environment Programme and the World Meteorological Organization to assess the science related to climate change. Its main objectives are to provide governments and policy makers with the tools to estimate GHG emissions, regular assessment reports regarding climate change, the consequences and risks resulting from climate change, and to come up with recommendations about adaptation and mitigation strategies.

The global sources of N<sub>2</sub>O are presented in Figure 1 (Tian et al., 2020) and further detailed in Appendix A. As can be seen, global N<sub>2</sub>O emissions from aquatic environments were estimated to 4,300 kt N-N<sub>2</sub>O-year<sup>-1</sup> from 2007 to 2016, with anthropogenic emissions from aquatic environments representing 14% of these emissions.

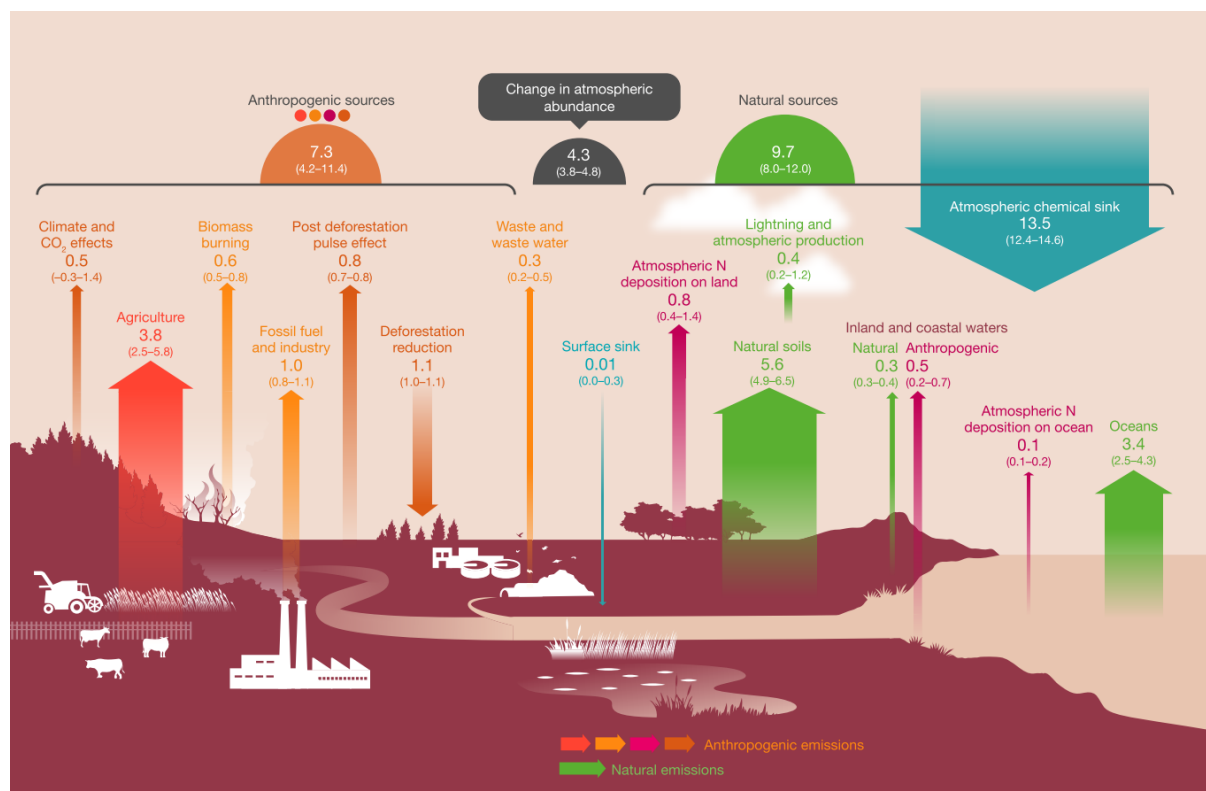


Figure 1: Global N<sub>2</sub>O budget. N<sub>2</sub>O fluxes are in Tg N-yr<sup>-1</sup> for 2007-2016.

Reprinted by permission from Springer Nature Customer Service Centre GmbH: Nature, A comprehensive quantification of global nitrous oxide sources and sinks, Tian et al., © 2020, doi: <https://doi.org/10.1038/s41586-020-2780-0>.

### 2.1.1. IPCC methodology

The IPCC assumes that anthropogenic N<sub>2</sub>O emissions from inland and coastal waters are indirectly caused by anthropogenic nitrogen (N) inputs from either i) managed lands (i.e. agricultural lands) following N leaching and runoffs from agricultural soils and N atmospheric deposition on water surfaces following N volatilization from land; and ii) N-rich wastewater discharge into water bodies (Figure 2). These indirect N<sub>2</sub>O emissions are then estimated by multiplying the N loads predicted to be received with specific emission factors (EFs, see Appendix B for further details) for specific activities and/or receiving environments. The EFs are calculated based on experimental measurements of N<sub>2</sub>O and N inputs in relevant ecosystems. There is currently no clear distinction between natural and anthropogenic N<sub>2</sub>O emissions from freshwater lakes and reservoirs (Appendix C.1) and, to our knowledge, Tian et al. (2020) are the first authors to propose a differentiation.

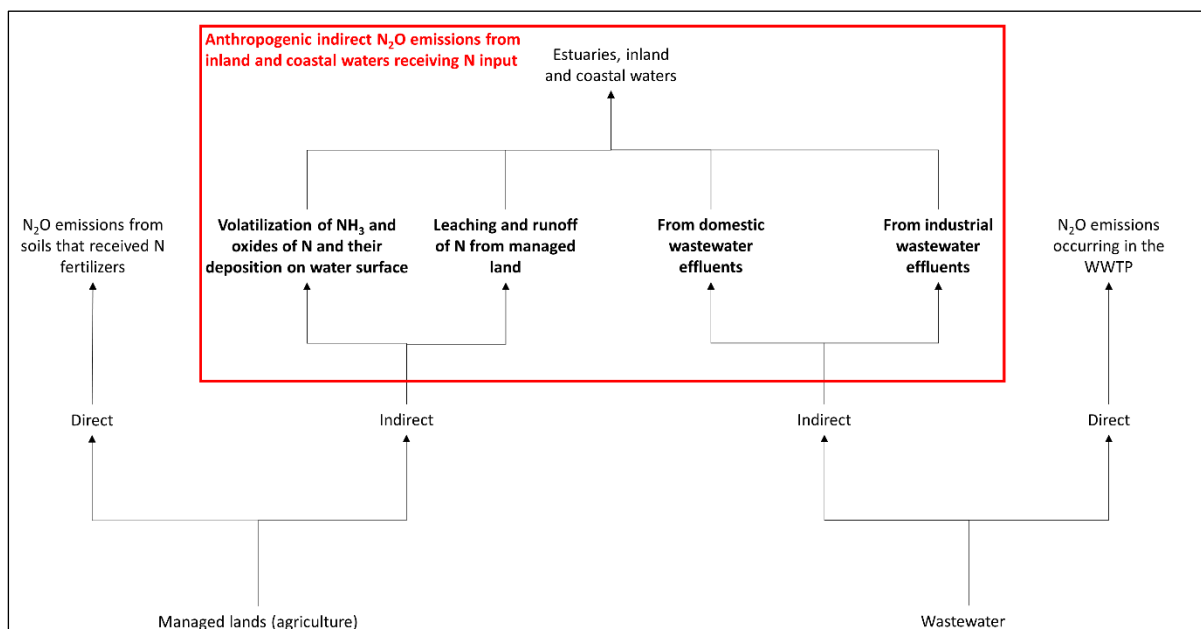


Figure 2: Anthropogenic N<sub>2</sub>O emissions from managed lands and wastewater according to the IPCC. The four N inputs to aquatic environments categories described by the IPCC causing indirect N<sub>2</sub>O emissions from water bodies are in bold character.

The IPCC advises to follow a three-tiered approach to calculate GHG emissions depending on the data available. The Tier 1 method enables to calculate GHG emissions using default EF values and, when applicable, default partitioning factors to estimate the N load (Appendix B). The Tier 2 method uses country-specific EFs and partitioning factors. Finally, the Tier 3 uses country-specific models and measurements (Hergoualc'h et al., 2019). For example, EF Tier 1 values of 0.01 and 0.011 kg N–N<sub>2</sub>O·kg N input<sup>-1</sup> are currently recommended to estimate indirect N<sub>2</sub>O emissions from aquatic environments receiving N via atmospheric deposition (EF<sub>4</sub>) and from agricultural runoff and leaching (EF<sub>5</sub>), respectively (Hergoualc'h et al., 2019) (Figure 3). A Tier 1 EF value of 0.005 kg N–N<sub>2</sub>O·kg N<sup>-1</sup> is also

recommended to estimate indirect N<sub>2</sub>O emissions from aquatic environments receiving N from domestic and industrial wastewater effluents, as exemplified in Box 2 (Bartram et al., 2019).

**Box 2:** N<sub>2</sub>O emissions from wastewater treatment plants and wastewater effluents (Bartram et al., 2019).

Direct N<sub>2</sub>O emissions from domestic wastewater treatment plants are calculated as:

$$N_2O \text{ Plants}_{DOM} = \left[ \sum_{i,j} (U_i \cdot T_{ij} \cdot EF_j) \right] \cdot TN_{DOM} \cdot \frac{44}{28}$$

Where N<sub>2</sub>O Plants<sub>DOM</sub> represents N<sub>2</sub>O emissions from domestic wastewater treatment plants (kg N<sub>2</sub>O·yr<sup>-1</sup>); TN<sub>DOM</sub> is the total amount of nitrogen in domestic wastewater (kg N·yr<sup>-1</sup>); U<sub>i</sub> is the fraction of population in income group i (with i = rural, urban high income and urban low income); T<sub>ij</sub> is the degree of utilization of treatment system j for each income group fraction i (with j = treatment/discharge pathway or system); EF<sub>j</sub> is the emission factor for treatment/discharge pathway or system j (kg N–N<sub>2</sub>O·kg N<sup>-1</sup>); and  $\frac{44}{28}$  is the conversion factor of kg N<sub>2</sub>O – N into kg N<sub>2</sub>O.

Indirect N<sub>2</sub>O emissions from wastewater effluents are calculated as a fraction of the nitrogen flux reaching the aquatic body. The removal of N from wastewater accounts for the transfer of N to sludge as well as N<sub>2</sub> and N<sub>2</sub>O losses to the atmosphere due to nitrification/denitrification processes during transport and treatment. Indirect N<sub>2</sub>O emissions from domestic wastewater effluents are therefore calculated as:

$$N_2O_{EFFLUENT,DOM} = N_{EFFLUENT,DOM} \cdot EF_{EFFLUENT} \cdot \frac{44}{28}$$

Where N<sub>2</sub>O<sub>EFFLUENT,DOM</sub> represents N<sub>2</sub>O emissions from domestic wastewater effluent (kg N<sub>2</sub>O·yr<sup>-1</sup>); N<sub>EFFLUENT,DOM</sub> is the nitrogen in the effluent discharged to aquatic environments (kg N·yr<sup>-1</sup>); and EF<sub>EFFLUENT</sub> (default = 0.005 kg N–N<sub>2</sub>O·kg N<sup>-1</sup>) is the emission factor for N<sub>2</sub>O emissions from wastewater discharged to aquatic systems (kg N–N<sub>2</sub>O·kg N<sup>-1</sup>).

$$N_{EFFLUENT,DOM} = \sum_j [(TN_{DOM} \cdot T_j) \cdot (1 - N_{REM,j})]$$

Where, T<sub>j</sub> is the degree of utilisation of treatment system j (with j corresponding to each wastewater treatment type used) and N<sub>REM</sub> is the fraction of total wastewater nitrogen removed during wastewater treatment per treatment type j.

Similar equations are used to estimate indirect N<sub>2</sub>O emissions from industrial wastewater effluents.

### 2.1.2. Limitations of current EFs

Based on the methodology described above, the IPCC estimates that rivers, estuaries and coastal zones emitted 600 kt N–N<sub>2</sub>O·year<sup>-1</sup> in 2006 (Ciais et al., 2013). Most authors however agree there is a large uncertainty associated with the IPCC methodology due to a lack of data and many authors tested different approaches to estimate N<sub>2</sub>O emissions from freshwater ecosystems, as further detailed in Appendix C (Hu et al., 2016; Miao et al., 2020; Tian et al., 2019; Webb et al., 2021). An additional limitation is that N<sub>2</sub>O emissions are assumed to increase linearly with the N load and that N<sub>2</sub>O synthesis is mostly caused by bacterial nitrification and denitrification, as further discussed below.



There is also a lack of specificity for the EFs used in different environments. For example, the same EF of  $0.0026 \text{ kg N-N}_2\text{O}\cdot\text{kg N leached and runoff}^{-1}$  is used for reservoirs and rivers despite a significant difference, documented at  $0.17 - 0.44 \text{ kg N-N}_2\text{O}\cdot\text{kg N leached and runoff}^{-1}$  and  $0.004 - 0.005 \text{ kg N-N}_2\text{O}\cdot\text{kg N leached and runoff}^{-1}$  respectively (Maavara et al., 2019). The difference could be explained by different hydraulic retention time (HRT) influencing the biological processes behind  $\text{N}_2\text{O}$  production (Maavara et al., 2019). Consequently, the current IPCC estimates for  $\text{N}_2\text{O}$  emissions from rivers might be overestimated.

The same issue has been highlighted regarding the EFs used to compute  $\text{N}_2\text{O}$  emissions from groundwater and surface drainage. Webb et al. (2021) argued that indirect emissions coming from artificial agricultural surface waters (drains, ditches, canals, farm dams and farm impoundments) were overestimated when using EFs developed by the IPCC, especially because the assumption that  $\text{N}_2\text{O}$  emissions linearly increase with N concentration does not take into account the drivers of  $\text{N}_2\text{O}$  production in these ecosystems (e.g. HRT and substrate availability to support microbial activity) (Webb et al., 2021). Besides, it has been pointed out that most EFs have been calculated using data from eutrophic waters, which could lead to an overestimation for non-eutrophic waters (Tian et al., 2019). Finally, even if the EFs do not seem to be dependent on climate, most data available come from temperate regions and might not be representative of the emissions under other climates (Tian et al., 2019).

### **2.1.3. The IPCC acknowledges higher $\text{N}_2\text{O}$ emissions occur in “nutrient-impacted environment”.**

Interestingly, the IPCC recently increased the EF associated with wastewater discharge into “nutrient-impacted waters” nearly 4-folds, from  $0.005$  to  $0.019 \text{ kg N}_2\text{O-N}\cdot\text{kg N}^{-1}$  (Table 1). This change was based on “research published between 1978 and 2017 [...] indicating that higher  $\text{N}_2\text{O}$  emissions occur when wastewater is discharged to nutrient-impacted (eutrophic) or hypoxic<sup>1</sup> aquatic receiving environments” (Bartram et al., 2019). It is however clear from the IPCC rationale for this change that the higher  $\text{N}_2\text{O}$  emissions experimentally recorded are largely due to the formation of hypoxic zones prone to  $\text{N}_2\text{O}$  emissions from bacteria. For this reason, the IPCC does not recommend to increase the EF used to compute  $\text{N}_2\text{O}$  emissions associated with N runoffs from agriculture as it states that a

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<sup>1</sup> Oxygen-depleted condition

“combination of reducing conditions and high organic loading [...] are unlikely to exist in agricultural runoff systems” (Bartram et al., 2019).

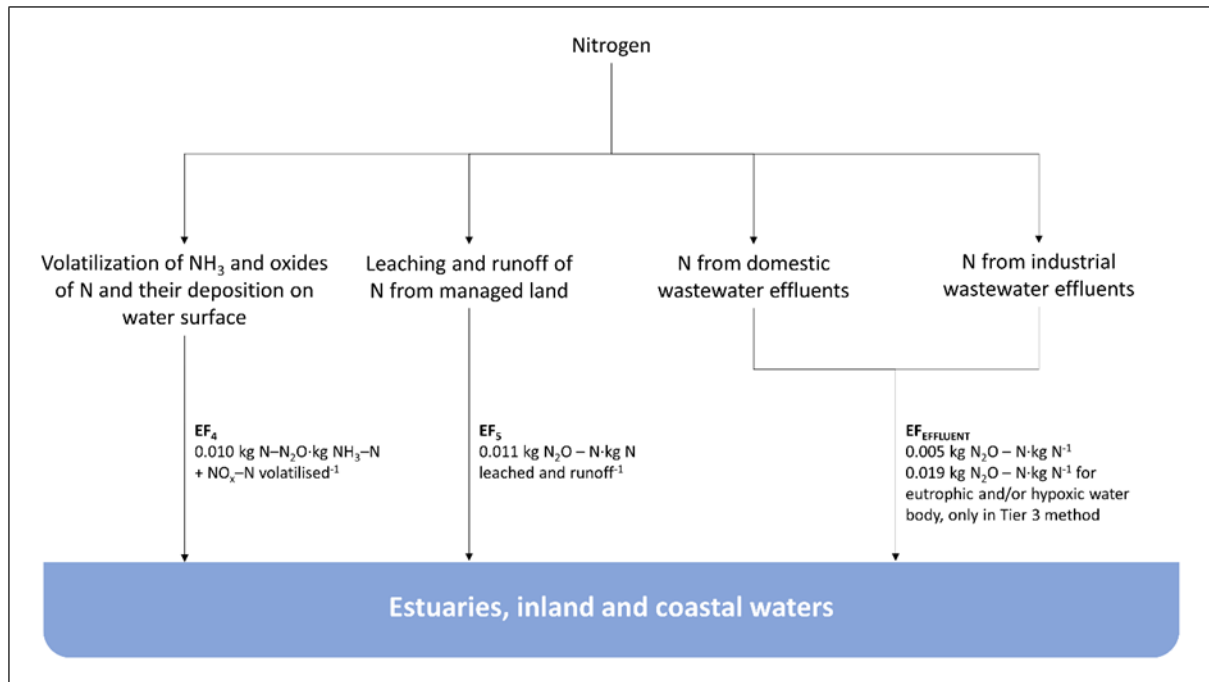


Figure 3: Current emission factors used for the different N flux causing indirect anthropogenic N<sub>2</sub>O emissions from aquatic environments.

Table 1: Default emission factor values recommended by the IPCC (values are in kg N–N<sub>2</sub>O·kg N<sup>-1</sup>).

		1996	2006	2019
Tier 1	$EF_5$	0.025	0.0075	0.011
	$EF_4$	0.01	0.01	0.01
	$EF_{\text{EFFLUENT}}$	NA	0.005	0.005
Tier 3	$EF_{\text{EFFLUENT}}$	NA	0.005	0.005 0.019 for eutrophic waters

The IPCC considers that N<sub>2</sub>O is mostly produced via microbial nitrification and denitrification processes. That assumption implies that mitigation strategies typically solely rely on reducing N inputs. However, microalgae have also been showed to produce N<sub>2</sub>O (Plouviez et al., 2019a; Weathers, 1984; Weathers & Niedzielski, 1986; Guieysse et al., 2013; Burlacot et al., 2020), although their contribution to N<sub>2</sub>O emissions from aquatic environments has barely been investigated. Besides, emissions have been recorded in oxic<sup>2</sup> environments that cannot always be attributed to the bacterial pathways (Table

<sup>2</sup> Oxygen-replete conditions

2). For example, Miao et al. (2020) observed that N<sub>2</sub>O emissions were high under normoxia<sup>3</sup>, especially when algae blooms occurred.

Table 2: N<sub>2</sub>O emissions reported during published studies acknowledging algal N<sub>2</sub>O synthesis in natural environments, simplified from Plouviez et al. (2019a).

Ecosystem	N <sub>2</sub> O fluxes	O <sub>2</sub> conditions <sup>1</sup>	Reference
Ocean	115 nmole·m <sup>-2</sup> ·h <sup>-1</sup>	Normoxic	Cohen & Gordon, 1978
Ocean	409 nmole·m <sup>-2</sup> ·h <sup>-1</sup>	Hypoxic	Pierotti & Rasmussen, 1980
Coastal wetland	125 – 228 nmole·m <sup>-2</sup> ·h <sup>-1</sup>	Anoxic and hypoxic	Smith et al., 1983
Ocean	123 – 132% saturation	Normoxic	Oudot et al., 1990
Lakes	300 – 700 nmole·m <sup>-2</sup> ·h <sup>-1</sup>	From anoxic to normoxic	Mengis, 1997
Ocean	88 nmole·m <sup>-2</sup> ·h <sup>-1</sup>	Not specified	Morell et al., 2001
Lake	357 – 2450 nmole·m <sup>-2</sup> ·h <sup>-1</sup>	Not specified	Wang et al., 2006
Lakes	0 – 10,057 nmole·m <sup>-2</sup> ·h <sup>-1</sup>	Oxic	McCrackin & Elser, 2011
Lake	12.5 – 2233 nmole·m <sup>-2</sup> ·h <sup>-1</sup>	Normoxic and Hyperoxic	Miao et al. (2020)

<sup>1</sup> Anoxic conditions are assumed when the oxygen concentration ([O<sub>2</sub>]) of the water is inferior to 3 μM, hypoxic conditions are assumed when 3 μM < [O<sub>2</sub>] < 200 μM, normoxic conditions are assumed when 200 μM < [O<sub>2</sub>] < 400 μM, and hyperoxic conditions are assumed when [O<sub>2</sub>] > 400 μM in freshwater (Diaz, 2016).

Eutrophic environments are inclined to algal blooms due to their N but mostly P supply that can trigger algae proliferation (Kalff, 1983; Nalewajko & Murphy, 2001; Elser et al., 2007; Abell et al., 2010). Consequently, evidencing the role microalgae play in N<sub>2</sub>O emissions is critical to adjust current methodologies for computing N<sub>2</sub>O emissions from aquatic ecosystems; establish the best mitigation strategies for eutrophic environments; and generate knowledge that can be applied to further research in similar ecosystems (e.g. eutrophic marine areas).

## 2.2. N<sub>2</sub>O synthesis by microorganisms

Many organisms, including phytoplankton and plants, can produce N<sub>2</sub>O via processes such as denitrification, nitrification and nitrogen uptake from nitrates (Hahn & Junge, 1977; Hallin et al., 2018). The following section reviews these mechanisms with special emphasis on microalgae.

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<sup>3</sup> Air-saturated conditions

### 2.2.1. Bacteria

N<sub>2</sub>O can be produced by bacteria through nitrification, denitrification, coupled nitrification-denitrification, nitrifier-denitrification, and anaerobic ammonia oxidation in the benthic<sup>4</sup> zone of different aquatic ecosystems.

**Nitrification:** Autotrophic ammonia-oxidizing bacteria (AOB), such as *Nitrosomonas* and *Nitrosococcus* species, use ammonia (NH<sub>3</sub>) as energy source under oxic conditions. AOB thus carry out the first step of nitrification, the aerobic oxidation of NH<sub>3</sub> into nitrite (NO<sub>2</sub><sup>-</sup>) by the enzyme ammonia monooxygenase (AMO). AMOs are copper-containing membrane-associated monooxygenases encoded by the genes *amoA*, *amoB* and *amoC* (Hayatsu et al., 2008). During nitrification, NH<sub>3</sub> is oxidised into hydroxylamine (NH<sub>2</sub>OH) by AMO and NH<sub>2</sub>OH is then oxidised to NO<sub>2</sub><sup>-</sup> by hydroxylamine oxidoreductase (HAO). During this process, N<sub>2</sub>O can be formed from the spontaneous chemical decomposition of NH<sub>2</sub>OH or NO<sub>2</sub><sup>-</sup> as illustrated in Figure 4 (Hayatsu et al., 2008; Wrage et al., 2001; Zhu-Barker et al., 2015). Nitrite-oxidizing bacteria (NOB) such as *Nitrobacter* and *Nitrospina* genera are responsible for the final step of nitrification, the oxidation of NO<sub>2</sub><sup>-</sup> into nitrate (NO<sub>3</sub><sup>-</sup>) by nitrite oxidoreductase (NXR). Microorganisms capable of oxidizing NH<sub>3</sub> all the way to NO<sub>3</sub><sup>-</sup> steps have recently been discovered and called complete ammonia oxidizer (comammox) (Daims et al., 2016). These organisms belong to the genus *Nitrospira* and can produce N<sub>2</sub>O from the abiotic conversion of NH<sub>2</sub>OH or NO<sub>2</sub><sup>-</sup> as described above (Han et al., 2021). Heterotrophic nitrifiers (e.g. *Paracoccus*, *Pseudomonas* and *Staphylococcus* genera) carry out heterotrophic ammonium oxidation to nitrate under normoxia or hypoxia, but urea can be used instead of ammonia (Wrage et al., 2001). AOB, NOB, heterotrophic nitrifiers and comammox can all be found in marine and freshwater ecosystems (Ward et al., 2011; Huang et al., 2020).

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<sup>4</sup> Located at the bottom of the water body

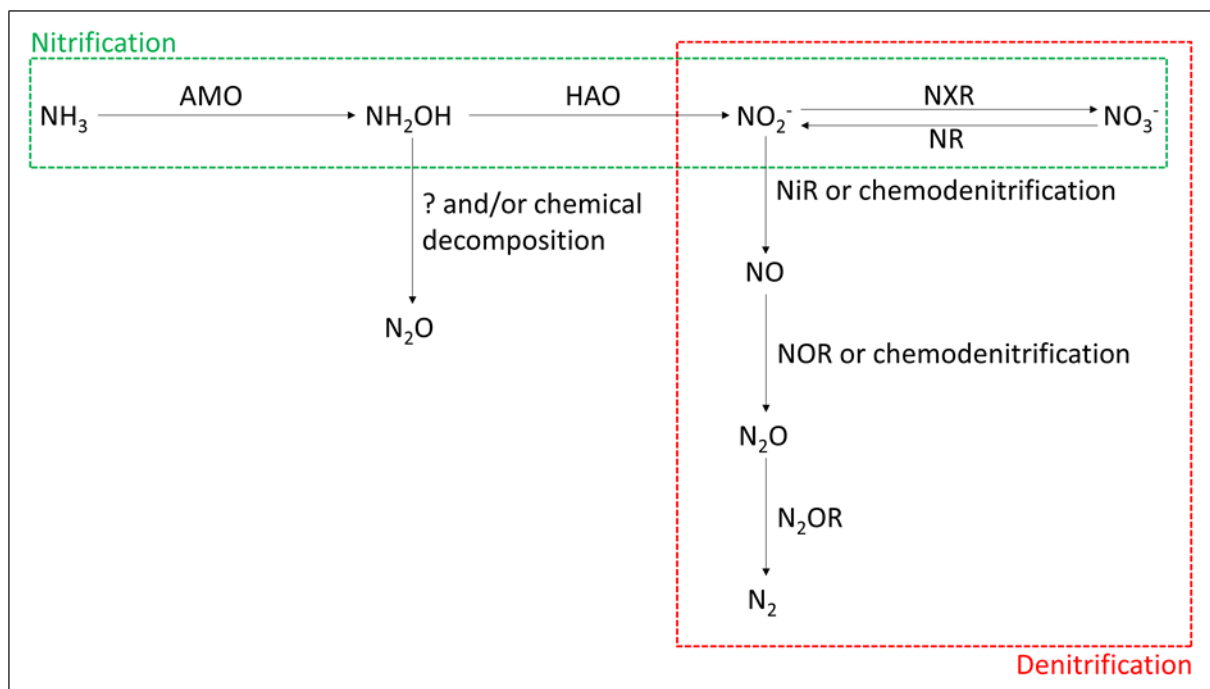


Figure 4: Nitrification and denitrification pathways in bacteria.

Abbreviations/symbols:  $\text{NH}_3$ , ammonia;  $\text{NH}_2\text{OH}$ , hydroxylamine;  $\text{NO}_2^-$ , nitrite;  $\text{NO}_3^-$ , nitrate;  $\text{NO}$ , nitric oxide;  $\text{N}_2\text{O}$ , nitrous oxide;  $\text{N}_2$ , dinitrogen; AMO, ammonia monooxygenase; HAO, hydroxylamine oxidoreductase; NXR, nitrite oxidoreductase; NR, Nitrate reductase; NiR, nitrite reductase; NOR, nitric oxide reductase;  $\text{N}_2\text{OR}$ , nitrous oxide reductase; ?, unknown enzyme.

**Denitrification:** Denitrifying bacteria (e.g. *Pseudomonas* and *Alcaligenes* genera) are responsible for  $\text{N}_2\text{O}$  production during partial denitrification under hypoxia or anoxia. Denitrification is a respiratory pathway during which  $\text{NO}_3^-$  is first reduced to  $\text{NO}_2^-$  by nitrate reductase (NR) and  $\text{NO}_2^-$  is then reduced to nitric oxide (NO) by nitrite reductase (NiR). NO is in turn reduced to  $\text{N}_2\text{O}$  by nitric oxide reductase (NOR) and  $\text{N}_2\text{O}$  is finally reduced to nitrogen ( $\text{N}_2$ ) by nitrous oxide reductase (symbolized as either  $\text{N}_2\text{OR}$  or NOS in the literature) as illustrated in Figure 4. The genes *narG*, *nirS* and *nirK*, *norB* and *norZ*, and *nosZ* encode NR, NiR, NOR and  $\text{N}_2\text{OR}$ , respectively (Garrido-Amador et al., 2021; Ren et al., 2019). During denitrification,  $\text{N}_2\text{O}$  production increases with the presence of  $\text{O}_2$  and at low pH as these conditions inhibit  $\text{N}_2\text{OR}$  activity. Denitrification is a widespread biological process that occur in most aquatic environments (Piña-Ochoa & Álvarez-Cobelas, 2006).

**Nitrifier-denitrification:** Certain AOB and methanotrophs can oxidise  $\text{NH}_3$  until  $\text{NO}_2^-$  and then denitrify  $\text{NO}_2^-$  to  $\text{N}_2$  via  $\text{N}_2\text{O}$  (as described above). Such coupled nitrification-denitrification process may also involve bacterial consortia where the nitrification and denitrification steps are carried out by different bacteria (Wrage et al., 2001). During nitrifier-denitrification under hypoxia,  $\text{N}_2\text{O}$  is released as an intermediate of nitrite reduction as in the case of 'classical' denitrification.

**Anaerobic ammonia oxidation:** Under anoxia, anammox bacteria oxidize ammonia (as energy source) using nitrite as electron acceptor. In the presence of  $\text{NH}_4^+$  and  $\text{NO}_2^-$ ,  $\text{N}_2\text{O}$  emissions have been

reported during anammox-based wastewater treatment but the pathway of N<sub>2</sub>O synthesis by these bacteria has not yet been identified (Kampschreur et al., 2009; Ali et al., 2016). Anammox bacteria can be found in the anoxic<sup>5</sup> and hypoxic zones of freshwater and marine environments<sup>6</sup> (Ward et al., 2011). These bacteria have so far only been reported in one lake and, therefore, their significance to N<sub>2</sub>O emissions in freshwater ecosystems is still poorly understood (Ward et al., 2011).

### 2.2.2. Fungi

Fungi such as *Fusarium oxysporum* can produce N<sub>2</sub>O via denitrification and co-denitrification (Figure 5). The fungal denitrification pathway takes place under either aerobic and anaerobic conditions and is similar to the bacterial denitrification pathway described above. The fungal *nirK* gene encoding its NirK enzyme is an orthologue<sup>7</sup> of the bacterial *nirK* gene. The fungal NOR enzyme, called cytochrome P450 nitric oxide reductase (P450nor), is encoded by the gene *p450nor* and is a member of the CYP55 family of the P450 cytochrome superfamily. Therefore, fungi can also produce N<sub>2</sub>O from NO<sub>2</sub><sup>-</sup> and a nitrogen co-substrate (e.g. NH<sub>4</sub><sup>+</sup>, amino acids, urea) in a process catalysed by the fungal NOR and known as co-denitrification (Hayatsu et al., 2008; Shoun et al., 2012; Higgins et al., 2016). Fungi are present in lentic<sup>8</sup> bodies' sediments and N<sub>2</sub>O emissions from Poyang Lake, China, have been correlated with fungal biomass concentration (Liu & Xu, 2016).

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<sup>5</sup> Severe oxygen depleted conditions

<sup>6</sup> i.e. sediments and oceanic OMZs

<sup>7</sup> An orthologous gene is a gene inherited in two species by a common ancestor and which diverged after the creation of the two species. The gene and its main function are conserved despite the divergence.

<sup>8</sup> Body of standing water

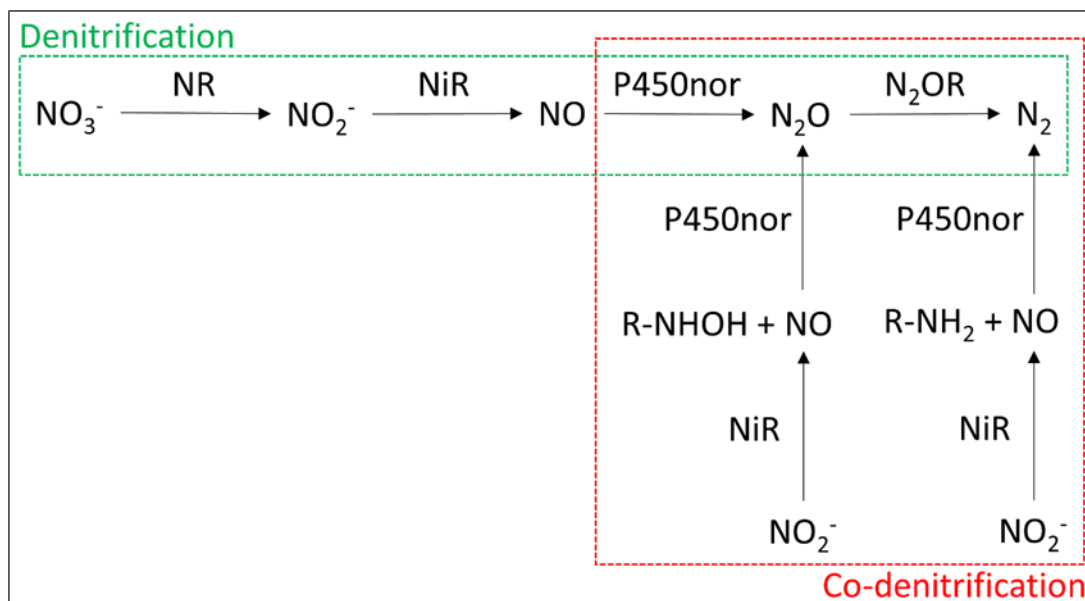


Figure 5: Denitrification and co-denitrification pathways in fungi.

Abbreviations/symbols:  $\text{NO}_3^-$ , nitrate;  $\text{NO}_2^-$ , nitrite; NO, nitric oxide;  $\text{N}_2\text{O}$ , nitrous oxide;  $\text{N}_2$ , dinitrogen; R-NHOH, imine; R-NH<sub>2</sub>, amine; NR, Nitrate reductase; NiR, nitrite reductase; P450nor, cytochrome P450 nitric oxide reductase;  $\text{N}_2\text{OR}$ , nitrous oxide reductase.

### 2.2.3. Archaea

Certain ammonia-oxidizing Archaea (AOA) harbour *amo*-like genes, *amoA*, *amoB*, and *amoC* encoding the three subunits of the AMO enzyme. AOA are thus capable of catalysing the first step of nitrification following the same pathway as AOB (Hatzenpichler, 2012). AOA were found to significantly contribute to  $\text{N}_2\text{O}$  emissions from oceans (Santoro et al., 2011; Löscher et al., 2012; Qin et al., 2017). While the pathways of archaeal  $\text{N}_2\text{O}$  synthesis are still debated (Wu et al. 2020; Stein et al. 2021), Jung et al. (2019) recently suggested that a strain of AOA could synthesise  $\text{N}_2\text{O}$  enzymatically at low pH in a reaction involving NO reduction into  $\text{N}_2\text{O}$  by the cytochrome P450. Some archaea are also capable of denitrification via a pathway similar to bacteria, and this mechanism could be a potential  $\text{N}_2\text{O}$  source (Hayatsu et al., 2008). AOA are generally found in marine waters but their presence has been reported in lakes as well (Wenk et al., 2016; Prosser et al., 2020).

### 2.2.4. Microalgae

Photosynthetic microorganisms include prokaryotes (e.g. cyanobacteria) and eukaryotes (e.g. Chlorophyceae, diatoms, etc.), including unicellular algae (existing as individual cells or in groups) known as microalgae. While microalgae are taxonomically all eukaryotic phototrophs, cyanobacteria are often considered as microalgae members in the engineering literature: This broader definition will

be used in this review for simplicity. Microalgae are mostly found in aquatic environments (Borowitzka, 2012), and their contributions to  $N_2O$  emissions have mainly been confirmed and reported in the last decade (Plouviez et al., 2019a).

While the study of  $N_2O$  synthesis in microalgae is still at its infancy, various pathways have been proposed in the literature, especially in the green unicellular microalga *Chlamydomonas*, as summarized in Figure 6. Succinctly, microalgae  $N_2O$  synthesis is currently believed to occur via the successive reduction of nitrate into nitrite and NO. NO is thus seen as the substrate of  $N_2O$  via various putative pathways as further detailed below.

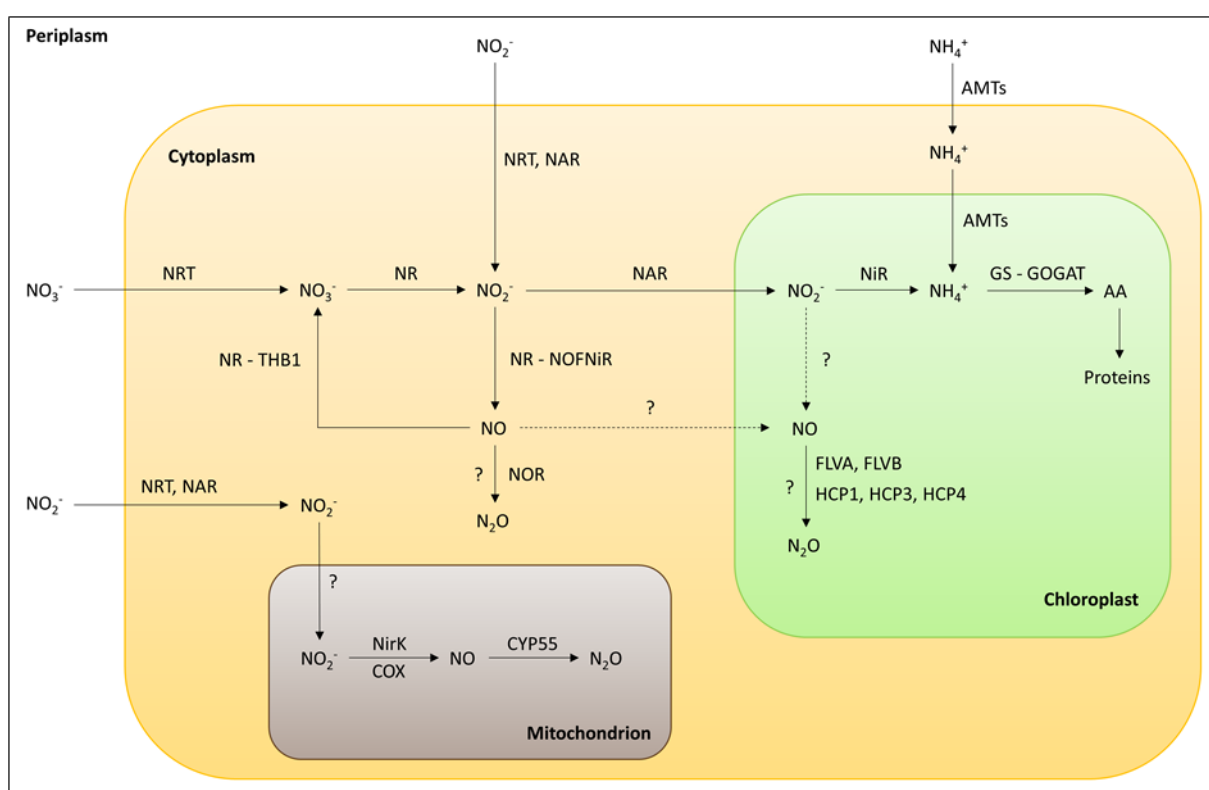


Figure 6:  $N_2O$  putative production pathways in *Chlamydomonas*.

Abbreviations/symbols:  $NH_4^+$ , ammonium;  $NO_2^-$ , nitrite;  $NO_3^-$ , nitrate; NO, nitric oxide;  $N_2O$ , nitrous oxide; AA, Amino acids; NRT, High-affinity nitrate transporter; NAR, Nitrite transporter; AMT, High-affinity ammonium transporter; NR, Nitrate reductase; NiR, Nitrite reductase; GS-GOGAT, Glutamate synthase - Glutamine oxoglutarate aminotransferase; NOFNiR, NO-forming nitrite reductase; THB1, Truncated hemoglobin; NirK, copper-containing nitrite reductase; COX, Cytochrome c oxidase; CYP55, Cytochrome P450 (Nitric oxide reductase); NOR, Nitric oxide reductase; FLV, Flavodiiron protein; HCP, Hybrid cluster protein.

#### 2.2.4.1. NO synthesis

NO is an important signal molecule in plants and algae. It is also involved in nitrate assimilation and can be synthesized by the enzymes NirK and NR in *Chlamydomonas* (Sanz-Luque et al., 2015a). NR is



encoded by the gene *NIA1* in *Chlamydomonas* and is made of 2 identical subunits each containing three prosthetic groups: flavin adenine dinucleotide (FAD),  $b_{557}$  heme, and molybdenum cofactor (Sanz-Luque et al., 2015a). NR expression is repressed by  $\text{NH}_4^+$  but induced by light and  $\text{NO}_3^-$ . NR plays a crucial role in the intracellular  $\text{NO}_3^-/\text{NO}_2^-$  cycle as its activity changes depending on its partner protein (Figure 6): When associated with NOFNiR (NO-forming nitrite reductase), the NR-NOFNiR dual enzyme reduces  $\text{NO}_2^-$  into NO (Chamizo-Ampudia et al., 2016). In association with THB1 (Truncated Hemoglobin 1), the NR-THB1 dual enzyme catalyses NO oxidation into  $\text{NO}_3^-$  (Sanz-Luque et al., 2015b; Bellido-Pedraza et al., 2020).

NirK is not broadly conserved in eukaryotes (Kim et al., 2009) but *Chlamydomonas* harbours a gene homologous to the bacterial and fungal genes encoding the enzyme NirK involved in reducing  $\text{NO}_2^-$  into NO (Figure 7). The function of the *Chlamydomonas* enzyme, called CrNiRK, has not been described yet but its presence suggests a pathway for NO synthesis from  $\text{NO}_2^-$  in the mitochondria (Figure 6). NirK is a copper-containing nitrite reductase composed of three subunits (Zumft, 1997) and NirK proteins are divided into two classes. CrNiRK belongs to the class II within which all eukaryotic and some prokaryotic NirK proteins are gathered. Kim et al. (2009) suggested that eukaryotic *NirK* homologs have a common origin: the protomitochondrion<sup>9</sup>.

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<sup>9</sup> The ancestral bacterial endosymbiont from which all mitochondria are thought to be derived

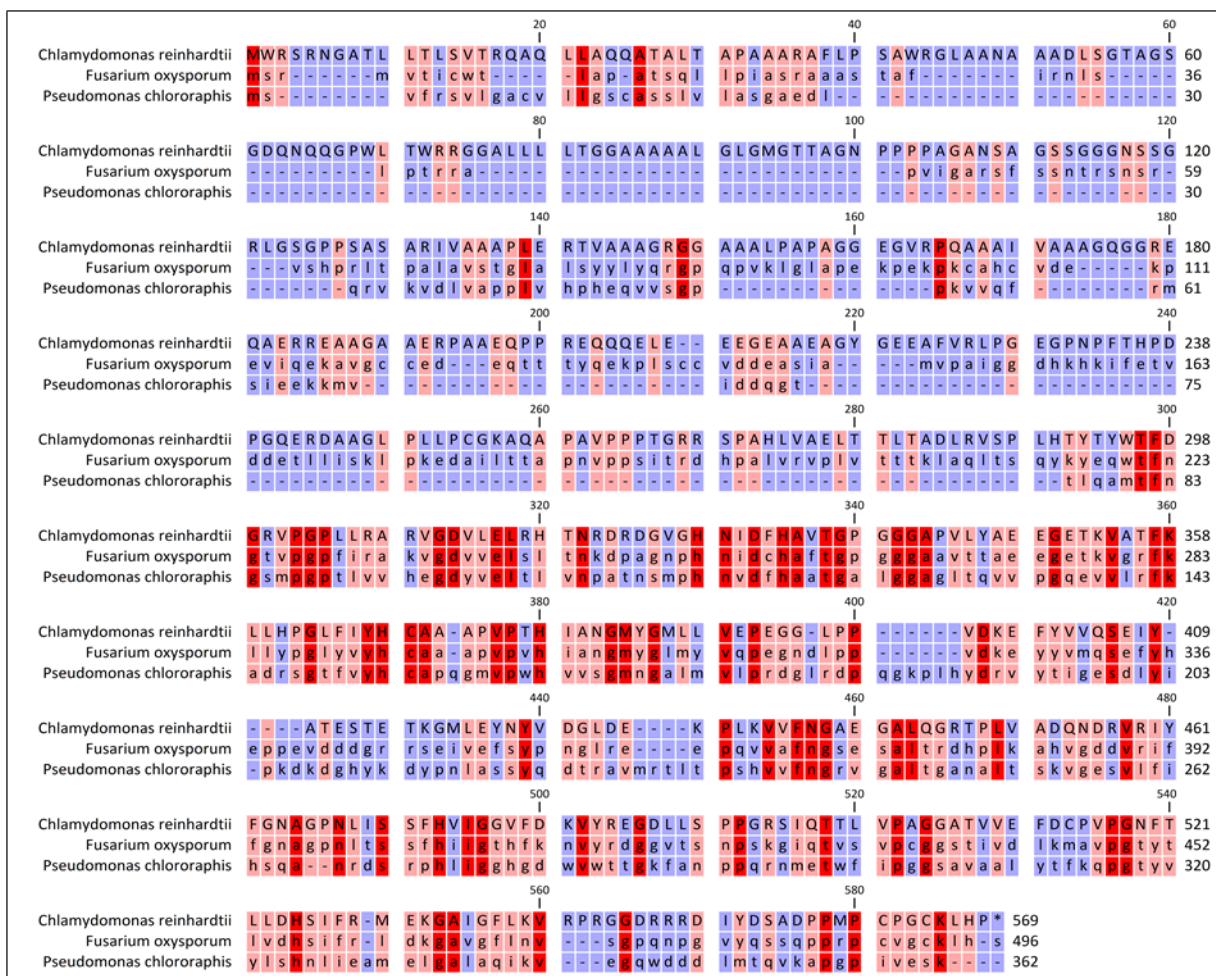


Figure 7: Peptides alignment of Nirk proteins from *Chlamydomonas reinhardtii* (microalga), *Fusarium oxysporum* (fungus) and *Pseudomonas chlororaphis* (bacterium) generated using the software CLC Sequence Viewer.

### 2.2.4.2. FLVA and FLVB

Flavodiiron proteins (FDPs) are commonly found in bacteria, archaea and many eukaryotes. Here, we focus on the class-C flavodiiron proteins which are found in photosynthetic organisms and are named FLVs (Alboresi et al., 2019). In prokaryotes, FLV2 and FLV4 are unique to  $\beta$ -cyanobacteria while FLV1 and FLV3 are present in all cyanobacteria (Alboresi et al., 2019). *Chlamydomonas* has two FLVs, named FLVA and FLVB, which are structurally similar to the cyanobacterial proteins. A nitric oxide and oxygen reductase activity has been reported for FLVs. In cyanobacteria, FLV1 and FLV3 have an electron sink role by reducing oxygen into water under light conditions. This reaction protects the photosystem I (PSI) during light fluctuation (Chaux et al., 2017; Jokel et al., 2015, 2018). In *Chlamydomonas*, O<sub>2</sub> reduction into H<sub>2</sub>O by FLVs constitutes an alternative electron transport pathway which limits the overproduction of cell-damaging redox power and maintains ATP synthesis immediately after an increase in illumination and under constant illumination (Alboresi et al., 2019; Saroussi et al., 2019).

In *Chlamydomonas*, FLVs have recently been suggested to reduce NO into N<sub>2</sub>O using electrons from photosynthesis under illumination (Burlacot et al., 2020). Additionally, FLVs have been reported to play a role during sulfur deprivation in *Chlamydomonas* as their activity increased when the cells were switching from aerobiosis<sup>10</sup> to anaerobiosis<sup>11</sup> to produce ATP without generating reducing power (Alboresi et al., 2019).

#### 2.2.4.3. CYP55

*Chlamydomonas* harbours a gene analogous of the fungal *CYP55* gene encoding a nitric oxide reductase (NOR) and possibly localized in the mitochondria (Plouviez et al., 2017a). *CYP55* is upregulated in presence of nitrate (Gérin et al., 2010), and the use of *CYP55*-deficient mutants appears to indicate that *CYP55* is involved in NO reduction into N<sub>2</sub>O in the dark, as illustrated in Figure 6 (Burlacot et al., 2020; Plouviez et al., 2017a). Alternatively, Burlacot et al. (2020) proposed that *CYP55* may be located in the chloroplast although reducing power can also be transferred from chloroplasts to mitochondria where mitochondrial proteins (like the alternative terminal respiratory oxidase AOX1) can help alleviate light energy excess (Kaye et al., 2019).

#### 2.2.4.4. HCP

*Chlamydomonas* also possess four Hybrid Cluster Proteins (HCPs) found in a large range of prokaryote and eukaryote organisms (Almeida et al., 2006). These enzymes are metalloproteins characterized by the presence of an iron-sulfur-oxygen cluster and, while their physiological functions are still uncertain in *Chlamydomonas*, different roles have been proposed based on bacterial HCPs. In *Escherichia coli*, these proteins have a peroxidase activity (Almeida et al., 2006) and also seem to be responsible for the production of N<sub>2</sub>O in anaerobic conditions, as shown in studies using HCP-deficient *E. coli* mutants (Wang et al., 2016). Furthermore, Van Lis et al. (2020) suggested that HCPs might exhibit a nitrate-reducing S-nitrosylase activity in *Chlamydomonas* that would catalyse NO<sub>3</sub><sup>-</sup> transformation into nitrosonium (NO<sup>+</sup>), but these authors argued that a NO-reductase role of HCP is unlikely in *Chlamydomonas*. The roles of these proteins in eukaryotic microalgae are still poorly understood and require further investigation.

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<sup>10</sup> Growth in an environment containing oxygen.

<sup>11</sup> Growth in an environment devoid of oxygen.

Van Lis et al. (2020) hypothesised that *HCP* genes in *Chlamydomonas* appeared via duplication events. HCP1, 3 and 4 have been predicted to be found in chloroplasts while HCP2 was located in mitochondria. In *Chlamydomonas*, *HCP* transcripts were reported to be highly upregulated under dark anoxia conditions but not translated into proteins (van Lis et al., 2020). Although no physiological explanations have been formulated for this lack of translation, it could indicate that HCPs are regulated at different transcriptional and translational levels to prime cells before exposure to conditions when HCPs are required. The mRNA generated during HCPs transcription was also reported to accumulate in the presence of  $\text{NO}_3^-$  (Gérin et al., 2010).

#### 2.2.4.5. Other mechanisms

$\text{N}_2\text{O}$  emissions from several microalgae species have been reported in the literature (Table 3) but, to our knowledge, the molecular mechanisms involved have not been further studied for these strains. Bellido-Pedraza et al. (2020) suggested that almost 33% of microalgae contain at least one of the proteins involved in  $\text{N}_2\text{O}$  synthesis in *Chlamydomonas* (NirK, CYP55, FLVs, HCP).

Table 3: Algal species involved in  $\text{N}_2\text{O}$  emissions reported during published studies (modified from Plouviez et al. (2019a)).

Alga division	Algal species	$\text{N}_2\text{O}$ flux	Reference
Green algae	<i>Chlorella vulgaris</i>	109 – 1480 $\text{nmole}\cdot\text{h}^{-1}\cdot\text{g DW}^{-1}$	Guieysse et al. (2013)
	<i>Chlorella rubescens</i>	1,200 – 2,500 $\text{nmole}\cdot\text{h}^{-1}\cdot\text{g DW}^{-1}$	Weathers (1984)
	<i>Coelastrum sp.</i>	560 – 1,100 $\text{nmole}\cdot\text{h}^{-1}\cdot\text{g DW}^{-1}$	Weathers (1984)
	<i>Chlorococcum vacuolarum</i>	150 – 290 $\text{nmole}\cdot\text{h}^{-1}\cdot\text{g DW}^{-1}$	Weathers (1984)
	<i>Neochloris sp.</i>	9.60 – 38,000 $\text{nmole}\cdot\text{m}^{-2}\cdot\text{h}^{-1}$	Plouviez et al. (2017b)
	<i>Scenedesmus dimorphus</i>	6 – 73 $\text{nmole}\cdot\text{h}^{-1}\cdot\text{g DW}^{-1}$	Bauer et al. (2016)
	<i>Scenedesmus obliquus</i>	0 – 1,000 $\text{nmole}\cdot\text{h}^{-1}\cdot\text{g DW}^{-1}$	Weathers (1984)
	Diatoms	<i>Skeletonema marinoi</i>	0.039 – 0.31 $\text{nmole}\cdot\text{h}^{-1}\cdot\text{aggregate}^{-1}$
<i>Thalassiosira weissflogii</i>		0.087 – 0.3 $\text{nmole}\cdot\text{L}^{-1}\cdot\text{h}^{-1}$	Kamp et al. (2013)
Cyanobacteria	<i>Aphanocapsa 6308</i>	0 – 1,500 $\text{nmole}\cdot\text{h}^{-1}\cdot\text{g DW}^{-1}$	Weathers & Niedzielski (1986)
	<i>Aphanocapsa 6714</i>	0 – 5,700 $\text{nmole}\cdot\text{h}^{-1}\cdot\text{g DW}^{-1}$	Weathers & Niedzielski (1986)
	<i>Nostoc sp.</i>	0 – 1,500 $\text{nmole}\cdot\text{h}^{-1}\cdot\text{g DW}^{-1}$	Weathers & Niedzielski (1986)

To summarize the current state of the art on biological  $\text{N}_2\text{O}$  synthesis, nitrification and denitrification are the two major biological pathways involved in  $\text{N}_2\text{O}$  production. These processes have traditionally been considered to be undertaken by bacteria and archaea. However, increasing evidence suggests that other organisms, such as microalgae, might be able to carry out at least part of these processes

and be responsible for N<sub>2</sub>O emissions from aquatic environments. The lack of knowledge on the pathways involved in microalgal N<sub>2</sub>O synthesis could mean that current N<sub>2</sub>O inventories are missing or misallocating significant N<sub>2</sub>O sources. Studying algal N<sub>2</sub>O emissions from eutrophic ecosystems is especially relevant now that the IPCC considers using a new higher EF to estimate N<sub>2</sub>O emissions from eutrophic waters receiving N from wastewater effluents. Indeed, the ability of microalgae to synthesise N<sub>2</sub>O combined with their ubiquity in eutrophic ecosystems suggest that microalgae play a significant role in these N<sub>2</sub>O emissions. Besides, as wastewater effluents are commonly discharged in lakes prone to eutrophication, lakes may thus constitute potential significant N<sub>2</sub>O sources. Finally, N<sub>2</sub>O emissions from terrestrial plants have been reported during NO<sub>3</sub><sup>-</sup> assimilation (Dean & Harper, 1986; Hakata et al., 2003; Smart & Bloom, 2001; Timilsina et al., 2020a; Timilsina et al., 2020b), which means the study of N<sub>2</sub>O production pathways in microalgae, a higher plant ancestor, could improve our understanding of the mechanisms involved in plants.

## 2.3. Diversity of freshwater lakes

### 2.3.1. Dimension/Morphology

Freshwater lakes, rivers and reservoirs hold 0.26% of the liquid freshwater found on Earth. Freshwater lakes cover 2.7 million km<sup>2</sup> and hold 175,000 km<sup>3</sup> of water (Carpenter et al., 2011). Natural lakes are formed during events such as volcanic eruption, glacier retreat, river movement, landslides, tectonic movement, shorelines movement and the action of wind on land, and solution lakes are a specific type of natural lakes formed by dissolution and collapse of rock structures. Artificial lakes, called reservoirs or impoundments, are typically created following the damming of a river. How a lake is formed establishes the lake's morphology, especially water depth and the hydraulic retention time (HRT) of the water flowing through the lake. The depth and HRT in turn impact the biological productivity of the lake (Wetzel, 2001a) through the interactions between the water, nutrient-laden sediments, and microorganisms (Wetzel, 2001a). Generally, biological productivity increases with the HRT and the area of sediments in contact with water.

A distinction is often made between deep (> 10 m) and shallow (< 10 m) lakes in the literature: Around 40% of the volume of freshwater held in freshwater lakes is found in great lakes basins (i.e. deep) but most lakes are small and shallow (Wetzel, 2001a; Meybeck, 1995; Downing et al., 2006). Shallow lakes are characterized by the absence of sustained thermal stratification (Kar, 2013). Compared to deep lakes, shallow lakes are also characterized by a relatively low HRT, low nutrient loss to the sediments, and high sediments resuspension due to the influence of wind (Wetzel, 2001b). These characteristics lead to a high nutrient load to the aqueous phase causing shallow lakes to be more prone to eutrophication than deep lakes. As further discussed below, lakes can be classified based on different characteristics and the trophic state is often used when considering water quality (see Appendix D for further details).

### 2.3.2. Trophic states

Carlson (1977) developed the Trophic State Index (TSI) to facilitate communication with the public and classify lakes based on their trophic state. This index is based on the chlorophyll a concentration (Chl<sub>a</sub>), the Secchi depth (SD), and the total phosphorus (TP) concentration of the lake. The Chl<sub>a</sub> concentration provides an indication of the algal biomass concentration in the water and is often correlated to TP concentration when phosphorus is the limiting nutrient for algae growth (Elser et al., 2007). The Secchi depth is used to estimate water transparency by measuring the depth below which a disc of 20 cm in diameter vanishes from sight when immersed in the water. This variable is also related to the algal

biomass concentration as it defines the depth of the layer inside which photosynthesis can be performed by primary producers. In New Zealand, a similar index ranging from 0 to 7 called the Trophic Level Index (TLI) is used to better fit the lake conditions of the country as explained in Appendix E. In contrast to the TSI, the TLI includes the total nitrogen (TN) concentration in its calculations because nitrogen has also been reported to be limiting algae growth (Abell et al., 2020; White et al., 1985). The different lake trophic states according to the TLI score are described in Table 4. The trophic state of a water body can give an idea of the water quality for different water uses.

Table 4: Description of the different trophic states of a lake according to its TLI score (Lake Trophic Level Index, 2020).

Trophic states	Description	TLI Score
Microtrophic	Very low levels of nutrients and algae, the water is transparent.	0 – 2
Oligotrophic	Water is transparent, oxygen is available in all layers, the nutrient level is low, and the water is poor in phytoplankton.	2 – 3
Mesotrophic	The water has moderate levels of nutrient and algae.	3 – 4
Eutrophic	The water is less transparent, rich in nutrients and the sediments are rich in organic matter. Hypolimnial oxygen depletion and frequent algae bloom can occur during summer.	4 – 5
Supertrophic	The water is cloudy, very rich in nutrients, especially nitrogen and phosphorus, with a high risk of algae blooms.	5 – 7

## 2.4. Eutrophication

Eutrophication is a natural phenomenon characterized by the overgrowth of primary producers (e.g. microalgae including the prokaryotic cyanobacteria, macroalgae etc.) in an aquatic ecosystem. Over geological timescales, eutrophication can lead to the filling, and ultimately the disappearance, of the water basin. Anthropogenic activities have accelerated the rate of eutrophication due to excess N and P inputs to aquatic ecosystems leading to many environmental issues as further discussed below (Pinay et al., 2018). Natural lakes, impoundments and ponds cover more than 3% of the Earth's surface (Downing et al., 2006) and as many of 30 - 40% of these aquatic bodies are eutrophic (Yang et al., 2008). Thus, it has been estimated that 0.9 - 1.2 % of the Earth's surface is covered with eutrophic lakes, which is equivalent to 5.36 million km<sup>2</sup> or 20 times the area of New Zealand.

### 2.4.1. Parameters leading to eutrophication

Eutrophication is characterized by the excessive growth of primary producers. Primary producers are photosynthetic organisms using light energy to convert CO<sub>2</sub> into organic matter via photosynthesis. Their growth can be promoted by different parameters, as described below.

**Nutrients:** Eutrophication is generally linked to the excessive input of N and/or P into aquatic bodies. These nutrients can originate from point and nonpoint sources. Point sources are usually continuous and relatively easy to identify and include wastewater effluents, runoffs from waste disposal or construction sites, storm overflows and sanitary sewer outfalls. Nonpoint sources are harder to identify because of their diffuse nature and include agriculture runoffs, urban runoffs from unsewered areas, septic tank leachate, atmospheric deposition caused by fossil fuel combustion, and land activities (Carpenter et al., 1998). Carpenter et al. (1998) estimated that most of the nutrients inputs leading to eutrophication originate from nonpoint sources in the USA (*1998 National Water Quality Inventory Report to Congress, 1998*; Carpenter et al., 1998). Beusen et al. (2016) estimated that the global contribution of agriculture to N and P inputs to freshwater has increased from 19% to 51% and 35% to 56%, respectively between 1900 and 2000. As explained above, nutrient availability is higher in shallow lakes than in deep lakes, which means nutrient concentration is often non-limiting in shallow lakes. Consequently, phytoplankton growth in shallow lakes will be mainly regulated by the four factors described below (Wetzel, 2001b).

**Light:** Wind and water flow (i.e. turbulence) resuspend particles (e.g. silt, clay, microorganisms) that reduce light penetration in water. Thus, the growth of benthic<sup>4</sup> primary producers is limited to low-depth area with clear water and low turbulence (Pinay et al., 2018). In many lakes, photosynthetic activity is mainly located near the surface as the primary producers themselves shade the water column below.

**Temperature:** Temperature (>15°C) accelerates biological reactions and promotes microbial growth (Paerl et al., 2001). For example, Wiedner et al. (2007) observed that water temperature rise promoted the spread of the invasive tropical freshwater cyanobacterium *Cylindrospermopsis raciborskii* in lakes Langer See and Melangsee (Germany). Temperature also reduces the solubility of dissolved oxygen in the water, thereby increasing the symptoms of eutrophication during spring and summer, when temperatures are higher (Pinay et al., 2018).

**Water retention and water mixing:** A longer HRT means that phytoplankton suspended in water stays longer in the aquatic ecosystem before being flushed out. Thus, a longer HRT can favour nutrient exchanges between phytoplankton, sediments, and water as more time is given for these exchanges



to occur (Pinay et al., 2018). The occurrence of algal blooms can also increase in poorly mixed water and during sustained water stratification. When turbulence becomes excessive, phytoplankton growth is reduced or even stopped due to cells damage caused by shear. Turbulence can also disrupt interactions between bacteria and cyanobacteria (Paerl et al., 2001). Thermal stratification increases the risk of algal bloom because phytoplankton can be trapped in a nutrient-rich layer near the surface, where light is also available (Pinay et al., 2018).

**O<sub>2</sub> concentration:** internal eutrophication occurs when lakes become hypoxic. In aerobic conditions, the decomposition of organic matter involves redox reactions where oxygen is usually the final electron acceptor. However, in hypoxic or anoxic conditions, chemical species such as nitrate, iron (hydr)oxides and sulphate can act as final electron acceptors. As P often forms stable complexes with iron (hydr)oxides in sediments, the reduction of iron (hydr)oxides can cause the release of P to the water (Smolders et al., 2006).

#### **2.4.2. Environmental, societal, and political implications of eutrophication**

The proliferation of microalgae has critical ecological implications: Blooms increase water turbidity and, consequently, attenuate light penetration in the water column. This leads to a switch to respiratory processes leading to hypoxia/anoxia, which is accelerated by the decay of algal organic matter. Low oxygen concentration in turn impacts the biological community structure and, thus, the trophic network of the ecosystems (Glibert, 2017). For example, Roberts et al. (2009) noticed that the seasonal hypolimnetic hypoxia of Lake Erie (USA) impacted yellow perch distribution and foraging. Eutrophication can therefore have profound and long-term effects on an ecosystem (Pinay et al., 2018).

Eutrophication has also several economic and political implications. The proliferation of microalgae, and especially toxic microalgae as harmful algal bloom (HAB), can indeed impact the economy via restrictions for fisheries, industry intake, and recreational water use. HABs are a threat to human and animal health due to the toxins released by microalgae such as *Microcystis*, as it happened in Lake Rotoiti (New Zealand) in 2004 (Wood et al., 2006). When a drinking water supply is affected, the local authorities can advise the population not to drink water from the tap for safety reasons, as it happened in Toledo (Ohio) during summer 2014 (Jetoo et al., 2015).

Most governments have adopted regulations that seek to prevent or mitigate eutrophication, such as the international Convention for the Protection of the Marine Environment of the North-East Atlantic in 1998 (*Convention for the protection of the marine environment of the North-East Atlantic*, 1998)

and the National Policy Statement for Freshwater Management in New Zealand in 2020 (*National Policy Statement for Freshwater Management*, 2020). Increasing public awareness is also putting pressure on regional councils in charge of water management (Pham et al., 2019).

### **2.4.3. Water quality control**

Monitoring lakes is of the utmost importance to ensure the safety of water users. Besides, lakes are typically found in the upstream part of the freshwater – marine continuum, meaning that an excess of nutrients or an algal bloom in a lake can propagate to rivers, estuaries and marine environments located downstream (Aubriot et al., 2020). Monitoring reservoirs is also important to efficiently manage drinking water treatment.

#### **2.4.3.1. Parameters monitored**

In recreational water, monitoring is mostly done during the swimming season (e.g. from November/December to February/March in New Zealand). Several parameters are regularly measured to evaluate water quality such as the water pH, turbidity, and the presence of pathogens (Almeida et al., 2012). The guidelines and standards used to assess microbiological water quality depend on the water use and are country-specific (see Table 5 and Table 6 for New Zealand guidelines).

The World Health Organization (WHO) recommends to monitor the presence of HAB in freshwater by i) assessing the carrying capacity of the ecosystem for cyanobacteria; ii) visually inspecting sites to detect any scum development; and iii) quantifying the algal biomass to evaluate the bloom risk (Bartram & Rees, 1999). The first step involves the quantification of the concentration of TP and TN (as these support algae growth) and, sometimes, the chemical oxygen demand (COD) (Almeida et al., 2012). The second step involves the measurement of the Secchi depth and the search for any color irregularity and scums on the water surface. Biomass quantification is done either via chlorophyll a concentration measurement or cell count under the microscope.

Considerable research efforts have aimed at developing models that predict the occurrence of algae blooms and simulate the effects of physical (light, temperature, turbulence, HRT), chemical (nutrients), biological (microbial interactions), and ecological (food network) parameters on water quality (Janssen et al., 2019). These models could be used later in this project to make predictions on water quality and, thus, algae proliferation in water bodies.

Table 5: New Zealand guidelines for cyanobacteria in recreational fresh waters (Wood et al., 2009)

Mode	Trigger level for potentially toxic cyanobacteria		Management response
	Benthic	Planktonic	
Surveillance	Up to 20% coverage of potentially toxigenic cyanobacteria attached to substrate.	Total cyanobacteria cell count < 500 cells/mL; or The biovolume equivalent for the combined total of all cyanobacteria does not exceed 0.5 mm <sup>3</sup> /L.	Routine monitoring.
Alert	20–50% coverage of potentially toxigenic cyanobacteria attached to substrate.	Biovolume equivalent of potentially toxic cyanobacteria between 0.5 and 1.8 mm <sup>3</sup> /L; or Total biovolume of all cyanobacterial material between 0.5 and 10 mm <sup>3</sup> /L.	Notify public, erect signs with information on appearance of mats and potential risks. Increase monitoring and consider toxin testing.
Action	Greater than 50% coverage of potentially toxigenic cyanobacteria attached to substrate; or Up to 50% where potentially toxigenic cyanobacteria are visibly detaching from the substrate, accumulating as scums along the river's edge or becoming exposed on the river's edge as the river level drops.	More than 12 µg/L total microcystins; or Biovolume equivalent greater than 1.8 mm <sup>3</sup> /L of potentially toxic cyanobacteria; or Total biovolume of all cyanobacterial material greater than 10 mm <sup>3</sup> /L; or Cyanobacterial scums consistently present.	Notify public, erect signs with information on appearance of mats and potential risks, notify the public of potential risk to health. Increase monitoring and consider toxin testing.

Table 6: Microbiological water quality guidelines for marine and freshwater recreational areas (New Zealand &amp; Ministry for the Environment, 2003)

Mode	Trigger level for faecal indicator bacteria (per 100 mL)		Management response
	Beach	River/Lake	
Surveillance	Enterococci ≤ 140	<i>E. coli</i> ≤ 260	Routine monitoring
Alert	140 < Enterococci < 280	260 < <i>E. coli</i> < 550	Increase monitoring and investigation of contaminant source
Action	Enterococci > 280	<i>E. coli</i> > 550	Public warnings if required, increased monitoring and investigation of contaminant source

### 2.4.3.2. Limiting nutrients in lakes

Phosphorus availability can limit microalgae growth in lakes. Consequently, it has been suggested that many algal blooms can be prevented by reducing phosphate inputs. Schindler et al. (2008) studied the trophic state of oligotrophic Lake 227 (Canada) under 3 different nutrient loading regimes: 1) N and P supply with a high N:P ratio (around 12) for 6 years; 2) N and P supply with a low N:P ratio (around 5) for 15 years; and 3) No N supply for 16 years without changing the P supply. When both N and P were supplied during the first 21 years of the study, the lake became highly eutrophic. When only P was supplied for the 16 following years, no significant change in the biomass of phytoplankton was observed. These results suggested that N input had little impact on eutrophication and that, therefore, P input was always limiting algae growth. The authors then extrapolated their results to freshwater

lakes and low-salinity estuaries under the assumption that N shortage is mitigated by N<sub>2</sub> fixation from cyanobacteria, causing P to remain the limiting nutrient. The same authors also pointed out the lack of whole-ecosystem data to attempt to prevent eutrophication by reducing N input (Schindler & Hecky, 2009). While this study is one of the most cited in the literature and has been a reference for years for mitigation strategies in regulatory frameworks, the “P limitation paradigm” has been challenged by several authors (Glibert, 2017; Howarth & Marino, 2006; Paerl, 2009; Conley et al., 2009; Xu et al., 2010; Paerl et al., 2011, 2016; Smith et al., 2016) who argued that only controlling P supply is not always sufficient to prevent HABs, and that a dual control of P and N must instead be established:

- Despite efforts to reduce P inputs, Lake Taihu (China) still regularly undergoes HABs during summer. Bioassays performed on lake samples showed phytoplankton growth<sup>12</sup> was limited by N input during summer (Paerl et al., 2011; Xu et al., 2010). As toxin production by non N<sub>2</sub>-fixing cyanobacteria (e.g. *Microcystis*) can be promoted by N supply (Paerl et al., 2016), reducing N input could have the double effect of decreasing the occurrence of blooms and reducing the hazard associated with HABs.
- P was shown to not be the limiting nutrient in tropical lakes in the study of Sterner (2008).
- The claim that N<sub>2</sub> fixers can alleviate N paucity has been strongly contested because N<sub>2</sub> fixation is an anaerobic process that can only be performed by certain species under oxic when the water is stratified. Also, N<sub>2</sub> fixation is seasonal and unlikely to supply enough N for algal growth (Paerl, 2009). As *Microcystis* have the capacity to move through the water column and, therefore, consume P directly at the sediment layer and bloom at the surface, reducing P loading might not be enough to prevent algae growth in lakes that accumulate P in sediments.
- Controlling only one nutrient can change the nutrient stoichiometry of the water, which can lead to undesirable effects on the biogeochemical cycles and algal biodiversity (Glibert, 2017) as observed in the San Francisco Estuary (USA) and in Lake Taihu in China (Glibert et al., 2011; Glibert et al., 2014).

As noted above, Schindler et al. (2008) extrapolated their results on the impact of P control to all freshwater lakes and low-salinity estuaries. However, these extrapolations have been challenged because of the differences in chemical and physical characteristics of the ecosystems compared. Indeed, N is lost via ammonification and denitrification along the freshwater to marine continuum. As P is not lost, the P:N ratio increases, which causes N limitation to increase along the continuum (Paerl,

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<sup>12</sup> especially the non-N<sub>2</sub> fixing cyanobacteria *Microcystis* spp often responsible for HABs

2009). Another important consequence of only limiting P input in lakes is that all the N that could have been consumed (at high P input) at the start of the freshwater - marine continuum, will now be consumed in estuarine and coastal waters, increasing eutrophication at the end of the freshwater - marine continuum. There is therefore a need to evaluate the full impact of nutrient management on the entire continuum (Conley et al., 2009; Paerl, 2009). This issue will not be further discussed as it is outside the scope of this study.

#### **2.4.3.3. Preventing/mitigating eutrophication in lakes**

Strategies to fight HABs and their symptoms include artificial oxygenation to reduce hypoxia or the addition of chemicals (e.g. lime, aluminium salts, ferrous sulfates or ferric chloride) to precipitate P (Pinay et al., 2018). The use of algicides can be efficient in small lakes not used for recreational activities. In small lakes, disturbing the thermal stratification by mechanical mixing or air bubbling can also be an alternative (Paerl et al., 2001). Biomanipulation via, e.g., the removal of planktivorous fishes to promote the growth of zooplankton can also be efficient to reduce eutrophication and restore biodiversity (Mehner et al., 2002; Pinay et al., 2018). Regardless the successes of these strategies, the main approach to prevent or mitigate eutrophication in lakes remains to reduce P and/or N inputs. This involves the identification of any point-source pollution that would directly increase the nutrient loads to the water. Nonpoint-source pollution, especially those from agricultural runoffs, needs to be well monitored and, when possible, reduced. However, Duarte et al. (2009) argued that a reduction of nutrient input does not necessarily lead to a cessation of eutrophication because of a delay in algal production response due to internal nutrient storage/recycling within the ecosystem (e.g. polyphosphate, (Plouviez et al., 2021)). Consequently, the objectives behind mitigation strategies should be carefully selected and a return to a historical lower trophic state might not be feasible in case of significant baseline shifts (e.g. increased CO<sub>2</sub> concentration, global warming, sea level rise, habitat loss etc.). To better understand how N and P impact algae growth and, by extension, algal N<sub>2</sub>O synthesis, it is therefore important to understand N and P metabolism in algae.

## **2.5. Algae cell biology**

Algae are a polyphyletic group including photosynthetic prokaryotes and eukaryotes (Table 7). This section focuses on the N and P metabolisms of microalgae (unicellular algae) because these nutrients trigger algae blooms and their combined impact on algal N<sub>2</sub>O synthesis is still unknown.

Table 7 : Main features of recognized algae divisions (simplified from Borowitzka (2012)).

Alga division	Cell forms	Habitats	Selected genera
<b>Prokaryotic algae</b>			
Cyanophyta (Cyanobacteria), 2 classes: 'blue-green algae' and 'prochlorophytes'	Unicellular, filaments and colonial	Freshwater, marine and terrestrial	<i>Microcystis</i> , <i>Spirulina</i> , <i>Prochloron</i>
<b>Eukaryotic algae</b>			
Glaucophyta (Glaucophytes)	Unicellular, flagellates	Freshwater	<i>Cyanophora</i>
Rhodophyta 'red algae', 2 classes	Unicellular, filamentous or thalloid	Marine and freshwater	<i>Bangia</i> , <i>Ceramium</i>
Cryptophyta	Mainly unicellular, one filamentous genus, flagellates	Marine and freshwater	<i>Cryptomonas</i>
Heterokontophyta, 7 classes			
- Class Phaeophyceae 'brown algae'	Multicellular	Marine	<i>Ascophyllum</i>
- Class Bacillariophyceae 'diatoms'	Unicellular and colonial	Freshwater and marine	<i>Navicula</i>
Dynophyta 'dinoflagellates'	Predominately unicellular, some coccoïd filamentous, flagellates	Mainly marine	<i>Alexandrium</i>
Haptophyta, 2 classes	Unicellular, flagellates	Mainly marine	<i>Emiliana</i>
Euglenophyta 'euglenoids'	Unicellular with single flagellum	Mainly freshwater	<i>Euglena</i>
Chlorophyta 'green algae', 10 classes			
- Class Chlorophyceae	Coccoïd, unicellular, or colonial flagellates, multicellular or multinucleate filaments.	Mainly freshwater	<i>Chlamydomonas</i> , <i>Chlorella</i>

## 2.5.1. Nitrogen metabolism

### 2.5.1.1. N cycle

N is an essential macronutrient often limiting primary production in aquatic environments (Elser et al., 2007). N is especially needed for the synthesis of amino acids, which are the building blocks of proteins, and nucleotides, which compose nucleic acids (Deoxyribonucleic Acid, DNA; Ribonucleic Acid, RNA). In aquatic environments, N can be found as dissolved inorganics ( $\text{NH}_4^+$ ,  $\text{NO}_3^-$ ,  $\text{NO}_2^-$ ), particulate organic matter, and dissolved gases ( $\text{N}_2$ ,  $\text{N}_2\text{O}$ ). Although the main N reservoir is the atmosphere,  $\text{N}_2$  can only be assimilated by specialized  $\text{N}_2$ -fixing bacteria and archaea. These microorganisms convert  $\text{N}_2$  into inorganic N ( $\text{NH}_3$ ) in a process that can generate the inorganic N used by other organisms (Kuypers et al., 2018). The industrial production of N fertilizer from  $\text{N}_2$  via the Haber-Bosch process has considerably increased the flux of N from the atmosphere to the biosphere (Erisman et al., 2008) (Figure 8).

The total nitrogen (TN) concentration in freshwater lakes is highly variable and can range from 0.1 to 8  $\text{mg}\cdot\text{L}^{-1}$  during summer in temperate lakes (Guildford & Hecky, 2000; Jin et al., 2005; McCauley et al., 1989). A high variability is also reported in tropical-subtropical lakes (Huszar et al., 2006). Seasonal

fluctuation in dissolved N concentration can be observed due to changes in the microorganisms populations involved in ammonification, nitrification and denitrification processes in lakes (Domogalla et al., 1926).

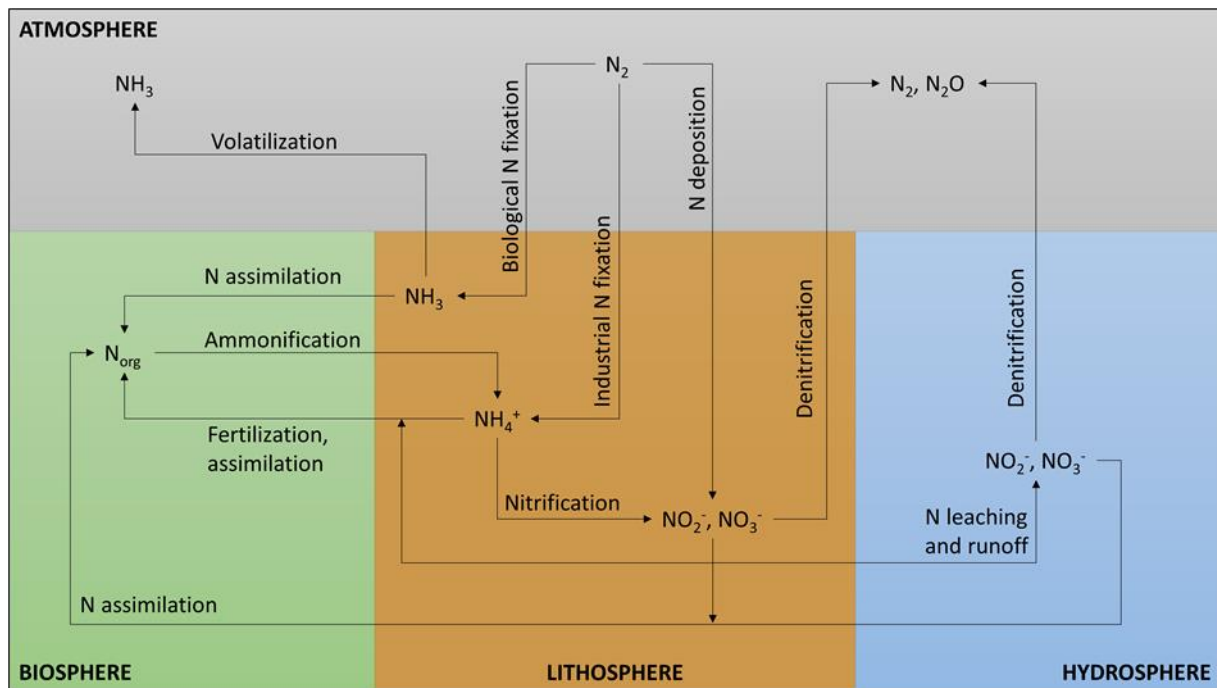


Figure 8: Nitrogen biogeochemical cycle.

Abbreviations/symbols:  $NH_3$ , ammonia;  $NH_4^+$ , ammonium;  $NO_2^-$ , nitrite;  $NO_3^-$ , nitrate;  $N_2$ , dinitrogen;  $N_2O$ , nitrous oxide;  $N_{org}$ , organic nitrogen.

### 2.5.1.2. Microalgae N metabolism

Microalgae can assimilate different forms of N from the environment. For example, *Chlamydomonas* species can assimilate both inorganic (e.g.  $NO_2^-$ ,  $NO_3^-$ ,  $NH_4^+$ ) and organic (e.g. urea, purines, L-amino acids) N (Calatrava, 2018). Ammonium and nitrate are the two main sources of N for algae, with a preference for  $NH_4^+$  in *Chlamydomonas* and other microalgae that repress  $NO_3^-$  assimilation in the presence of ammonium (Gonzalez-Ballester et al., 2004). *Chlamydomonas* genome encodes eight ammonium transporters (AMTs) with different regulation patterns (Gonzalez-Ballester et al., 2004). Some of these transporters are induced under deprivation, suggesting a putative role as high-affinity transporters, whereas other transporters are mainly expressed in N-rich media, indicating a possible role as low-affinity transporters (Gonzalez-Ballester et al., 2004). Once inside the cell, ammonium is transported into the chloroplast where it is incorporated into amino acids through the GS-GOGAT cycle (Glutamine Synthetase – Glutamine OxoGlutarate Amino Transferase).

When  $NH_4^+$  supply is insufficient to sustain growth, microalgae can also assimilate  $NO_3^-$ .  $NO_3^-$  first enters the cell cytosol via  $NO_3^-$  transporters (NRTs and NARs) where it is then reduced into  $NO_2^-$  by

nitrate reductase (NR). Cytosolic  $\text{NO}_2^-$  is transported to the chloroplast where it is further reduced into  $\text{NH}_4^+$  and  $\text{NH}_4^+$  is then incorporated into amino acids as described above. Under specific conditions, some  $\text{NO}_2^-$  can be reduced into NO and this NO is then either reduced into  $\text{N}_2\text{O}$  or converted to  $\text{NO}_3^-$  by truncated haemoglobins (e.g. THB1) associated to reductases (e.g. NR) (see Figure 6 in section 2.2) (Bellido-Pedraza et al., 2020; Calatrava et al., 2017; Kuypers et al., 2018).

## 2.5.2. Phosphorus metabolism

### 2.5.2.1. P cycle

P is an essential macronutrient used in the synthesis of genetic material (DNA, RNA) and intracellular energy carriers in cells (Adenosine Triphosphate, ATP). It is also involved in cellular signalling through phosphorylation and it is a constituent of lipids in cells membranes (Ruttenberg, 2014).

In nature, P is mainly stored as orthophosphate ( $\text{PO}_4^{3-}$ ) in rock sediments. Weathering causes the introduction of this mineral P in soils and rivers from where it is transported to lakes and oceans. There, P can be adsorbed to mineral matter and/or integrated in organic matter until it finally builds up at the bottom of water basins and forms sedimentary rocks under the action of pressure. Over time, the uplifting and exposure of the rock formation reintroduces P into the environment (Figure 9) (Ruttenberg, 2014). The flux of  $\text{PO}_4^{3-}$  to the biosphere has however recently increased due to P mining and use as P fertilizers. The spreading of these fertilizers leads to P run-offs and leaching to lakes, streams and coastal waters, which in turns can cause eutrophication as explained in section 2.4 (Ciais et al., 2013).

Microalgae can assimilate P under the form of dissolved inorganic  $\text{PO}_4^{3-}$  (Ruttenberg, 2014) but phytoplankton can synthesise alkaline phosphatases also enabling the use of dissolved organic  $\text{PO}_4^{3-}$  (Li et al., 2015). Following desorption from sediments or leaching from crops fertilization, particulate P can be a source of dissolved P, especially in the case of internal eutrophication as explained in section 2.4 (Paerl et al., 2001).

The total P (TP) concentration is usually around  $100 \mu\text{g}\cdot\text{L}^{-1}$  in eutrophic lakes (Canfield & Bachmann, 1981; Jacoby et al., 1982; Jin et al., 2005), which is more than twice the typical concentration found in mesotrophic lakes (Jin et al., 2005; Schindler, 1977). The soluble reactive P concentration found in mesotrophic and eutrophic lakes, which is comprised of  $\text{PO}_4^{3-}$ , polyphosphate (PolyP) and organic P compounds, ranges from  $1 - 80 \mu\text{g}\cdot\text{L}^{-1}$  and varies seasonally (Jin et al., 2005; Zhu et al., 2007). Different  $\text{PO}_4^{3-}$  concentrations can be reported with different analysis methods and chromatographic analyses, which are usually used for routine analyses, have been reported to overestimate  $\text{PO}_4^{3-}$  concentration



(Hudson et al., 2000; Rigler, 1966). Using radiobioassays, Hudson et al. (2000) measured the  $\text{PO}_4^{3-}$  concentrations in 53 North American lakes with different trophic status and reported  $\text{PO}_4^{3-}$  at picomolar concentrations ( $0.003$  and  $0.084 \mu\text{g}\cdot\text{L}^{-1}$ ).

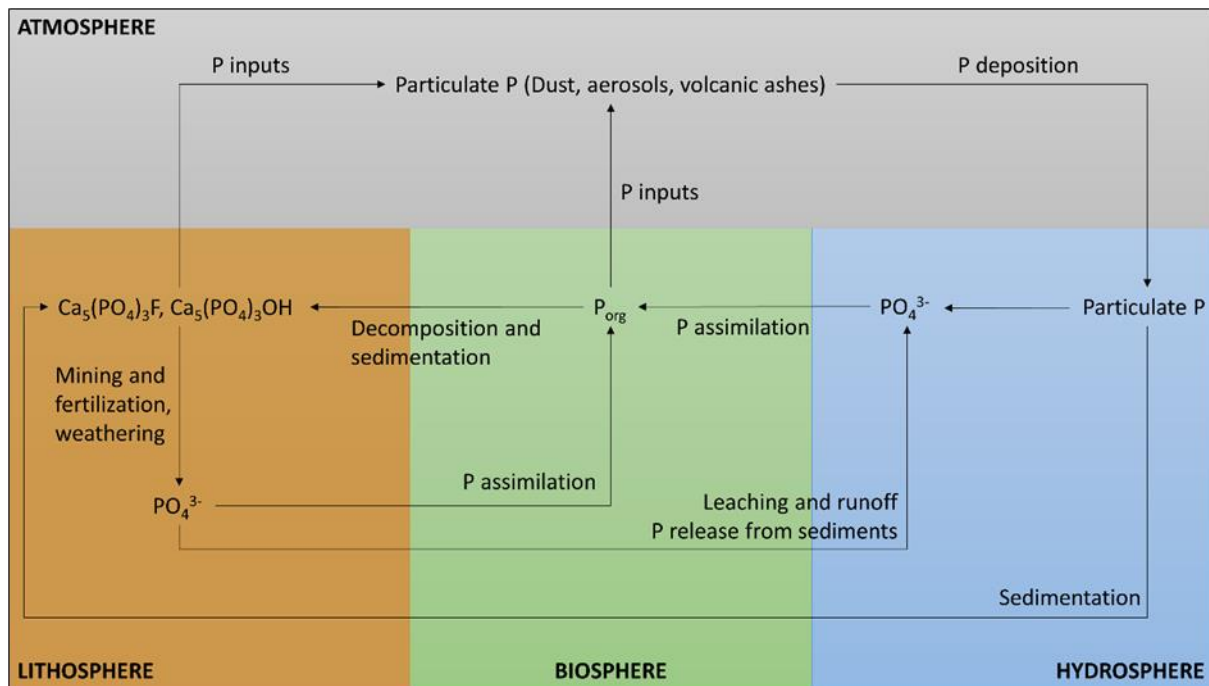


Figure 9: Phosphorus biogeochemical cycle.

Abbreviations/symbols:  $\text{Ca}_5(\text{PO}_4)_3\text{F}$ , fluorapatite;  $\text{Ca}_5(\text{PO}_4)_3\text{OH}$ , hydroxyapatite;  $\text{PO}_4^{3-}$ , phosphate;  $\text{P}_{\text{org}}$ , organic P.

### 2.5.2.2. Microalgae P metabolism

As discussed earlier, P is one of the main nutrients limiting microalgae growth in aquatic ecosystems (Kalf, 1983; Nalewajko & Murphy, 2001; Elser et al., 2007; Abell et al., 2010). Wang et al. (2020) identified 25 putative PHTs (phosphate transporters) in *Chlamydomonas* which are divided into four subfamilies: CrPTA, crPTB, CrPHT3 and CrPHT4. Using four algorithms predicting subcellular localization of the CrPHTs via the detection of targeting motifs in their N-terminal peptides, the PTA family proteins were predicted to be located in the plasma membrane. PTB proteins were thus predicted to target either the secretory pathway or the plasma membrane; PHT3 was predicted to target to mitochondria; and most of the PHT4 family proteins were predicted to be located in chloroplasts. Wang et al. (2020) also identified two transport systems that could be represented by PTB proteins in green algae  $\text{Na}^+/\text{PO}_4^{3-}$  symporters (Reid et al., 2000), and  $\text{H}^+/\text{PO}_4^{3-}$  symporters (Cauthier & Turpin, 1994) that could correspond to PTA proteins. Following the qRT-PCR (quantitative Reverse Transcription-Polymerase Chain Reaction) analyses of CrPHT genes under P depletion, these authors suggested that PTA1 and PTA3 may be low-affinity PHTs while CrPTB2, CrPTB3, CrPTB4, CrPTB5, CrPTB6, CrPTB7, and CrPTB8 were candidates for high-affinity PHTs. Wykoff et al. (1999) suggested

that CrPHT genes may be regulated by the transcription factor *CrPSR1* (phosphate starvation regulator 1) in P-limited medium. Another putative low-affinity transporter encoded by *PTC1* and similar to the yeast *Saccharomyces cerevisiae* has been identified in *Chlamydomonas* but its location and function have not yet been confirmed (Grossman & Aksoy, 2015).

Microalgae can accumulate  $\text{PO}_4^{3-}$  under the form of polyphosphates (polyPs), which are chains of tens to hundreds of  $\text{PO}_4^{3-}$  groups linked by anhydrous orthophosphate bonds (P-O-P). ATP is the  $\text{PO}_4^{3-}$  donor in this reaction and, therefore, polyP synthesis requires energy (Sanz-Luque et al., 2020a; Smith, 1966). In prokaryotes, polyP is synthesised by the enzyme PolyP kinase (PPK) encoded by the gene *ppk* and regulated by the *pho* regulon (Santos-Beneit, 2015; Sanz-Luque et al., 2020a). In cyanobacteria, polyP is degraded by polyphosphatase encoded by the gene *ppx* (Sanz-Luque et al., 2020a). In most eukaryotes, polyP is synthesised by the VTC4 subunit of the Vacuolar Transporter Chaperone (VTC) complex, a transmembrane protein that also translocates the growing polyP into the vacuole where they are stored (Plouviez et al., 2021; Sanz-Luque et al., 2020a). The regulation of polyP synthesis is linked to the inositol phosphate metabolism in eukaryotes (Plouviez et al., 2021; Sanz-Luque et al., 2020a). If phytoplankton is cultivated in a medium deprived of P and is then transferred to a P-rich medium, cells synthesize and accumulate polyP via a mechanism called overcompensation uptake (Aitchison & Butt, 1973). Algae also store polyP without previous deprivation and this accumulation is called luxury uptake (Powell et al., 2011). Luxury uptake has been suggested to be a protective process to prevent future P limitations in ecological niches where strong fluctuations in P availability may occur (Sanz-Luque et al., 2020a). PolyP accumulation has also been observed in microalgae during nutrient (sulfur and nitrogen) deprivation (Sanz-Luque et al., 2020a; 2020b).

To conclude, N metabolism is responsible for  $\text{N}_2\text{O}$  synthesis and N is often limiting primary producers' growth. Thus, it would be tempting to only link algal  $\text{N}_2\text{O}$  emission to N supply in eutrophic ecosystems, as currently done by the IPCC for bacterial  $\text{N}_2\text{O}$  production. However, P is also a critical nutrient often limiting microalgae growth in natural ecosystems. The fluctuations in the availability/supply of N and P together with the algae's capacity to manage their internal N and P cellular pools have a significant ecological impact on the eutrophication of lakes. As P supply can support eutrophication (sometimes synergistically with N supply), simplifying  $\text{N}_2\text{O}$  emissions as only an output of N supply may lead to erroneous prediction of  $\text{N}_2\text{O}$  emissions from eutrophic ecosystems. Understanding the impact of P supply on algal  $\text{N}_2\text{O}$  emissions is, thus, critical to estimate these  $\text{N}_2\text{O}$  emissions accurately.

## 2.6. Conclusions from the literature review

The IPCC currently estimates N<sub>2</sub>O emissions from aquatic environments by assuming that bacterial nitrification and denitrification processes leading to N<sub>2</sub>O synthesis are linearly related to the N inputs received by the aquatic body assessed. Thus, the N<sub>2</sub>O emissions are calculated as a fraction of the N flux reaching the aquatic body defined as emission factors (EFs). However, this ‘bacteria-centric’ assumption that N<sub>2</sub>O emissions only depends on the N input received by an aquatic system may be challenged by the ability of algae to bloom in response to P inputs (or combined N and P inputs) and produce N<sub>2</sub>O. Eutrophic aquatic environments are already known to be a higher source of N<sub>2</sub>O than mesotrophic environments and the IPCC recently acknowledged this fact by increasing the Tier 3 factor used to compute indirect N<sub>2</sub>O emissions from wastewater discharge into eutrophic and nutrient-impacted aquatic environments from 0.005 to 0.019 kg N<sub>2</sub>O-N emitted per kg of N received. However, the higher emissions from eutrophic environment are only considered at a Tier 3 level for wastewater discharge and is not considered in the case of indirect N<sub>2</sub>O emissions from eutrophic aquatic ecosystems receiving N inputs from agricultural N leaching and runoff. In addition, microalgae are still not considered as the potential cause of these N<sub>2</sub>O emissions and therefore, emissions are still only computed from N inputs. Thus, we hypothesize that microalgae could significantly contribute to N<sub>2</sub>O emissions in eutrophic aquatic environments and that it may be necessary to also consider P input to accurately estimate N<sub>2</sub>O emissions from these systems.

Understanding the exact source of the N<sub>2</sub>O emissions in eutrophic aquatic environments is critical for two main reasons:

1. ‘microalgal-N<sub>2</sub>O’ emissions can vary greatly and are affected by very specific parameters (e.g. light). This means that past monitoring based on ‘bacteria-centric’ methodologies may have missed the contribution of microalgal metabolism.
2. while ‘bacterial N<sub>2</sub>O’ is typically directly related to the N load received by a specific environment, ‘microalgal-N<sub>2</sub>O activity’ could be boosted by P (phosphorus supply) triggering eutrophication, meaning that N<sub>2</sub>O emissions from affected ecosystems could no longer be based solely on N-loadings (as currently done).

### 3. Research plan

#### 3.1. Research objectives

This thesis seeks to answer the following questions: 1) Do eutrophic ecosystems generate significant N<sub>2</sub>O emissions and what are the conditions potentially triggering these emissions? 2) Do microalgae contribute to these emissions and, if they do, what are the main chemical (especially the combination of nitrogen and phosphorous supply), environmental (e.g. light, temperature) and biological (associated microorganisms) parameters influencing microalgal N<sub>2</sub>O emissions in eutrophic ecosystems? To reduce the scope of the research, emphasis will be given to the study of eutrophic lakes as a 'worse-case environment' due to occurrence of 'hyper-eutrophication' in New Zealand lakes receiving agricultural runoffs (e.g. Lake Horowhenua), the high potential relevance on these ecosystems (Plouviez et al., 2019a), and practical reasons (site accessibility, health and safety, existing monitoring data). We nevertheless expect that the knowledge generated will be of broader relevance as it may, for example, suggest that other eutrophic environments should be better studied. Besides, this thesis may also improve the accuracy of natural N<sub>2</sub>O emissions estimates, which is important for policy. Finally, the work will ultimately improve our ability to monitor, predict, and mitigate N<sub>2</sub>O emissions from aquatic environments.

To answer the questions listed above, it will be necessary to achieve two main objectives divided in 5 sub-objectives (Figure 10):

#### **Objective 1: Assess the magnitude of N<sub>2</sub>O emissions from eutrophic lakes**

##### Sub-objective 1.1: Field monitoring

N<sub>2</sub>O emissions from selected New Zealand eutrophic and control 'clean' lakes will be determined using a bespoke high-temporal resolution monitoring strategy. This activity will also provide lake microcosms for laboratory assays and inform on possible correlations between water quality and N<sub>2</sub>O emissions to guide the design of sub-objectives 1.2 and 2.1.

##### Sub-objective 1.2: Assess the impact of N and P supply in laboratory assays

The influence of nitrogen and phosphate supply regimes on microalgal N<sub>2</sub>O synthesis will be evaluated using lake microcosms as well as representative model microalgae species *C. reinhardtii*. This will inform us on the potential impact of N and P supply on N<sub>2</sub>O emissions, better inform field monitoring (Sub-objective 1.1) and provide evidence as of the biological origin of these emissions (Objective 2).

These laboratory assays will also provide sample to study the impact of these nutrients on the expression of key genes associated with N<sub>2</sub>O synthesis in *C. reinhardtii* for sub-objective 2.2.

Sub-objective 1.3: Assess the potential global significance N<sub>2</sub>O emissions from eutrophic lakes

Together, sub-objectives 1.1 and 1.2 will provide the experimental data needed to establish the potential of N<sub>2</sub>O emissions from eutrophic lakes from water quality data.

## **Objective 2: Assess the significance of microalgal N<sub>2</sub>O synthesis in eutrophic lakes**

Sub-objective 2.1: Assess the role of microalgae synthesis on N<sub>2</sub>O emissions in laboratory assays

Lake microcosms assays using specific inhibitors as well as assays using knockout and/or knockdown mutants of *C. reinhardtii* will be performed to evaluate the microalgal contribution to N<sub>2</sub>O emissions at a laboratory scale. This preliminary activity will enable us to develop the methods that will then be used to 'isolate' algal production from artificial consortia (with focus on *C. reinhardtii*) and 'natural' lake microcosms. The results from these tests will also inform us on the genes that are more likely to be expressed in field samples (Sub-objective 2.2).

Sub-objective 2.2: Confirm the expression of key genes in selected field samples

Following the identification of key enzymes and molecules involved in the algal N<sub>2</sub>O production pathway (from past and ongoing research with our collaborators), quantitative PCR analyses will be performed on selected field samples to confirm if algae are indeed actively synthesising N<sub>2</sub>O in the monitored ecosystems and/or if other microorganisms significantly contribute to these emissions (Sub-objective 2.1). Using the data gathered before, the potential global environmental significance of N<sub>2</sub>O emissions from eutrophic lakes will be assessed and a "refine" methodology to estimate N<sub>2</sub>O emissions from eutrophic lakes and guidance for mitigation strategies will be proposed (Sub-objective 1.3).

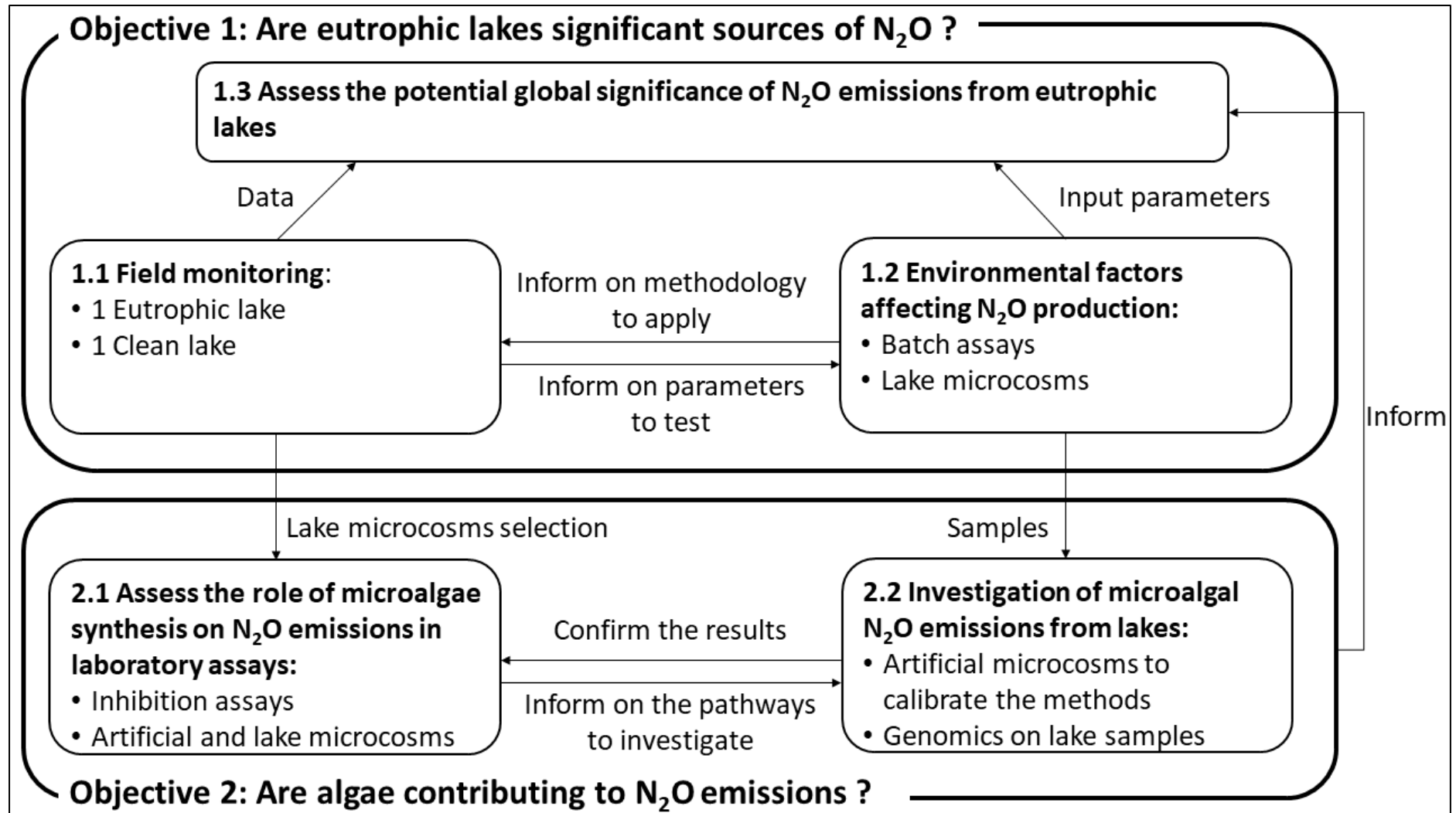


Figure 10: Thesis framework.

### 3.2. Adaptation to the circumstance of the COVID-19 pandemic

Due to the COVID-19 pandemic, I had to start my research in the Department of Biochemistry and Molecular Biology at Cordoba University in Spain for 7 months, where I was hosted by my co-supervisor Dr Emanuel Sanz-Luque. Consequently, I started by conducting experiments on axenic *Chlamydomonas* cultures to train in the laboratory and acquire a better understanding of the molecular pathways involved in N<sub>2</sub>O production under varying conditions and obtain preliminary data (Sub-objective 2.1). The analytical equipment (G2508 gas concentration analyser (PICARRO)) used to measure N<sub>2</sub>O emissions being the property of another department (Department of Botanic, Plant Ecology and Physiology), access to this equipment was limited. I was also able to conduct my literature review while I was there. Once in New Zealand, I started working on building more field knowledge (Sub-objective 1.1) and the influence of environmental factors on microalgal N<sub>2</sub>O emissions (Sub-objective 1.2). However, due to unforeseen circumstances, the PhD thesis was converted into a Master thesis.

Consequently, I could not extend the monitoring and laboratory activities initially scheduled. The following thesis therefore focuses on the presentation of approximately 5 months of field monitoring data with one associated bioassay done in New Zealand, as well as the laboratory data on axenic *Chlamydomonas* cultures done in Spain. While this limited dataset limits my ability to conclude if microalgae indeed contribute to the N<sub>2</sub>O emissions from eutrophic aquatic ecosystems, it provides evidence on how to best direct future research as discussed below.

## **4. Proposed methodology**

### **4.1. Significance of N<sub>2</sub>O emissions from eutrophic ecosystems**

#### **4.1.1. Assessing the magnitude of N<sub>2</sub>O emissions from New Zealand eutrophic lakes**

Sampling sites (n= 2) were selected depending on their trophic state history using data from Horizons Regional Council (<https://www.lawa.org.nz>) and the Ministry for the Environment (<https://data.mfe.govt.nz>), the literature (Wood et al., 2017), and the availability of local water quality and weather data. Accessibility and distance to Palmerston North (Appendix F) were also taken into consideration for health and safety reasons. Lake Horowhenua had thus been selected to be our main case study as it was supertrophic, only 48 km away from Massey University, easy to access, and most importantly, because we were working in collaboration with members of the Lake Horowhenua Trust (members of the Trust have worked hard to develop restoration solutions for the lake and they were interested in the proposed work). Five easily accessible potential back-up locations for field sampling were identified (oligotrophic: Lake Tarawera; mesotrophic: Lake Rotoiti (Tasman), Lake Okareka; eutrophic: Lake Horowhenua, Lake Wiritoa, Lake Dudding). Lake Horowhenua was sampled once per week during Autumn and Winter (when algae growth is typically low) and at least twice per week during Spring and Summer (when algae growth is typically the fastest). The Turitea drinking water reservoir has been selected to be our “clean” control lake as it is considered mesotrophic, only 8 km away from Massey University and easy to access. Access to the reservoir has been allowed by Palmerston North City Council. For each lake, pH, dissolved oxygen and conductivity were measured on site and chlorophyll a, nutrients (e.g. nitrite, nitrate, phosphate) and total suspended solid were quantified from water sampled and N<sub>2</sub>O was quantified using the head space method (Table 8). Methods for routine sampling and analysis were tested during Winter using samples from Lake Horowhenua.



Table 8: Rationale behind the selected parameters to monitor.

Parameter	Rationale	Reference
pH	Influence biological N <sub>2</sub> O synthesis by influencing enzymatic activity. N <sub>2</sub> O solubility is also influenced by pH.	Section 2.2
Dissolved oxygen		Section 2.2, Appendix C
Water temperature	Impact N <sub>2</sub> O solubility and biological reactions.	Paerl et al. (2001)
Conductivity	Impact the nutrients form (ions concentration of the water) and thus inform on nutrient availability which impacts the water trophic state.	García-Lledó et al. (2011)
Chlorophyll a	Proxy for algae biomass concentration which impacts the water trophic state and was shown to be positively correlated to N <sub>2</sub> O emissions	Bartram & Rees (1999), DelSontro et al. (2018)
Total suspended solid	Information on water turbidity which impacts the water trophic state and light availability.	Almeida et al. (2012)
Nutrient concentration	Support algal growth, evaluation of the trophic state.	Almeida et al. (2012)
Radiant energy	Influence enzymatic activity and was shown to influence N <sub>2</sub> O emissions.	Plouviez et al. (2017b)
Wind speed	Influence the average transfer velocity of N <sub>2</sub> O.	Miao et al. (2020)

This field monitoring enabled us to gather data regarding the water quality of the lakes and their N<sub>2</sub>O emissions. These data will be used to establish correlations between the nutrient load of the lake and the N<sub>2</sub>O emissions in future study. They will also inform us on what are the important parameters to look for during further laboratory work (Sub-objective 1.2). These data will also be used in future study to understand what inputs (e.g. TN concentration, temperature) could be needed to model N<sub>2</sub>O emissions (Sub-objective 1.3).

#### 4.1.2. Identification of environmental factors affecting N<sub>2</sub>O production

*Chlamydomonas reinhardtii* wild type (WT) and mutants were cultivated in batch assays to determine how environmental parameters influence N<sub>2</sub>O synthesis. Because parameters influencing microalgae growth have been extensively investigated (Carpenter et al., 1998; Smith et al., 1999; Pinay et al., 2018), this first task was achieved by testing varying nutrient (N and P) concentrations, nutrition modes (mixotrophic<sup>13</sup> or autotrophic<sup>14</sup>) and light conditions. Once the triggers and their thresholds are determined, the experiments will be extended to selected lake samples (microcosms) and seek to systematically evaluate the impact of P load under various conditions known to cause/repress N<sub>2</sub>O

<sup>13</sup> Combination of two different modes of nutrition to acquire carbon by microorganisms: photosynthesis for inorganic carbon (CO<sub>2</sub>) fixation and use of organic sources.

<sup>14</sup> Assimilation of inorganic carbon via photosynthesis.

synthesis during future work. These experiments also provided samples to perform qPCR to study the impact of N and P supply on the expression of key genes involved in N<sub>2</sub>O synthesis.

Overall, these laboratory assays partially enabled us to understand and demonstrate the interactions between P and N supply on N<sub>2</sub>O emissions. Further research is still needed to achieve this Sub-objective. These data informed us on the parameters to consider and monitor during our field work (Sub-objective 1.1). If needed, this information will also serve to establish how P supply should be considered to estimate N<sub>2</sub>O emissions (Sub-objective 1.3).

### **4.1.3. Determination of the potential global environmental significance of N<sub>2</sub>O emission from eutrophic lakes**

Based on the data obtained during the previous tasks, future work will aim at determining the key parameters (e.g light, nutrients) needed to estimate N<sub>2</sub>O emissions from lakes. Using the field data and new knowledge, and especially if future study demonstrate microalgae are playing an important role in these emissions (Sub-objectives 2.1 and 2.2), a specific methodology for N<sub>2</sub>O emissions from eutrophic lakes will be proposed. This method will finally be applied during further study to estimate the global significance of N<sub>2</sub>O emissions from eutrophic lakes and explore how these new tools and data could guide new strategy for monitoring and mitigation.

## **4.2. Microalgal contribution to N<sub>2</sub>O emissions from eutrophic lakes**

### **4.2.1. Evaluation of the microalgal contribution to N<sub>2</sub>O emissions in the lab**

The key pathways and parameters influencing algal N<sub>2</sub>O synthesis were partially confirmed using *Chlamydomonas* parent mutants' strains as positive and negative controls to calibrate the methods for further work. New candidate proteins involved in the N<sub>2</sub>O synthesis pathways were also investigated using specific *Chlamydomonas* mutants and chemical inhibitors. It is important to note that due to the need to express a strong N<sub>2</sub>O synthesis in these assays, the nutrient concentrations tested may be irrelevantly high in comparison to the concentrations found in natural lakes. This means that the N<sub>2</sub>O emissions rates reported should not be directly extrapolated to estimate field conditions without extensive verification. Negative controls (i.e. killed cells or sterile media supplied with NO<sub>2</sub><sup>-</sup>) were used to confirm that N<sub>2</sub>O is biologically produced by the microalgae and not via abiotic reactions.

To identify the 'microalgal contribution' to N<sub>2</sub>O emissions within diverse microcosms, N<sub>2</sub>O emissions from artificial algal-bacterial microcosms and selected lake samples will be tested during future work

under a range of conditions selected to enhance or repress a specific activity (e.g. light vs darkness, anoxia vs hyperoxia etc). This will also include the testing of various substrate and inhibitor, which will be challenging given the similarities in the biological pathways involved (see Section 2.2). Wild type *Chlamydomonas* and knockout mutants unable to produce N<sub>2</sub>O will therefore be cultured with various N<sub>2</sub>O-producing model bacteria (e.g. the nitrifier *Nitrosomonas europaea* (Arp et al., 2003)). The comparison of the N<sub>2</sub>O production between the microcosms with WT and mutants and between pure and mixed culture will make it possible to quantify the contribution of algae to N<sub>2</sub>O emissions while also assessing the potential ecological interaction effects between these organisms (e.g. are algal and bacterial N<sub>2</sub>O activities complementary or competitive?). Batch assays will also investigate the effect of different inhibitors on N<sub>2</sub>O emissions from algal-bacterial artificial co-cultures. Methods using specific inhibitors which could be then applied to real lake samples will thus be developed.

Overall, this second objective will provide new fundamental knowledge on the direct and/or indirect role of microalgae in N<sub>2</sub>O emissions from eutrophic ecosystems and inform us on the N<sub>2</sub>O synthesis pathways that should be investigated during the field work (Sub-objective 2.2). The biological origins of N<sub>2</sub>O emissions will thus be identified and these emissions will be quantified more accurately (Sub-objective 1.3) during future work.

#### 4.2.2. Investigation of microalgal N<sub>2</sub>O emissions from lakes

To evaluate the microalgal contribution to N<sub>2</sub>O emissions from lakes, qPCR will be used during future research to estimate the relative abundance of N<sub>2</sub>O-specific bacterial (*nirS*, *nirK*, *Nos*), archaeal (*AmoA*), fungal (*p450nor*) and algal (*NR*, *Cyp55* and potentially *FLVs*) genes over reference genes (e.g. 16S rRNA, ITS and 18S rRNA) in selected samples (Table 9). This work will aim at identifying the microbial community present in our samples and identifying their potential biochemical functions.

Table 9: Molecular biology principles

Method	Principle	Critical point to consider	Reference
qPCR	Measure the amount of PCR products (i.e. specific genes) in a DNA sample detected by fluorescence using a DNA binding dye as it accumulates during PCR cycles.	Long process Selection of controls Requires optimization (DNA extraction)	Smith & Osborn, 2009

To circumvent the inherent limitations potentially occurring when dealing with environmental DNA<sup>15</sup> (eDNA) (Buxton et al., 2017; Goldberg et al., 2016; Taberlet et al., 2018) and/or RNA (Ritalahti et al., 2010), guidelines recently developed by an international expert panel will be followed during future work (Goldberg et al., 2016). In order to design the eDNA study, a pilot study should be conducted first using pure *Chlamydomonas* cultures and artificial microcosms to assess the suitability of the methods for water collection, sample preservation, extraction, and analysis.

While the evidence obtained during this work is necessary to identify the microorganisms present in our samples, it is not sufficient to discriminate the organisms responsible for N<sub>2</sub>O emissions in complex microcosms. This is why these data will be complemented with specific inhibition assays (Sub-objective 2.1) to target the putative microorganisms producing N<sub>2</sub>O.

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<sup>15</sup> Complex mixture of genomic DNA from many different organisms found in an environmental sample such as water or soil (Taberlet et al., 2018).

## 5. Material and methods

### 5.1. Lake sampling

In this work, the water quality and the N<sub>2</sub>O concentration of the samples were quantified to determine the magnitude of N<sub>2</sub>O emissions and gather data to study the interaction between selected parameters (e.g. N and P load) and N<sub>2</sub>O emissions as described in section 4.1.1 (Sub-objective 1.1).

Lakes monitoring for GHGs have increased in the last decades. Various methods have been used to measure N<sub>2</sub>O emissions as further described in Table 10 and Table 11 below.

Table 10: Methods for sampling to measure N<sub>2</sub>O emissions from water bodies.

Sampling technique	Principle	Advantages	Limits	Reference
Floating chamber	N <sub>2</sub> O flux is calculated by measuring the gaseous concentration of N <sub>2</sub> O in a closed chamber floating at the surface of the water	Easy to deploy, cheap, suitable for large area (chamber size up to 150 m <sup>3</sup> )	Causes disturbance such as shading, water turbulence	Stow et al. (2005), Rapson & Dacres (2014), Wang et al. (2020)
Headspace	Dissolved N <sub>2</sub> O concentration is estimated by measuring the gaseous concentration of N <sub>2</sub> O in an airtight container headspace assuming the dissolved N <sub>2</sub> O concentration in the aqueous phase is in equilibrium with the gas phase	Easy, non-disturbing, cheap	Sampling intensive to estimate the N <sub>2</sub> O flux	Butler & Elkins, (1991), Mengis, (1997), Reay et al. (2003)
Micrometeorological techniques	N <sub>2</sub> O flux is measured from a large area based on N <sub>2</sub> O measurements from a gas sensor and meteorological data (e.g. wind speed, temperature)	Non-disturbing, suitable for large area (1 – 10 km <sup>2</sup> )	Expensive and require fast and sensitive gas analyzers	Rapson & Dacres (2014)

The headspace sampling technique has been selected for our study because this method is easy to deploy, cheap and, contrary to floating chamber, this technique does not cause disturbance such as shading. This is critical for the study of algal N<sub>2</sub>O emissions because algae N<sub>2</sub>O synthesis was shown to be influenced by light (Plouviez et al., 2017b).

Table 11: Methods for analysis of N<sub>2</sub>O in gas and liquid samples, simplified from Rapson & Dacres (2014).

Analytical technique		Sensitivity	Advantages	Limits
Chromatography	Gas chromatography equipped with an electron capture detector (ECD)	Limit of detection: 30 ppb Precision: 0.18–0.4 ppb	Low operational cost, widely used	Non-continuous measurements, standard needs to be run every 3 samples, long run time
Optical	Fourier-transform infrared spectroscopy (FTIR)	Limit of detection: <1 ppb Precision: 0.03–0.1 ppb	Continuous, portable	High capital costs
	Laser absorption spectroscopy: Lead Salt Lasers	Precision: <1 ppb	Rapid and highly sensitive, low interference from other trace gases	High capital and operational costs (e.g. require cryogenic cooling)
	Laser absorption spectroscopy: Quantum Cascade Lasers (QCLs)	Precision: 0.05 ppb	Fast, portable	High capital costs
Amperometric	Microsensor	Limit of detection: 22 ppb	Low capital costs, portable	Sensor drift over long-term use, fragile

The chromatographic analysis method has been selected because our laboratory has extensive experience using this technique which is cheap, easy to maintain and to operate (Hamilton & Lewis, 2006).

Superficial water samples of 500 mL were taken weekly at 5 m and 15 m from the edge of lake Horowhenua and immediately put on ice in darkness to preserve the chlorophyll. For N<sub>2</sub>O quantification, 6 mL of water samples were injected in triplicates using 12 mL vials (Labco Exetainer® 12ml Vial, Lampeter, UK) previously flushed with dinitrogen gas as described by Plouviez et al. (2019b). 200 µL of zinc chloride (50% w/v) were added to the vials to inhibit the metabolism of the present organisms (Kamp et al., 2013). Two aquatic environments were selected for this research:

**Lake Horowhenua**, the largest natural lake in the Manawatū-Whanganui region, was selected as our main case study. It is located at 2 km west of Levin and 7 km east of the coast of the Tasman Sea (Latitude: 40° 36' 36.301" S; Longitude: 175° 15' 19.718' E). The lake is a polymictic aeolian shallow lake measuring 304 ha with a maximum depth of 2 m and considered supertrophic. The lake used to receive the treated sewage waters from Levin, which participated in its nutrients' enrichment. Access to the lake was kindly allowed by Jonathan Procter who is in close collaboration with the Lake Horowhenua Trust.

**The Turitea drinking water reservoir** was selected as a control "clean" lake. This is an artificial lake formed following the damming of the Turitea stream and located at 8 km from Massey University. The reservoir measures 12 ha with a maximum depth of 10 m and is considered mesotrophic. It constitutes the main water supply of Palmerston North which is why access to the reservoir is restricted. Access to the reservoir was kindly allowed by Steve Langevad who was the Water Treatment Supervisor from Palmerston North City Council.

## 5.2. Strains selection and maintenance

### 5.2.1. Strains used

The model unicellular green alga *Chlamydomonas reinhardtii* was selected to study N<sub>2</sub>O production pathways because this microorganism has already been extensively studied, its genome has been sequenced (<https://phytozome.jgi.doe.gov>), there are numerous molecular techniques optimized for this organism to create mutants, and mutants libraries are readily available (relevant strain have already been selected and tested (Table 12)). All strains were maintained on a sterilised solid Tris-Acetate-Phosphate (TAP) medium containing 16 g·L<sup>-1</sup> of agar.

Table 12: Strains used in that study.

Strain	Genetic background	Gene common name	Notes	Source
M3	nii1 <sup>-</sup>	NiR (Cre09.g410950)	Originate from the cross between G1 and 04-1	Navarro et al., 2000
CC-4533 (CMJ030)	Wild type (cw15, mt <sup>-</sup> )	cw15	Originate from the cross between D66+ and 4A-. It cannot grow in nitrate.	Zhang et al., 2014
LMJ.RY0402.052321	cw15, flvb <sup>-</sup> , mt <sup>-</sup>	FLVB (Cre16.g691800)	Background strain: CC-4533	Li et al., 2019
LMJ.RY0402.177695	cw15, cyp55 <sup>-</sup> , mt <sup>-</sup>	CYP55 (Cre01.g007950)	Background strain: CC-4533	Li et al., 2019
6145c	Wild type		CC-2895 mt <sup>-</sup> : derivative of Ruth Sager's strain 6145 (see CC-1691)	Chlamycollection.org
CC-1691	Wild type		mt <sup>-</sup> equivalent to Sager's basic wild type strain (see CC-1690). It has the wild type alleles at the NIT1 and NIT2 loci and can grow on nitrate. Phenotype: yellow in the dark.	Chlamycollection.org
CC-1690 mt + (21gr)	Wild type		Sager's basic wild type strain. It has the wild type alleles at the NIT1 and NIT2 loci and should be green in the dark.	Chlamycollection.org

### 5.2.2. Microalgae cultivation

The protocols for *Chlamydomonas* culture and maintenance are well established (Harris, 2009). An E-flask of 100 mL filled with 40 mL of buffered TAP medium was inoculated with axenic cultures of the chosen strain. The cultures were then incubated for 2 days under continuous agitation (100 rpm), constant illumination ( $54 \mu\text{mol photons}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ ) and a temperature of 24°C. The cultures were then transferred into a 1 L E-flask to which buffered TAP medium was added up to 400 mL and the cultures were then incubated in the same conditions as the inoculum culture for 2 days. The cultures were then refreshed daily and used as the inoculum culture for the experiments for a maximum of 5 days.

### 5.3. Batch assays

During these assays, the N<sub>2</sub>O production of the cultures were quantified to determine the impact of environmental factors (e.g. N and P supply) on microalgal N<sub>2</sub>O emissions as described in section 4.1.2 (Sub-objective 1.2).

As described by Harris (2009), axenic *Chlamydomonas* cultures were previously grown in buffered TAP medium under light in a 1 L E-flask. For experiments in High Salt Medium (HSM) or minimal medium (MM), cells were induced for 6 h in the chosen medium: around 200 mL of cells in exponential phase were taken and put in 50 mL sterile falcon tubes to be centrifuged for 3 minutes at 1,200 x g. The supernatant was discarded, and the cells were then cleaned twice with the final medium. The cells were then grown in this medium under light for 6 h in a 1 L E-flask.

The cultures were diluted at the beginning of each experiments to obtain a chlorophyll concentration of  $10 \mu\text{g}\cdot\text{mL}^{-1}$  (equivalent to  $0.25 \text{ g DCW}\cdot\text{L}^{-1}$ , see Appendix G). To get N-free algae culture, the cells growing in ammonium were cleaned twice with N-free medium.

### 5.3.1. Experimental set-up in Spain

For  $\text{N}_2\text{O}$  quantification, cells were resuspended in 250 mL of medium supplied with nitrite and put in a 1 L Duran bottle.  $\text{N}_2\text{O}$  emissions were measured from the cultures for 24 h using cavity ring-down spectroscopy (CRDS) in the G2508 gas concentration analyser (PICARRO).

Note: This equipment belongs to another Department of the University of Cordoba (Department of Botanic, Plant Ecology and Physiology) where it is used to measure  $\text{N}_2\text{O}$  emissions from agricultural soils. Therefore, its coupling to bottles containing axenic cultures of *Chlamydomonas* required optimization.

During the experiment, cells were grown in an incubator (Versatile Environmental Test Chamber MLR-351, Sanyo) at  $25^\circ\text{C}$  with continuous agitation at 100 rpm (Rotamax 120, Heidolph, Schwabach, Germany) and under continuous illumination using fluorescent lamps and a light irradiance around  $130 \mu\text{mol photons}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ . A summary of the conditions tested can be found in Table 13.

Table 13: Summary of the conditions tested in the batch assays conducted in Spain.

Strain	Medium	Conditions	Nutrient addition	Number of replicates
<b>M3</b>	HSM without nitrogen	Light	10 mM $\text{NO}_2^-$	3
	TAP medium without nitrogen	Light	0.01 mM $\text{NO}_2^-$	1
			0.1 mM $\text{NO}_2^-$	1
<b>6145c</b>	HSM without nitrogen	Light	10 mM $\text{NO}_2^-$	3
				1
	HSM without nitrogen + DCMU	Light	10 mM $\text{NO}_2^-$	4
	HSM without nitrogen + DMSO	Light	10 mM $\text{NO}_2^-$	1
<b>cc-4533 (CMJ030)</b>	HSM without nitrogen	Light	10 mM $\text{NO}_2^-$	8
				1
<b>052321 (flvb mutant)</b>	HSM without nitrogen	Light	10 mM $\text{NO}_2^-$	3
<b>177695 (cyp55 mutant)</b>	HSM without nitrogen	Light	10 mM $\text{NO}_2^-$	3
<b>21gr</b>	HSM without nitrogen	Light	10 mM $\text{NO}_2^-$	1
<b>No cells</b>	HSM without nitrogen	Light		1



### 5.3.2. Experimental set-up in New Zealand

While the actual conditions changed depending on the purpose of the experiment, the following protocol developed by Plouviez et al. (2017a) was used to study N<sub>2</sub>O synthesis in pure cultures: Cells were resuspended at 10 µg chlorophyll·mL<sup>-1</sup> in the treatment medium supplied with the chosen concentration of nitrite and put in serum flasks sealed with rubber septa. Cells were incubated in a Multitrom (Infors HT) at 25°C under continuous agitation and either under continuous illumination (120 µmol photons·m<sup>-2</sup>·s<sup>-1</sup>) or in the dark. Cell concentration was assessed by measuring both chlorophyll concentration and the dry cell weight.

### 5.4. Inhibition assays

In these inhibition assays, the N<sub>2</sub>O production from axenic microalgal culture and lake samples were quantified to evaluate the contribution of each type of microorganisms (especially bacteria and algae) as described in section 4.2.1 (Sub-objective 2.1).

The experiments were performed using the protocol described above but DCMU (3-(3,4-dichlorophenyl)-1,1-dimethylurea), an inhibitor of the photosystem II, was added to the medium after 15 h to investigate the contribution of the photosynthetic transport electron chain to N<sub>2</sub>O production. Other inhibitors (especially bacterial inhibitors) were planned to be used in similar assays conducted with artificial microcosms and lake samples (Table 14). Experiments will be conducted during future work to check the inhibition efficiency and the specificity of each inhibitor on the microbial cultures. The use of microalgae mutants that are not producing N<sub>2</sub>O will be preferred to the use of microalgae inhibitors because inhibitors are not always specific and could thus affect the growth of other organisms.

Table 14: Inhibition assays.

<sup>1</sup> essential posttranslational modification of the translation initiation factor 5A that stimulates ribosomal peptidyltransferase in archaea.

<sup>2</sup> Cell damage due to the oxidation of proteins, lipids, and DNA.

Organism	Inhibitor	Main effect	Reference
Fungi	Carbendazim	Inhibit the microtubule polymerization, disrupting cell division	Mahan et al. (2005)
	Thiophanate-methyl		
	Benomyl		
Archaea	N1-guanyl-1, 7- diaminoheptane	Hypusination <sup>1</sup> inhibitor	Jansson et al. (2000)
Bacteria	Penicillin G potassium	Affects bacterial cell wall synthesis	Guieysse et al. (2013)
	Streptomycin sulfate	Affects bacterial ribosomes	
Microalgae	Hydrogen peroxide (cyanobacteria)	Induce a severe oxidative stress <sup>2</sup>	Hanson & Stefan (1984)
	Copper sulfate	Interacts with proteins, nucleic acids and metabolites	Zhou et al. (2020)

## 5.5. Primers design and optimisation

In selected experiments, the expression of key genes was quantified to determine the impact of test conditions (e.g. P supply, presence of inhibitor as described in sections 4.1.2 and 4.2.1) on N<sub>2</sub>O synthesis regulation (Sub-objectives 1.2 and 2.1) .

The expression of genes encoding putative proteins involved in N<sub>2</sub>O production in *Chlamydomonas* (*FLVA*, *FLVB*, *NIRK*, *HCP1*, *HCP2*, *HCP3*, *HCP4* and *CYP55*) was measured using qPCR. Primers for these genes have been designed using QuantPrime online tool (<https://quantprime.mpimp-golm.mpg.de/>) and the efficiency of the PCR using these primers was checked (Table 15).

Based on the kinetics of N<sub>2</sub>O production from the axenic *Chlamydomonas* cultures, gene expression was measured at 6 different times (0 h, 1 h, 3 h, 6 h, 17 h and 24 h) in M3 and CMJ030 cultures in HSM-N (control) and HSM-N + 10 mM NO<sub>2</sub><sup>-</sup> (treatment).

Table 15: Sequence of the designed primers.

\* This gene had been designed and the efficiency of the PCR had been checked before this study.

Gene	Transcript identifier	Primer forward sequence	Primer forward length	Primer forward Tm	Primer forward GC	Amplicon length
FLVA 2	Cre12.g531900.t1.1	ACTGCAGGCAATCGCTCCTTTG	22	64,315	54,545	62
FLVB	Cre16.g691800.t1.1	GCATGAGCGATGTGGTCATGTG	22	62,788	54,545	130
CYP55*	Cre01.g007950.t1.1	GATGGCCGCTTCTCCAAGGTC	21			
HCP1	Cre09.g391450.t1.1	TCGCATTCCAGAAGAGCACTGG	22	63,232	54,545	119
HCP2	Cre02.g129550.t1.1	CGAATGGCATAACGGTACTAAGGC	24	64,186	54,167	68
HCP3	Cre02.g129500.t1.1	CGCGATGCCTCTATTAAGCCAAG	24	64,254	54,167	72
HCP4	Cre09.g391650.t1.1	CAATAAAGAGGCAGGCGGTAG	22	62,28	54,545	80
NIRK	Cre08.g360550.t1.1	AGGAGTTCTACGTGGTGCAGTC	22	62,33	54,545	70

Lyophilised primers (obtained from Invitrogen) were resuspended in ultrapure distilled water at 100  $\mu$ M. Primer stock solution was prepared by mixing forward and reverse primers at a final concentration of 10  $\mu$ M of each in distilled water.

Primer efficiency was assessed by PCR, using a serial dilution of a mix of cDNA previously obtained as template. The log copy number (or log dilution times) against the number of PCR cycle (Ct) was plotted to calculate the slope, which was used to calculate the efficiency according to the following Equation 1:

$$Efficiency (\%) = 10^{\frac{1}{slope}} - 1 \quad \text{Equation 1}$$

Briefly, the PCR mix was prepared with 0.8  $\mu$ M of primer stock solution, cDNA, ultrapure water and 10  $\mu$ L of reagent (SsoFast™ EvaGreen® Supermix); for a total of 20  $\mu$ L per well as indicated in the manufacturer's instructions. Two replicates per dilution were run. As internal control for relative quantification (housekeeping genes) we used *CBLP*, *UBI* or *ACT1* genes, whose expressions remained unchanged upon different nitrogen sources. The PCR protocol entailed one cycle at 96 °C for 2 minutes followed by 40 cycles of 10 seconds at 96°C and 30 seconds at 60°C.

For relative quantification we used the  $\Delta$ Ct method to compare the level of changes in gene expression. To do so, the fold difference was used and calculated using Equation 2.

$$Fold\ difference = 2^{(\overline{Ct}_{reference\ gene} - \overline{Ct}_{gene\ of\ interest})} \quad \text{Equation 2}$$

Where the fold difference ( $\Delta$ Ct) is the relative expression of the gene of interest compared to the reference gene;  $\overline{Ct}_{gene\ of\ interest}$  is the mean threshold cycle value of the two replicates of the gene of interest and  $\overline{Ct}_{reference\ gene}$  is the mean threshold cycle value of the two replicates of the reference gene.

## 5.6. Analysis

### 5.6.1. Algal cell quantification

#### 5.6.1.1. Chlorophyll measurement

The total chlorophyll concentration in pure cultures and artificial microcosms was determined by spectrophotometry (DU® 800 UV/Visible Spectrophotometer, Beckman Coulter, Brea, California, USA) from triplicate samples as described by Wintermans & De Mots (1965). A sample of 750 µL of culture was transferred in a 1.5 mL Eppendorf and centrifuged for 75 seconds at 12,000 x g. The supernatant was discarded and 750 µL of ethanol was added to the pellet. The mixture was then mixed using a vortexer and centrifuged for 75 seconds at 12,000 x g. The absorption of the supernatant at 649 nm ( $A_{649}$ ) and 665 nm ( $A_{665}$ ) was then measured and the total chlorophyll concentration was calculated by Equation 3:

$$\text{Chlorophyll concentration } (\mu\text{g} \cdot \text{mL}^{-1}) = 6.1 \times A_{665} + 20.04 \times A_{649} \quad \text{Equation 3}$$

In lake samples, the total chlorophyll concentration was also determined using Equation 3 but to extract chlorophyll, a known volume was first filtered by vacuum suction through a 2.7 µm glass microfiber filter (47 mm, MicroScience MS GC filters, Australia) in triplicates as described by Papista et al. (2002) and Sartory & Grobbelaar (1984). The filters were then immersed in a solution of ethanol at 90% and left in a cold room at 4°C for 24 h. The extract was then centrifuged for 2 min at 3,000 x g and the supernatant was transferred to a 15 mL clean Falcon tube. The colour of the pellet was checked to be sure that no visible chlorophyll remained in it. The optical density of the clarified extract at 649 nm (chlorophyll b) and 665 nm (chlorophyll a) was then measured using a UV-1800 spectrophotometer (Shimadzu, Japan).

#### 5.6.1.2. Dry cell weight and Total suspended solid

As described by Béchet et al. (2015), 2.7 µm glass microfiber filters (47 mm, MicroScience MS GC filters, Australia) previously dried at 105°C for 24 h were weighted, then a known volume of cells were filtered and rinsed with the same volume of distilled water to remove any dissolved salts. The filters were then dried for 1h at 105°C and weighted. The mass concentration was then calculated using Equation 4:

$$\text{Mass concentration} = \frac{\text{Weight}_{\text{Filter with cells}} - \text{Weight}_{\text{Filter}}}{\text{Volume of cells}} \quad \text{Equation 4}$$

### 5.6.1.3. Cell counting

As described by Guillard & Sieracki (2005), the number of cells in the culture were counted using a Bürker pattern counting chamber. For this purpose, a sample of 10 µL of culture were put on the counting chamber and photos of the cells inside 12 squares diagonally arranged on the counting chamber using the microscope's magnification 40X were taken. The cells inside each square and on the upper and right edges were counted. The Mesurim software was used to count the cells when the number of cells per square exceeded 10. The final concentration was calculated using Equation 5:

$$\text{Cell concentration} = \frac{N_{\text{cells}}}{N_{\text{squares}} \times V} \times DF \quad \text{Equation 5}$$

Where  $N_{\text{cells}}$  is the total number of cells counted,  $N_{\text{squares}}$  is the number of squares in which cells were counted,  $V$  is the liquid volume in one square and  $DF$  is the dilution factor.

### 5.6.1.4. Optical density

The culture optical density was determined from triplicates by measuring the absorbance at 683 nm and 750 nm of 1 mL of culture. The blank was previously done using the same medium as the culture.

## 5.6.2. pH, Dissolved oxygen, and Conductivity

The dissolved oxygen (DO) concentration, conductivity and pH were monitored using a multimeter (Orion™ Star A329, Thermo Scientific) equipped with the relevant probes (Orion™ 4-Electrode Conductivity Cell, Thermo Scientific; Orion™ RDO™ Dissolved Oxygen Probe, Thermo Scientific; Orion™ ROSS Ultra™ Low Maintenance Triode™ pH/ATC electrode, Thermo Scientific). The calibration of the probes was checked before each sampling.

## 5.6.3. Nutrients

The concentrations of  $\text{PO}_4^{3-}$ ,  $\text{NO}_2^-$ ,  $\text{NO}_3^-$  and  $\text{NH}_3$  in filtered samples (0.2 µm, Sartorius Minisart filters, Gottingen, Germany) were quantified using a Lachat Quikchem 8500 Series 2 Flow Injection Analysis System (Lachat Instruments, Loveland, Colorado, USA). In the case that the equipment broke down,

well-described methods that we have experience with were used to measure  $\text{NO}_2^-$  i.e. the substrate of  $\text{N}_2\text{O}$  <sup>16</sup>.

A Dionex ICS-2000 Ion Chromatograph (Dionex Corporation, Sunnyvale, California, USA) that was used to measure the concentration of chloride and sulphate in filtered samples could also be used as back-up to measure concentrations of  $\text{NO}_3^-$ ,  $\text{NO}_2^-$  and  $\text{PO}_4^{3-}$ .

The dissolved carbon and dissolved nitrogen concentrations will be determined using a Total Organic Carbon Analyser (TOC-L, Shimadzu, Japan) during future work.

#### 5.6.4. Light intensity

The light intensity in the incubators was measured using a PAR meter taking the mean of 9 measurements inside the incubators. During lake sampling, meteorological data (solar irradiance, air temperature, wind velocity and direction, and relative humidity) were taken from a close weather station (Levin Electronic Weather Station, 1.2 km away from the Horowhenua lake; Palmerston North Electronic Weather Station, 7.4 km away from the Turitea reservoir). The data from the National Climate Database collected by the National Institute of Water and Atmospheric Research (NIWA) and MetService were accessed using the CliFlo web system (<https://cliflo.niwa.co.nz/>). An example of how data can be extracted and used is provided by the CliFlo web system.

#### 5.6.5. $\text{N}_2\text{O}$ measurement

$\text{N}_2\text{O}$  was quantified using either a G2508 gas concentration analyser (Spain) or an ECD-gas chromatograph (New Zealand).

**G2508 analyser:** The  $\text{N}_2\text{O}$  concentration in the gas phase of a Duran bottle in which the algal culture was growing was continuously (at least one measure per minute) measured by Cavity Ring-Down Spectroscopy (G2508 Gas Concentration Analyzer (Picarro), Santa Clara, California, USA). The accuracy

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<sup>16</sup> The nitrite concentrations in the culture supernatants were determined by a colorimetric assay based on Griess reaction. A sample of 750  $\mu\text{L}$  of algal culture was put in a 1.5 mL Eppendorf and centrifuged for 75 seconds at 12,000  $\times g$ . The cell-free supernatant was kept in the freezer. A working reagent mixture was prepared mixing equal volumes of 10  $\text{g}\cdot\text{L}^{-1}$  sulfanilamide with 200 mL of HCl 12 N and 200  $\text{mg}\cdot\text{L}^{-1}$  N-(1-naphthyl)ethylenediamine dihydrochloride. A volume of 150  $\mu\text{L}$  of the prepared working reagent mixture was put in a flat-bottom 96-wells microtiter plates and 75  $\mu\text{L}$  of supernatant was added to the wells. The mixture was incubated for 2 minutes and the absorbance at 540 nm was then measured using a microplate reader (iMark, Bio-Rad). Samples were diluted as needed and the nitrite concentration was calculated using previously made calibration curves containing seven points of known concentrations ranging from 0 to 4 mM  $\text{KNO}_2$ .

of the equipment was verified by measuring a standard mix of gas (499.5 ppm CO<sub>2</sub>, 0.498 ppm N<sub>2</sub>O, 1.99 ppm CH<sub>4</sub>). In this protocol, the gas entering the analyser first went through a desiccator and a filter to avoid the accumulation of water or particles inside the equipment and the gas coming from the pump was filtered with a hydrophilic PVDF 0.22 µm filter (Millipore Millex GV) before entering the Duran bottle to avoid any contamination.

**GC-ECD:** In the serum flasks, gaseous N<sub>2</sub>O concentration was measured by sampling 5 mL of the gas phase with a gas-tight syringe and a needle. This sample was then injected onto a ECD-gas-chromatography (Shimadzu GC-2010, Shimadzu, Japan) using a 1000 µL sample loop (380°C), a Alltech Porapak QS 80/100 column (70°C, Sigma-Aldrich, USA) and an electron conductivity detector (315°C) as described by Guieysse et al., (2013). Calibration was achieved using a 100 ppm N<sub>2</sub>O standard in N<sub>2</sub> (neutral gas).

**Computation and key assumptions:** In both serum flasks and Duran bottles, we assumed the gas and the liquid phase N<sub>2</sub>O concentrations were at equilibrium at the time of sampling. The concentration of N<sub>2</sub>O dissolved in the liquid phase was therefore calculated using Henry's law at 25°C and the total amount of N<sub>2</sub>O produced was then calculated by summing up the amount of N<sub>2</sub>O that was present in the gas and liquid phase as (Equation 6):

$$n_{N_2O_{total}}^t = x_{N_2O}^t \cdot P^t \cdot \left( \frac{V_g}{R \cdot T} + H_{N_2O} \cdot V_l \right) \quad \text{Equation 6}$$

Where  $n_{N_2O_{total}}^t$  is the total amount of N<sub>2</sub>O produced in the Duran bottle at time t (moles N<sub>2</sub>O);  $x_{N_2O}^t$  is the molar fraction of N<sub>2</sub>O in the gas phase at time t (mol N<sub>2</sub>O·mol gas<sup>-1</sup>); P<sup>t</sup> is the pressure in the gas headspace at time t (typically 101325 Pa unless otherwise stated); V<sub>g</sub> is the volume of gas in the bottle/serum flask (mL); V<sub>l</sub> is the volume of liquid in the bottle/serum flask (mL); R is the ideal gas constant (8.314 J·mol<sup>-1</sup>·K<sup>-1</sup>); T is the temperature inside the bottle (298.15 K); and H<sub>N<sub>2</sub>O</sub> is the Henry law constant of N<sub>2</sub>O at T (2.5·10<sup>-7</sup> mol·L<sup>-1</sup>·Pa<sup>-1</sup>). The pressure changes and the N<sub>2</sub>O losses caused by sampling were accounted for as described by Guieysse et al. (2013) and Plouviez (2017).

Emissions were computed as described by Plouviez et al. (2019b) using Equation 7:

$$\phi_{N_2O} = ([N_2O]^* - [N_2O]) \cdot Kl_{N_2O} \quad \text{Equation 7}$$

Where  $\phi_{N_2O}$  represents the N<sub>2</sub>O flux (mol·m<sup>-2</sup>·d<sup>-1</sup>), [N<sub>2</sub>O]<sup>\*</sup> represents the dissolved N<sub>2</sub>O concentration at equilibrium with earth atmosphere (1.3·10<sup>-8</sup> mol·L<sub>water</sub><sup>-1</sup>), [N<sub>2</sub>O] is the mean dissolved N<sub>2</sub>O concentration in our samples (mol·L<sub>water</sub><sup>-1</sup>) and Kl<sub>N<sub>2</sub>O</sub> is the N<sub>2</sub>O mass transfer coefficient (m·d<sup>-1</sup>).

The field value of  $Kl_{N_2O}$  was estimated from the mass transfer coefficient for oxygen at 20 °C and corrected for temperature based on the diffusivities of these gases and a temperature correction factor of 1.024 (Equation 8):

$$Kl_{N_2O} = 1.024^{T-20} \cdot Kl_{O_2} \cdot \sqrt{\frac{DF_{N_2O}}{DF_{O_2}}} \quad \text{Equation 8}$$

Where  $Kl_{N_2O}$  is the N<sub>2</sub>O mass transfer coefficient (m·d<sup>-1</sup>), T is the temperature of the liquid (°C),  $Kl_{O_2}$  is the O<sub>2</sub> mass transfer coefficient (m·d<sup>-1</sup>),  $DF_{N_2O}$  and  $DF_{O_2}$  are the molecular diffusion coefficients in water at 20°C of N<sub>2</sub>O and O<sub>2</sub> (1.98·10<sup>-9</sup> m<sup>2</sup>·s<sup>-1</sup> and 1.84·10<sup>-9</sup> m<sup>2</sup>·s<sup>-1</sup>, respectively).

The field value of  $Kl_{O_2}$  was calculated according to Schwarzenbach et al. (2005) calculations as (Equation 9):

$$Kl_{O_2} = (4 \cdot 10^{-4} + u^2 \cdot 4 \cdot 10^{-5}) \cdot 86,400 \cdot 0.01 \quad \text{Equation 9}$$

Where  $Kl_{O_2}$  is the O<sub>2</sub> mass transfer coefficient (m·d<sup>-1</sup>), u (m·s<sup>-1</sup>) is the average surface wind speed, 86,400 is used to convert the secondly rate into daily values (s·d<sup>-1</sup>) and 0.01 is used to convert the values from cm to m (m·cm<sup>-1</sup>).

### 5.6.6. RNA extraction from axenic *Chlamydomonas* cultures and qPCR

Cells were harvested and concentrated prior to extracting RNA from the samples. Four different harvest methods were tested to optimise the cell extract preparation protocol (Appendix H). This step is crucial as it determines the RNA concentration of the sample: RNA must be preserved, and sampling volume must be kept low enough to avoid large volume changes over time.

RNA extraction from axenic *Chlamydomonas* cultures was done following the protocol proposed by Sanz-Luque & Montaigu (2018). Briefly, RNA was extracted from the cells sample using a phenol (pH 4.5):chloroform:isoamyl alcohol solution and precipitated using a solution of lithium chloride (8 M). The contaminating genomic DNA was removed from the RNA extract using a kit of DNase I recombinant, RNase-free (Roche, Mannheim, Germany) according to the manufacturer's instructions. Once the RNA was purified, cDNA was synthesised from our RNA extract by RT-qPCR using the iScript™ cDNA Synthesis Kit (Biorad, Hercules, California, USA) following the manufacturer's instructions. The real time qPCR was then carried out using the reagent SsoFast™ EvaGreen® Supermix (Biorad, Hercules, California, USA) as stated in the manufacturer's manual. Genes *CBLP*, *UBIQUITIN* and *ACT1* were used as reference genes. The quality of the RNA sample was estimated by calculating the ratio



of the absorbance at 260 nm and 280 nm (260/280). Using this protocol, gene expression was quantified under various conditions and times (e.g. cells grown in HSM – N and in HSM – N + 10 mM  $\text{NO}_2^-$  after 0 h, 1 h, 3 h, 6 h, 17 h and 24 h of experiment).

### 5.6.7. Electrophoresis and DNA sequencing

Electrophoresis was used to verify RNA integrity after extraction and separate the DNA produced during primers optimisation. DNA sequencing was used to check that amplification worked properly (no unspecific amplifications).

The nucleic acids were separated using agarose gels containing 1%-2% agarose (w/v) in TAE buffer (4.84 g·L<sup>-1</sup> Tris; 1.1 mL·L<sup>-1</sup> acetic acid; 1 mM EDTA pH 8.0). Ethidium bromide at 0.5 µg·mL<sup>-1</sup> was added to the gel to dye the nucleic acids and allow detection using a UV-transilluminator. Prior to loading the samples in the gel, 2 µL of loading buffer (50% glycerol; 0.25% bromophenol blue; 0.25% xylene cyanol) was added to the samples and the wells were then filled with 15 µL of the mixture. As molecular weight marker, 10 µL of Invitrogen 1 kb Ladder Plus marker (ThermoFisher Scientific) was used in the first and the last well of the gel.

Once the migrations on the gel were over, nucleic acids were visualized using UV light and data analysed with the image system Gel Doc 2000 (BioRad). For the RNA, two discrete bands of ribosomal RNA were indicative of no significant degradation. For the DNA, the observed bands corresponding to DNA fragments of interest were extracted and purified from the gel by cutting the gel with a scalpel and then using the Ultrafree<sup>®</sup>-DA (Millipore) centrifugal filter units according to the manufacturer instructions. The DNA samples were sequenced using Sanger technology by the centralized research support service SCAI (Servicio Central de Apoyo a la Investigacion) at the University of Cordoba (Spain).

## 6. Results and discussion

### 6.1. N<sub>2</sub>O emissions from New Zealand lakes

Due to COVID-19, the New Zealand border was closed which delayed my arrival in NZ (March 2021). As a consequence, weekly sampling at Lake Horowhenua and Turitea reservoir only started in April and October 2021, respectively. As shown in Figure 11, algal blooms still occurred during autumn in lake Horowhenua.



Figure 11: Lake Horowhenua in June 2021 (left) and January 2018 (right).

Table 16 summarizes the experimental evidence gathered between April and October 2021. At this time, there is no clear relationship between chlorophyll concentration and dissolved N<sub>2</sub>O concentration. N<sub>2</sub>O was however synthesized in the presence of O<sub>2</sub>.

Based on the average of the mean values of chlorophyll, dissolved nitrite, dissolved nitrate and phosphate (dissolved phosphorus) concentrations, we calculated a TLI score of 7.5 for lake Horowhenua and 4.8 for the Turitea reservoir. Based on longer monitoring performed by the team in New Zealand, the value have been recalculated and indicate that Lake Horowhenua (12 months of data) is eutrophic (TLI score of 5.0) and the Turitea reservoir (6 months of data) is oligotrophic (TLI score of 2.2). The latter results correspond to historical data for lake Horowhenua provided in LAWA (*Lake Horowhenua Water Quality, 2021*).

The N<sub>2</sub>O emissions were estimated at 0.4 – 8.7 g N-N<sub>2</sub>O·ha<sup>-1</sup>·yr<sup>-1</sup> (n = 29 sampling events) between April and October 2021 for lake Horowhenua, which is low and can be explained by a lower biological activity due to lower temperature and light during winter. However, data from lake Horowhenua were already gathered in 2018 and presented at the IWA Conference of 2019 on Algal Technologies and

Stabilization Ponds for Wastewater Treatment and Resource Recovery. N<sub>2</sub>O emissions of Lake Horowhenua during Summer 2018 were thus estimated at 0.6 – 24.2 g N-N<sub>2</sub>O·ha<sup>-1</sup>·yr<sup>-1</sup> (n = 3 sampling events). The N<sub>2</sub>O emissions from the Turitea reservoir were estimated at 0.3 – 1.0 g N-N<sub>2</sub>O·ha<sup>-1</sup>·yr<sup>-1</sup> (n = 3 sampling events) between August and October 2021. These results show that the N<sub>2</sub>O emissions were higher at the eutrophic lake compared to the oligotrophic reservoir, which is consistent with our hypothesis. Also, these results tally with the observation from the IPCC that nutrient-impacted waters emit more N<sub>2</sub>O than clean waters. Besides, the recorded emissions between April and October 2021 (0 – 3.5 nmole·m<sup>-2</sup>·h<sup>-1</sup>) are in the lower range of the fluxes reported in the literature for lakes (Table 2, fluxes ranged from 0 to 10,057 nmole·m<sup>-2</sup>·h<sup>-1</sup>).

Lakes represent around 1.3% of New Zealand area (*Land Cover, 2021*) and around half of New Zealand monitored lakes are eutrophic or worse (Ministry for the Environment & Stats NZ, 2020), meaning that 167,500 ha of lakes are eutrophic. Based on our data from lake Horowhenua, we estimated that national N<sub>2</sub>O emissions from New Zealand eutrophic lakes represent 0.07 to 4.1 t N-N<sub>2</sub>O·yr<sup>-1</sup>. This corresponds to 0.0004 – 0.02% of New Zealand agricultural N<sub>2</sub>O emissions in 2020 (Ministry for the Environment, 2022) which is the main sector responsible for N<sub>2</sub>O emissions in New Zealand. Thus, emissions from eutrophic lakes represent a small portion of New Zealand N<sub>2</sub>O emissions.

Based on the data from lake Horowhenua, global N<sub>2</sub>O emissions from eutrophic lakes and reservoirs, which represent 126,000,000 ha, were estimated at 25 – 1,500 t N-N<sub>2</sub>O·yr<sup>-1</sup>. These emissions represent 0.008 to 0.3% of global anthropogenic N<sub>2</sub>O emissions from inland and coastal waters (Tian et al., 2020). The global contribution of eutrophic lakes to N<sub>2</sub>O emissions is thus low.

To conclude, the N<sub>2</sub>O emissions from eutrophic lakes reported in this study are low and therefore the national and global significances are low. However, the measured fluxes in this study were in the lower range of fluxes reported in the literature. Thus further research is required to understand which parameters (e.g. type of lake, nutrients, etc.) could explain the variability of fluxes reported in different studies.

Table 16: First sampling data from Lake Horowhenua (LH) and Turitea reservoir (TR).  
The mean values are based on triplicates.

Lake	Date	Time	Radiation (MJ·m <sup>-2</sup> )	Distance from the lake edge (m)	Water temperature	Air temperature (°C)	Conductivity (µS·cm <sup>-1</sup> )	Dissolved oxygen (mg·L <sup>-1</sup> )	pH	Chlorophyll (mg·L <sup>-1</sup> )		Total suspended solid (g·L <sup>-1</sup> )		N <sub>2</sub> O (nmol N <sub>2</sub> O·L water <sup>-1</sup> )		NO <sub>2</sub> <sup>-</sup> (µg·L <sup>-1</sup> )	NO <sub>3</sub> <sup>-</sup> (mg·L <sup>-1</sup> )	PO <sub>4</sub> <sup>3-</sup> (µg·L <sup>-1</sup> )		Chloride (mg·L <sup>-1</sup> )		Sulfate (mg·L <sup>-1</sup> )	
										Mean	Error	Mean	Error	Mean	Error			Mean	Error	Mean	Error	Mean	Error
LH	27/04/2021	10:10	1.09	15		15	220.4		7.88	1.38	0.10	0.23	0.01					520.0	22.9	27.8		8.5	
LH	06/05/2021	10:30	0.725	15	17.3	17	234.3	9.98	8														
				2	17.7		239.7	5.34	8.1	3.83	0.09	1.39	0.01					228.6	1.0	26.6		9.3	
LH	13/05/2021	10:40	1.48	15	12.1	12	246.4	6.39	7.41	0.05	0.03	0.01	0.00					208.7	2.0	24.6		7.1	
				5	12.9		250.3	7.03	7.56	0.47	0.04	0.21	0.02					120.6	1.5	26.7		7.2	
LH	19/05/2021	10:30	1.13	15	13.5	16	239.1	10.01	7.93	0.05	0.01	0.03	0.00			52.7	2.0	89.5	1.1	26.4		10.1	
				5	14.2		238.5	10.5	8.09	0.03	0.01	0.03	0.00			109.5	1.3	57.6	2.2	24.5		7.1	
LH	27/05/2021	10:30	1.095	15		12	262.3	10.78	7.58	0.06	0.00	0.09	0.01			23.6	1.2	66.8	1.2	26.5		10.8	
				5			255.2	12.65	7.65	0.20	0.01	0.16	0.01			59.6	1.2	88.9	0.7	28.1		13.2	
LH	03/06/2021	10:30	0.695	15	10.5	10	252.5	10.93	7.4	0.12	0.01	0.07	0.00			128.3	2.9	48.3	9.8	28.4		17.2	
				5	10.5		259.6	11.62	8.06	0.83	0.02	0.25	0.04			47.2	0.9	34.8	0.7	30.0		18.3	
LH	10/06/2021	10:30	0.535	15	12.6	14	235.7	10.81	7.54	0.11	0.00	0.02	0.00	25.6	2.7	86.9	2.0	34.9	5.4	28.7		21.2	
				5	12.6		242.4	15.42	8.53	2.75	0.05	0.55	0.02	16.3	5.2	61.0	1.1	57.6	2.2	29.4		20.2	
LH	16/06/2021	10:30	0.525	15	14.4	15	0.215	4.94	7.17	0.02	0.00	0.01	0.00	10.7	2.1	50.7	1.0	47.4	2.6	30.1	0.2	19.7	0.2
				5	14.5		284.5	7.19	7.14	0.08	0.01	0.04	0.01	4.7	0.5	54.8	0.4	54.8	5.3	34.4	0.2	18.6	0.1
LH	23/06/2021	10:30	0.72	15	7.6	7	235.4	8.75	6.83	0.03	0.00	0.01	0.00	7.7	0.5	16.4	1.1	28.5	0.4	27.1	0.1	20.3	0.1
				5	7.9		240.3	9.3	7.09	0.08	0.00	0.07	0.01	7.1	10.3	39.5	1.7	23.7	2.0	27.8	0.0	20.3	0.0
LH	01/07/2021	10:30	0.945	15	7.7	9	223.7	9.62	7.89					15.2	0.7	69.1	3.3	54.0	7.1	25.9	0.0	20.6	0.0
				5	7.9		222	9.77	7.74					18.1	3.2	10.4	3.2	52.2	1.5	25.7	0.1	20.1	0.1
LH	09/07/2021	10:30	0.975	15	7.6	9	228.4	9.58	6.73	0.04	0.00	0.02	0.00	22.5	0.3	39.3	4.4	40.1	0.8	26.0	0.1	22.9	0.0
				5	7.6		226.7	9.72	7.18	0.05	0.00	0.02	0.00	12.2	7.9	41.9	5.2	30.8	0.8	26.3	0.2	23.1	0.2
LH	14/07/2021	10:30	0.625	15	6.1	6	240.6	10.71	6.6	0.08	0.00	0.08	0.00	23.4	2.1	36.6	6.2	41.7	0.2	27.4	0.1	25.8	0.1
				5	6		243.4	11.52	7.07	0.13	0.00	0.14	0.00	17.3	1.8	43.6	5.8	46.0	0.3	28.1	0.1	25.5	0.1
LH	22/07/2021	10:30	0.44	15	11.9	13	227.5	8.87	8	0.01	0.00	0.02	0.00	12.2	1	106.6	2.4			29.7	0.1	19.3	0.1
				5	12		211.9	8.85	7.57	0.02	0.00	0.03	0.01	12.8	0.7	17.5	1.4			28.5	0.0	18.0	0.1
LH	30/07/2021	10:30	0.965	15	11.4	14	226.6	8.69	6.7	0.01	0.00	0.01	0.00	32.4	6.2	94.2	7.3			27.0	0.0	24.1	0.0
				5	11.9		228.2	8.59	6.82	0.02	0.00	0.01	0.00	37.2	2.9	56.7	6.9			27.1	0.0	23.8	0.1
LH	05/08/2021	13:00	1.02	15		10	226.8	10.9	6.89	0.04	0.00	0.07	0.01	25.8	1.7	25.6	9.7			27.6	0.0	23.7	0.0
				5			238.3	10.23	6.75	0.07	0.00	0.19	0.00	34.8	2.2	93.0	10.1			29.4	0.0	23.6	0.1
LH	12/08/2021	10:30	0.47	15	10.8	13	219.7	9.82	7.73	0.03	0.00	0.04	0.00	69.5	4.3	46.0	9.8			27.6	0.1	23.9	0.0
				5	10.4		227.8	9.78	8.07	0.03	0.00	0.01	0.00	65.5	4.5	37.3	9.2			27.3	0.0	23.7	0.0
TR	16/08/2021	10:00	1.19	5	9	12	93.4	11.2	6.15	0.02	0.00	0.00	0.00	7.1	1.8	16.7	1.0			17.9	0.0	4.2	0.0
LH	15/10/2021	11:00	1.51	15	15.5	14.2	215.2		7.1	0.03	0.00	0.01	0.00	16.2	2	37.8	9.7						
				5	15.2		225.3		7.28	0.04	0.00	0.04	0.01	14.4	0.8	146.1	4.5						
TR	15/10/2021	12:40	1.69	5	13	14.5	88.15		6.1	0.01	0.00	0.00	0.00	6.3	1	19.8	0.4						
LH	21/10/2021	11:00	1.77	15	17.5	15.8	212.3	10.24	7.28	0.02	0.00	0.01	0.00	12.5	1.4	59.3	6.8						
				5	17.3		214.5	10.62	7.35	0.02	0.00	0.01	0.00	13.8	0.8	86.1	5.8						
LH	27/10/2021	10:30	2.05	15	21.2	17.5	212.7	9.8	7.04	0.03	0.00	0.01	0.00	7.3	1.6	256.8	5.4						
				5	22.1		212.8	10.59	7.3	0.03	0.00	0.01	0.00	7.4	0.2	62.9	3.2						
TR	27/10/2021	12:30	2.15	5	17.8	19	87.5	10.4	6.73	0.01	0.00	0.00	0.00	3.2	0.4	8.1	0.0						

## 6.2. Physiological conditions leading to N<sub>2</sub>O synthesis in *Chlamydomonas*.

CYP55 has been proposed as the main enzyme involved in N<sub>2</sub>O production by *Chlamydomonas* in the dark (Plouviez et al., 2017a; Burlacot et al., 2020) while FLVs could be responsible for N<sub>2</sub>O production in the light (Burlacot et al., 2020). HCPs, have been evidenced to catalyze N<sub>2</sub>O synthesis in *E. coli* (a bacterium) under anaerobic conditions (Wang et al., 2016) but their putative N<sub>2</sub>O producing roles in *Chlamydomonas* have not been investigated yet.

The physiological conditions associated with the production of N<sub>2</sub>O by the model organism *Chlamydomonas reinhardtii* are still not well known. Nitrite has been identified as a precursor of N<sub>2</sub>O production as nitrite can be reduced into nitric oxide (NO) intracellularly, and NO is in turn reduced to N<sub>2</sub>O. Laboratory-scale experiments were therefore conducted to study the N<sub>2</sub>O synthesis by different axenic cultures of *Chlamydomonas* under different nutrient and light conditions.

Because the accumulation of NADPH and ATP blocks the electron flow and leads to the production of reactive oxygen species with detrimental effects when cells cannot use excess light; we hypothesised that *Chlamydomonas* cells produce N<sub>2</sub>O to maintain the electron flow and protect cells against the impact of light excess. Specifically, we believe that the production of N<sub>2</sub>O from NO is a mechanism to divert energy excess, possibly supported by FLVs (A and B) (Figure 12). To validate this hypothesis, we chemically inhibited the photosynthetic electron flow coming from PSII (using DCMU) in *Chlamydomonas* wildtype. In addition, FLV-deficient mutants were used to assess the role of FLVs in N<sub>2</sub>O production.

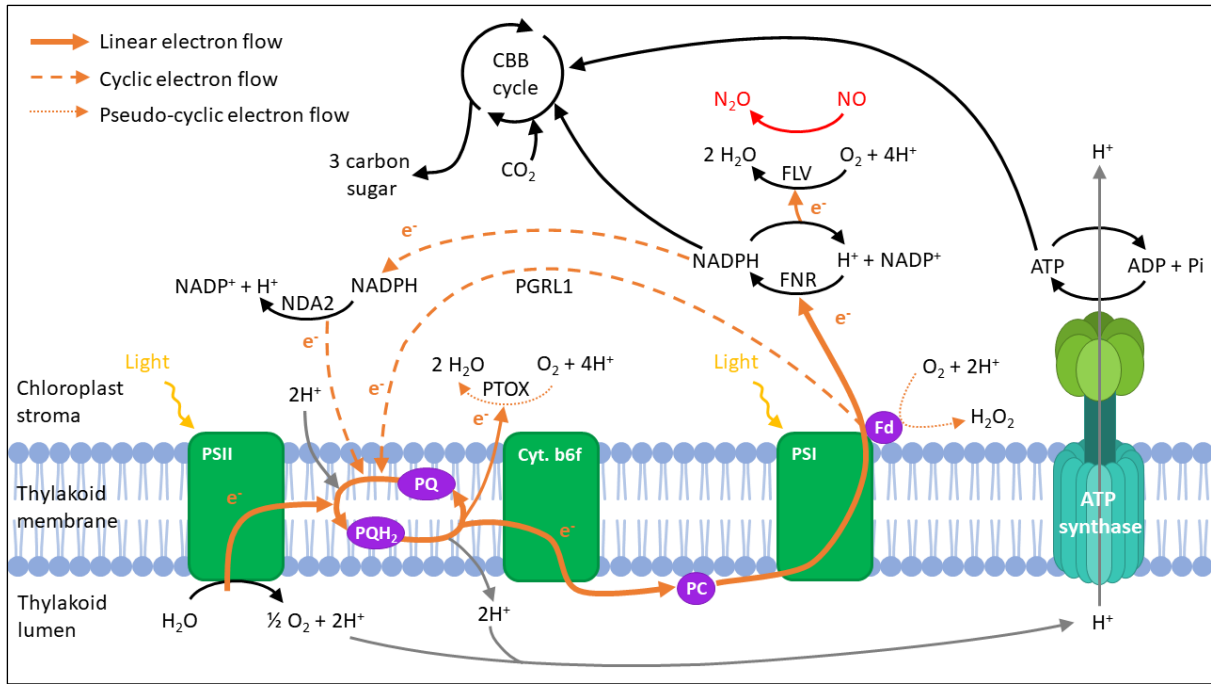
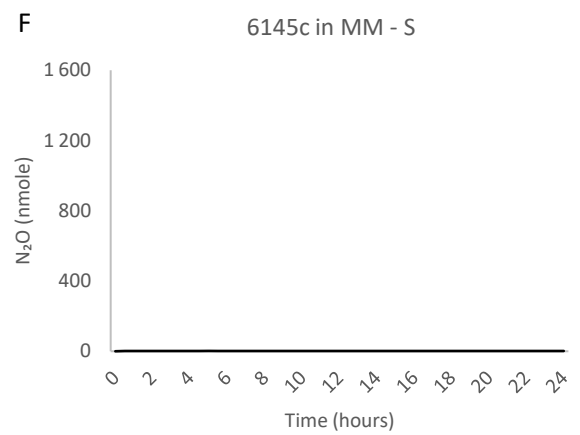
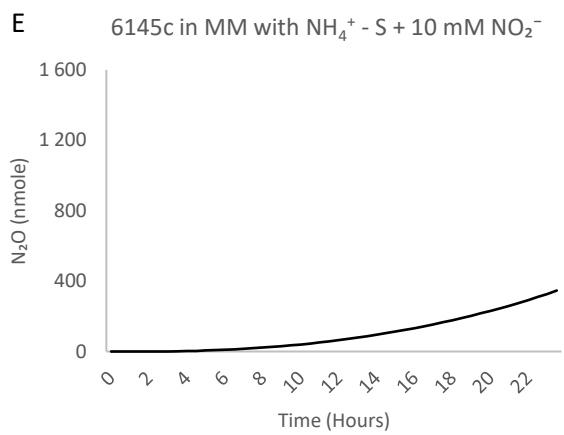
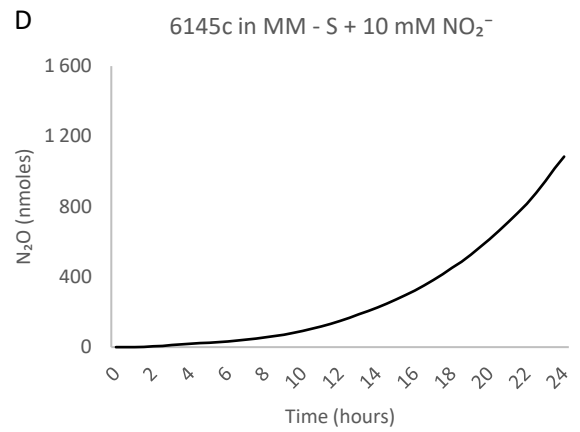
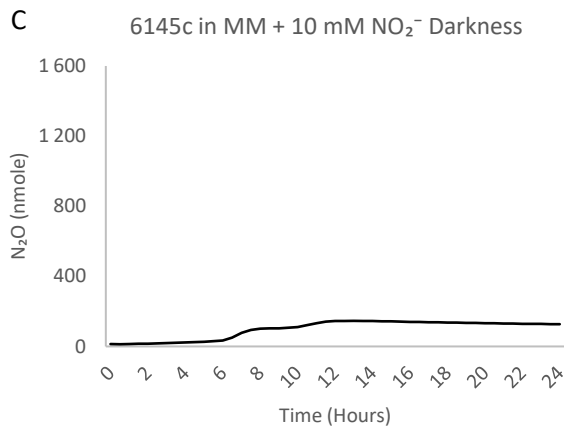
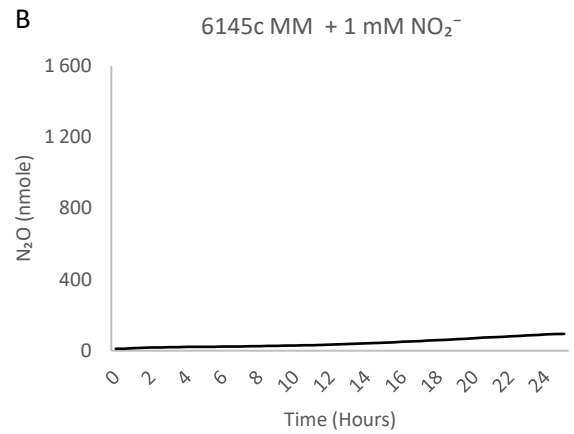
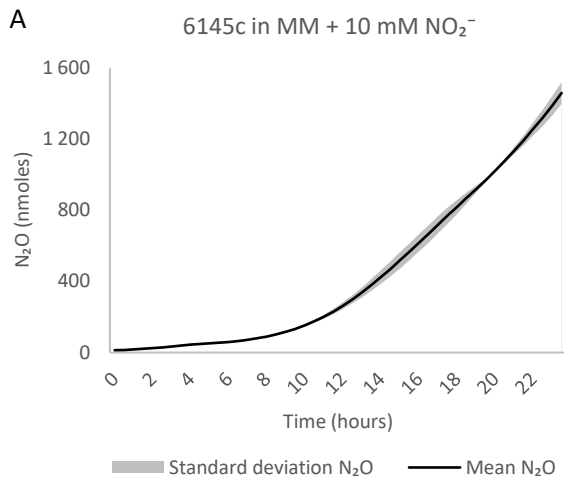


Figure 12: Oxygenic photosynthesis pathway in *Chlamydomonas reinhardtii*.

Photosynthesis enables the conversion of light energy into chemical energy to fix carbon. The photolysis of water releases electrons that are transported through the thylakoid membrane and used to generate reducing power (NADPH) and high-energy orthophosphate bonds (ATP). This energy drives the fixation of  $\text{CO}_2$  in the Calvin-Benson-Bassham cycle. Environmental light and/or nutrient fluctuations can trigger the production of reactive oxygen species which are detrimental to the photosynthetic apparatus. To acclimate to these fluctuations, photosynthetic organisms possess different alternative electrons routes to dissipate the excess energy (e.g. NDA2, PTOX). Here, we conjecture that the reduction of NO into  $\text{N}_2\text{O}$  by the FLVs (in red) could be an alternative pathway to divert the excess energy received by the microalga. Abbreviations/symbols:  $\text{H}_2\text{O}$ , water;  $\text{O}_2$ , dioxygen;  $\text{H}_2\text{O}_2$ , hydrogen peroxide; NO, nitric oxide;  $\text{N}_2\text{O}$ , nitrous oxide;  $\text{CO}_2$ , carbon dioxide;  $\text{H}^+$ , hydron;  $e^-$ , electron; PSII, photosystem II; PQ, plastoquinone;  $\text{PQH}_2$ , plastoquinol; Cyt. b6f, cytochrome b6f; PC, plastocyanine; PSI, photosystem I; Fd, ferredoxine; ATP, adenosine triphosphate; ADP, adenosine diphosphate; Pi, inorganic phosphate;  $\text{NADP}^+/\text{NADPH}$ , nicotinamide adenine dinucleotide phosphate; FNR, ferredoxin  $\text{NADP}^+$  reductase; FLV, flavodiiron protein; PGRL1, proton gradient regulation 5-like photosynthetic phenotype 1; PTOX, plastid terminal oxidase; NDA2, type II  $\text{NADPH}$  dehydrogenase; CBB cycle, Calvin-Benson-Bassham cycle.

The wild type strain 6145c was used as a control as it is the parental strain of the M3 mutant. The M3 mutant, which is deficient in nitrite reductase, was selected because it produces more  $\text{N}_2\text{O}$  than the parental strain 6145c (Plouviez et al., 2017a). This strain cannot convert nitrite into ammonium and, therefore, secretes most of the nitrite to the medium. This nitrite is being continuously taken up and excreted and can also be constantly reduced into NO, the substrate for  $\text{N}_2\text{O}$  synthesis, allowing us to use this mutant as a tool to investigate and enhance the impact of different conditions and molecular players on the  $\text{N}_2\text{O}$  production. We also used the "WT" strain CMJ030 (cc-4533) as a control because it is the parental strain for the *flv* and *cyp55* mutants.

### 6.2.1. N<sub>2</sub>O synthesis by the wild type *Chlamydomonas 6145c*



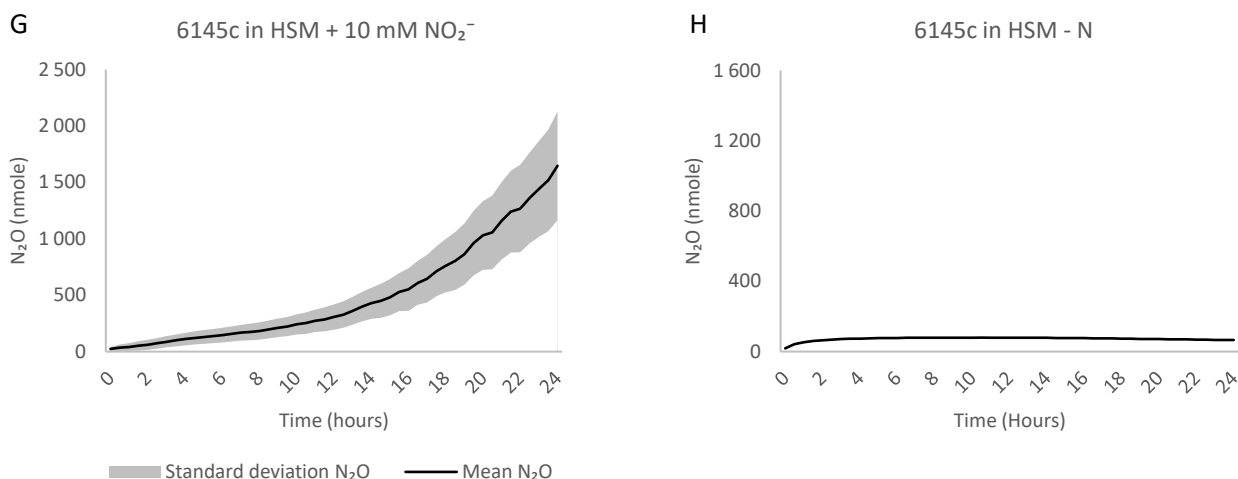


Figure 13: N<sub>2</sub>O batch test with *C. reinhardtii* 6145c (10 mg chlorophyll·L<sup>-1</sup> ~ 0.25 g DCW·L<sup>-1</sup>). 6145c in MM + 10 mM NO<sub>2</sub><sup>-</sup>, triplicate (A); 6145c MM + 1 mM NO<sub>2</sub><sup>-</sup> (B); 6145c in MM + 10 mM NO<sub>2</sub><sup>-</sup> in the dark (C); 6145c in MM - S + 10 mM NO<sub>2</sub><sup>-</sup> (D); 6145c in MM - S + 10 mM NO<sub>2</sub><sup>-</sup> (E); 6145c in MM - S (F); 6145c in HSM + 10 mM NO<sub>2</sub><sup>-</sup>, triplicate (G); 6145c in HSM - N, control (H). The graphs A to F come from another set-up in which the cells were grown in a temperate room without windows at 25°C under continuous light provided by four 300×300×40 mm LED panels of 30 W, with warm white light (3200 K) at a light irradiance around 145 μmol photons·s<sup>-1</sup>·m<sup>-2</sup>.

First, results from an initial batch assay experiments confirmed that N<sub>2</sub>O emissions had a biological origin in our set-up and required nitrite to be significant (see abiotic controls in Appendix I). These results, presented in Figure 13, also showed that axenic cultures of *Chlamydomonas* 6145c synthesised N<sub>2</sub>O mainly under illumination when supplied NO<sub>2</sub><sup>-</sup> in autotrophic conditions. N<sub>2</sub>O total production over 24 h was lower when NO<sub>2</sub><sup>-</sup> supply was decreased (Figure 13 A and B). We thus recorded a production of 31.5 ± 10.7 μmole N<sub>2</sub>O·g DW<sup>-1</sup> after 24 h when cells were grown autotrophically and supplied with 10 mM NO<sub>2</sub><sup>-</sup>. The N<sub>2</sub>O specific production rate between 4 and 24 h after nitrogen addition was 1215 nmol N<sub>2</sub>O·g DW<sup>-1</sup>·h<sup>-1</sup>. These values are similar to *Chlorella vulgaris* production of 37.9 μmole N<sub>2</sub>O·g DW<sup>-1</sup> reported under similar conditions by Weathers (1984), who also recorded a production rate of 1580 nmol N<sub>2</sub>O·g DW<sup>-1</sup>·h<sup>-1</sup> when supplying 33 mM NO<sub>2</sub><sup>-</sup> (Weathers, 1984). However, our production rate was one order of magnitude higher than *C. vulgaris* production of 109 nmol N<sub>2</sub>O·g DW<sup>-1</sup>·h<sup>-1</sup> when supplying 12 mM NO<sub>2</sub><sup>-</sup> under similar conditions (Guieysse et al., 2013). The discrepancy between the results from Weathers (1984) and Guieysse et al. (2013) might be due to a difference of light intensity during these experiments and/or a difference in the genetic background of the *C. vulgaris* strains used. There is therefore a possibility that the genetic background plays an important role in N<sub>2</sub>O production as we can observe a high variation of the production in our experiments between the different *Chlamydomonas* wild type strains (6145c, CMJ030 and 21gr) cultured in the same conditions (see sections 6.2.1, 6.2.3 and 6.2.4).



As can be seen in Figure 13 A and C, N<sub>2</sub>O production was significantly lower in the dark than in the light. An opposite trend was observed when other green algae species were grown mixotrophically (Weathers, 1984). This could be explained by a variation in oxygen level when a carbon source is available. We also suggest that in autotrophic conditions N<sub>2</sub>O production is triggered to get rid of the excess energy received when no carbon source is available.

To assess if a different nutrient stress could trigger similar N<sub>2</sub>O production, we also cultured the cells in a medium deprived of sulphur under autotrophic conditions in the light. As seen in Figure 13 D, E and F, in the absence of a usable N source (ammonium), the N<sub>2</sub>O total production after 24 h was slightly lower when cells were deprived S than when this nutrient was supplied. However, N<sub>2</sub>O production was much lower when NH<sub>4</sub><sup>+</sup> was available (Figure 13 E) than when this N source was added (Figure 13 D). This difference could be explained by the fact that NO<sub>3</sub><sup>-</sup> and NO<sub>2</sub><sup>-</sup> assimilation are inhibited in presence of NH<sub>4</sub><sup>+</sup> in *Chlamydomonas* (Sanz-Luque et al., 2015a).

For each experiment conducted in the light, the CO<sub>2</sub> present in the bottle was totally consumed after 2 to 3 h of experiments. Under CO<sub>2</sub> depletion, the light energy cannot be used for CO<sub>2</sub> fixation and the cells need to get rid of the excess NADPH by activating alternative pathways. N<sub>2</sub>O production, which consumes NADPH (Burlacot et al., 2020), could be one of the strategies used under these conditions.

### 6.2.2. N<sub>2</sub>O synthesis by a *Chlamydomonas* NiR knockout mutant (M3)

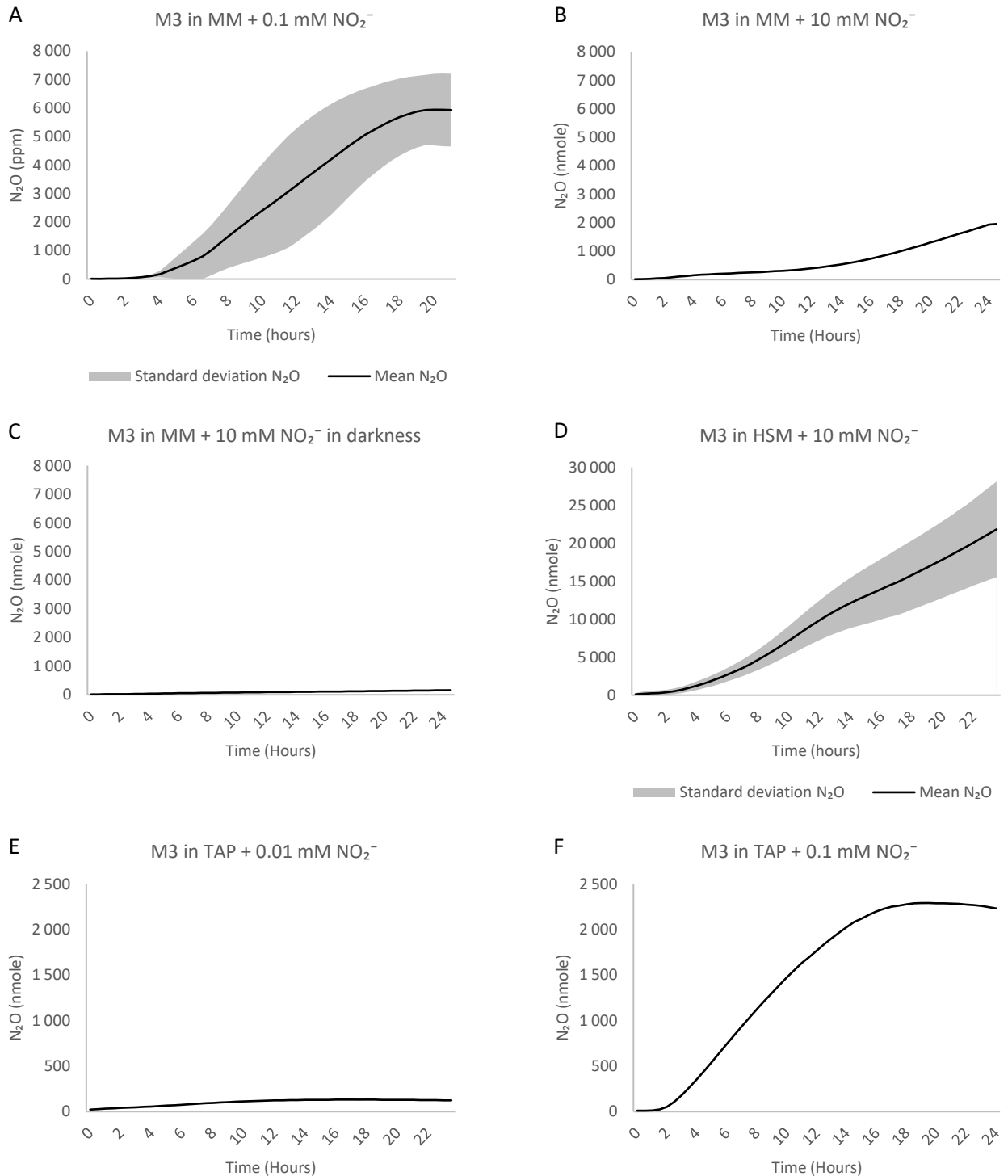


Figure 14: N<sub>2</sub>O batch test with *C. reinhardtii* M3 (10 mg chlorophyll·L<sup>-1</sup> ~ 0.25 g DCW·L<sup>-1</sup>). M3 in MM + 0.1 mM NO<sub>2</sub><sup>-</sup>, triplicate (A); M3 in MM + 10 mM NO<sub>2</sub><sup>-</sup> (B); M3 in MM + 10 mM NO<sub>2</sub><sup>-</sup> in the dark (C); M3 in HSM + 10 mM NO<sub>2</sub><sup>-</sup>, triplicate (D); M3 in TAP + 0.01 mM NO<sub>2</sub><sup>-</sup> (E); M3 in TAP + 0.1 mM NO<sub>2</sub><sup>-</sup> (F). The graphs A to C come from another set-up in which the cells were grown in a temperate room without windows at 25°C under continuous light provided by four 300×300×40 mm LED panels of 30 W, with warm white light (3200 K) at a light irradiance around 145 μmol photons·s<sup>-1</sup>·m<sup>-2</sup>.

The M3 *Chlamydomonas* strain is a NiR deficient mutant that cannot reduce  $\text{NO}_2^-$ . Thus, nitrite supply to the medium causes this substrate to be continuously taken and secreted by M3 cells. As  $\text{NO}_2^-$ , a key precursor to  $\text{N}_2\text{O}$ , accumulates intracellularly,  $\text{N}_2\text{O}$  production is higher in the nitrite-accumulating M3 than in the nitrite-reducing wild type strain (Figure 14 D) under illumination. In darkness however,  $\text{N}_2\text{O}$  production in M3 was similar to the wild type (Figure 14 C). This was expected as nitrite reduction into NO and the subsequent synthesis of  $\text{N}_2\text{O}$  is much lower in the dark than in the light (Figure 14 B and C), as reported by Plouviez et al. (2017a).

We can observe a high variability in the production of  $\text{N}_2\text{O}$  when supplying 10 mM  $\text{NO}_2^-$ . However, our results must be analysed with care considering that the spectroscope loses accuracy when the production is higher than 8,000 nmoles of  $\text{N}_2\text{O}$ . It was nevertheless necessary to verify that the variability was not due to experimental error. During our experiments, we switched from minimal medium (MM) to high salt medium (HSM) because the use of MM frequently caused the formation of precipitates after autoclaving. The only difference between the MM and HSM media is the calcium concentration which is five times lower in HSM compared to MM. A potential biological explanation to the lower production in MM (Figure 14 B and D) could therefore be that calcium is an inducer of non-photochemical quenching (i.e. the loss of excess energy via heat) (Petroustos et al., 2011). Under conditions of low  $\text{Ca}^{2+}$  supply, cells may need to quench more energy through pathways such as  $\text{N}_2\text{O}$  synthesis, explaining the high emission recorded in HSM.

With these experiments, we showed that microalgal  $\text{N}_2\text{O}$  synthesis varied with physiological conditions of the cell. In addition to the mixotrophic conditions leading to  $\text{N}_2\text{O}$  synthesis in the light and in the dark (Weathers, 1984; Guieysse et al., 2013; Plouviez, 2017), we showed that *Chlamydomonas* produced  $\text{N}_2\text{O}$  in autotrophic conditions under illumination. These results are consistent with the results obtained by Guieysse et al. (2013) with *Chlorella vulgaris*.

Once we identified the conditions under which *Chlamydomonas* was consistently producing  $\text{N}_2\text{O}$ , we could further investigate the  $\text{N}_2\text{O}$  synthesis pathways in this microalga.

### 6.2.3. N<sub>2</sub>O synthesis in CMJ030, *flvb*, and *cyp55* strains

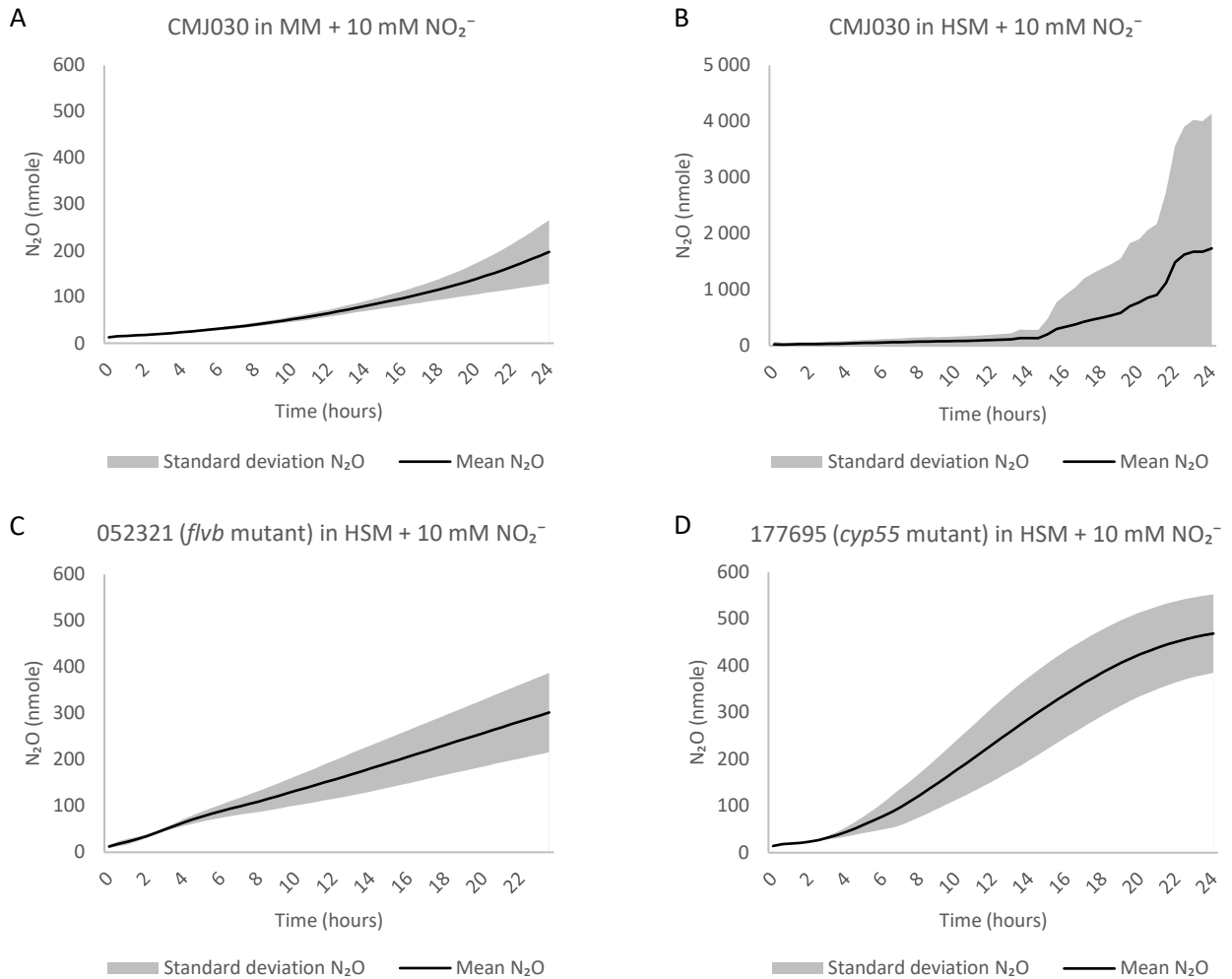


Figure 15: N<sub>2</sub>O batch test with *C. reinhardtii* CMJ030, LMJ.RY0402.052321 and LMJ.RY0402.177695 (10 mg chlorophyll·L<sup>-1</sup> ~ 0.25 g DCW·L<sup>-1</sup>). CMJ030 in MM + 10 mM NO<sub>2</sub><sup>-</sup>, triplicate (A); CMJ030 in HSM + 10 mM NO<sub>2</sub><sup>-</sup>, triplicate (B); LMJ.RY0402.052321 in HSM + 10 mM NO<sub>2</sub><sup>-</sup>, triplicate (C); LMJ.RY0402.177695 in HSM + 10 mM NO<sub>2</sub><sup>-</sup>, triplicate (D). The graph A comes from another set-up in which the cells were grown in a temperate room without windows at 25°C under continuous light provided by four 300×300×40 mm LED panels of 30 W, with warm white light (3200 K) at a light irradiance around 145 μmol photons·s<sup>-1</sup>·m<sup>-2</sup>.

In order to confirm the known conditions during which FLV and CYP55 are involved in N<sub>2</sub>O production (i.e. light and dark conditions, respectively), we obtained the previously described *flvb* and *cyp55* mutants (Burlacot et al., 2020) from the *Chlamydomonas* resource center ([www.chlamycollection.org](http://www.chlamycollection.org)). The strain CMJ030 was used as control because this wildtype has the same genetic background as the *flvb* and *cyp55* mutants. Unfortunately, as shown in Figure 15 B, N<sub>2</sub>O production by CMJ030 was highly variable and the experiments were not reproducible in HSM. Optimization is therefore required to determine why we are observing such variability. Nevertheless, the FLVs-deficient mutant produced a significant amount of N<sub>2</sub>O when NO<sub>2</sub><sup>-</sup> was supplied to the medium in autotrophic conditions (Figure 15 C). This production evidences that other proteins than

FLVs were involved in N<sub>2</sub>O production in this strain under the conditions tested. This challenges the conclusions of Burlacot et al. (2020) who proposed that FLVs were the main enzymes involved in N<sub>2</sub>O production in the light in *Chlamydomonas* (Burlacot et al., 2020).

Because CYP55 has been associated with N<sub>2</sub>O production in the dark, we tested the production of N<sub>2</sub>O from the *cyp55*-deficient mutant under illumination. As can be seen in Figure 15 D, this strain produced N<sub>2</sub>O in these conditions which confirms that CYP55 is not or poorly involved in N<sub>2</sub>O production in the light. This result requires to be confirmed by comparing the production with an adequate control strain. The production in darkness needs to be checked to confirm that no N<sub>2</sub>O is produced as previously suggested in the literature (Plouviez et al., 2017a; Burlacot et al., 2020). If these results are confirmed, the *cyp55*-deficient mutant could be a suitable strain to use during our microcosm assays in the dark.

#### 6.2.4. N<sub>2</sub>O synthesis in the *Chlamydomonas* wild type strain 21gr

Considering the lack of reproducibility observed with CMJ030, another suitable parental strain was investigated. The 21gr strain is the mt+ equivalent of 6145c which reproducibly produces N<sub>2</sub>O and whose genetic background is known. Besides, 21gr has been widely used to perform crossing and create double mutants. Thus, this strain has been selected to do a crossing between *flvb* and *cyp55* mutants to get the mutations in the 21gr background. As can be seen in Figure 16, this strain produced significant amounts of N<sub>2</sub>O.

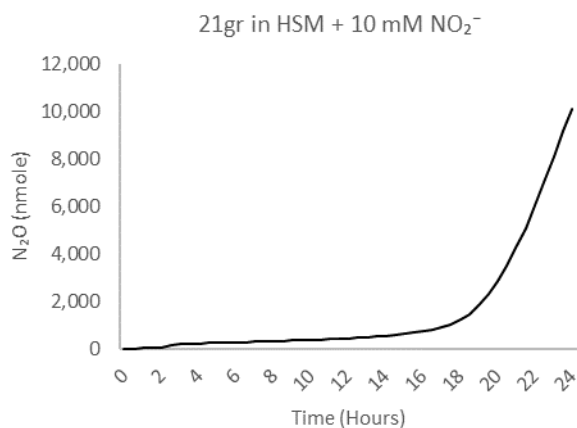


Figure 16: N<sub>2</sub>O batch test with *C. reinhardtii* 21gr (10 mg chlorophyll·L<sup>-1</sup> ~ 0.25 g DCW·L<sup>-1</sup>).

### 6.3. Inhibition assays

#### 6.3.1. Inhibition assay using the wild type *Chlamydomonas* 6145c

Exposure to excess light energy can lead to the degradation of the photosynthetic apparatus due to the generation of reactive oxygen species via the electron flow in the thylakoid membrane (Saroussi et al., 2019). Alternative pathways re-routing excess electrons are involved in the acclimation of *Chlamydomonas* to fluctuating environmental parameters such as light energy. We therefore conducted inhibition experiments using DCMU on wild type cells to investigate the contribution of the photosynthetic electron flow to N<sub>2</sub>O production.

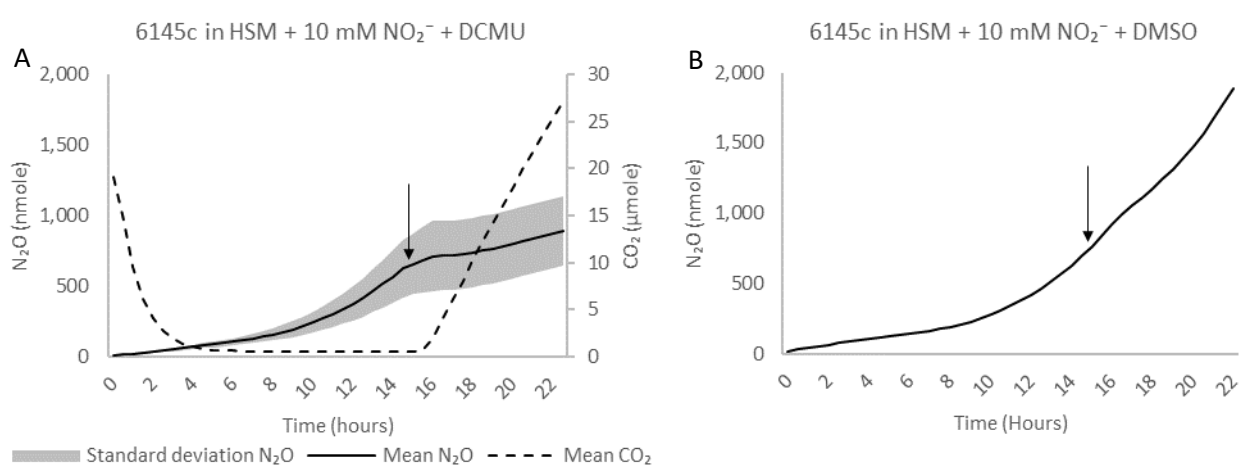


Figure 17: Inhibition assay with *C. reinhardtii* 6145c (10 mg chlorophyll·L<sup>-1</sup> ~ 0.25 g DCW·L<sup>-1</sup>). After 15 h of experiment, 10 μM DCMU (left) or DMSO (right) was added to the medium (arrow).

For this purpose, we also added DMSO (dimethyl sulfoxide) after 15 h in our control because the DCMU was diluted in DMSO. As shown in Figure 17 A, N<sub>2</sub>O production significantly decreased when DCMU is added to the medium whereas DMSO addition had not impact on the control. As DCMU blocks the electron flow coming from the photosystem II, the experimental evidence herein shown suggest that at least part of the N<sub>2</sub>O production depends on the photosynthetic electron flow.

#### 6.3.2. Inhibition test using lake samples

In order to study the biological source of N<sub>2</sub>O in lake Horowhenua, we performed inhibition assays on lake samples rich in microalgae (samples withdrawn during a bloom on 12/08/2021). Penicillin G potassium was selected because it has been previously shown to inhibit bacterial activity without inhibiting microalgal activity during N<sub>2</sub>O synthesis.

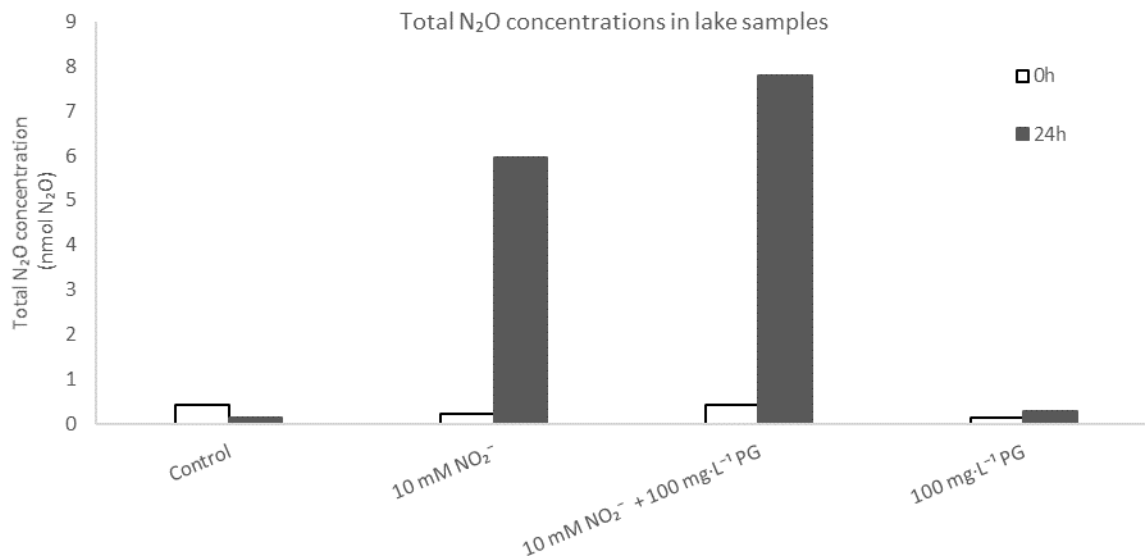


Figure 18: Inhibition assays using superficial samples from lake Horowhenua visibly rich in algae ( $28 \mu\text{g chlorophyll}\cdot\text{L}^{-1}$ ). At time 0h, either  $10 \text{ mM NO}_2^-$ ,  $10 \text{ mM NO}_2^- + 100 \text{ mg}\cdot\text{L}^{-1}$  Penicillin G potassium (PG), or  $100 \text{ mg}\cdot\text{L}^{-1}$  Penicillin G potassium was added to the sample.

As shown in Figure 18, N<sub>2</sub>O production increased when nitrite was added to the medium compared to the control after 24h of cultivation. This was expected as we observed the same trend during batch assays with pure algae cultures supplied with nitrite (see section 6.2). When the medium was supplied with Penicillin G potassium, a broad bacterial inhibitor, N<sub>2</sub>O production was similar to the production recorded in the control culture after 24h. We even obtained the highest N<sub>2</sub>O production when both nitrite and Penicillin G potassium were added to the culture medium, which suggests that bacteria are not the main producers of N<sub>2</sub>O.

However, these results need to be confirmed with further experiments performed on lake samples in triplicates.

## 6.4. qPCR

### 6.4.1. RNA extraction

Following RNA extraction from culture aliquots, 260/280 ratio values of around 2, which is expected for good quality samples, were measured for most samples (Appendix J). The shape of the absorbance curve of each sample was also checked as it informs on the quality of the sample.

### 6.4.2. qPCR results

As a first step, we decided to focus on the expression of *FLVA*, *FLVB*, *NIRK* and *CYP55* genes because the involvement of these genes in N<sub>2</sub>O production had already been demonstrated.

Following on primers optimization and efficiency tests, we performed qPCR on samples from M3 and CMJ030 cultures in HSM-N (control) and HSM-N + 10 mM NO<sub>2</sub><sup>-</sup> (treatment). The melting peak charts at different primer dilutions was compared to our samples to verify that the peak occurred at the expected temperature and, thus, that no unspecific amplification was performed during PCR (see Appendix K).

The qPCR results are not shown here because the expressions of our reference genes were highly variable both in the control and the treatment. Further optimization is therefore required. However, while we could not normalize the expression of *FLVA*, *FLVB*, *NIRK* and *CYP55*, their transcript numbers were similar in the controls and treatment. While these results are preliminary and specific to the conditions studied, the similarity between treatment and controls could mean that the proteins targeted were already present and active in the cells and/or that nitrite did not impact their expressions. If confirmed, these results could mean that RT-qPCR may not be suitable to identify the microalgae genes directly responsible for N<sub>2</sub>O synthesis following nitrite supply, but the technology may still be suitable to confirm the biological source of N<sub>2</sub>O in microcosms.



## 7. Conclusions

Preliminary data gathered from Lake Horowhenua show that the lake is a source of N<sub>2</sub>O (0.4 – 8.7 g N-N<sub>2</sub>O·ha<sup>-1</sup>·yr<sup>-1</sup>, n = 29 sampling events) and is eutrophic but no relationship between the algal biomass concentration and the N<sub>2</sub>O production could be evidenced at this stage. Based on our data, the national significance of N<sub>2</sub>O emissions from New Zealand eutrophic lakes is low compared to the fluxes reported in the literature. Further research is thus required to determine which parameters affect these emissions. Preliminary inhibition assays under illumination performed on lake samples supplied 10 mM NO<sub>2</sub><sup>-</sup> and 100 mg·L<sup>-1</sup> Penicillin G potassium suggest that bacteria are not the main producers of N<sub>2</sub>O in the lake.

While the ability of microalgae to produce N<sub>2</sub>O has been demonstrated in the laboratory and 2 pathways have been determined, new evidence suggest that other enzymes could be involved in N<sub>2</sub>O production: CYP55, which catalyses the reduction of NO to N<sub>2</sub>O, has thus been proposed as the main enzyme involved in *Chlamydomonas* in the dark (Plouviez et al., 2017a), while flavodiiron proteins (FLVs) have been suggested to be responsible for N<sub>2</sub>O production in the light (Burlacot et al., 2020). In additions, hybrid cluster proteins (HCPs) present in both bacteria and *Chlamydomonas* are producing N<sub>2</sub>O in *E. coli* in anaerobic conditions and could be involved in N<sub>2</sub>O production in *Chlamydomonas*. Consequently, investigating the different putative N<sub>2</sub>O production pathways in the laboratory is still needed to confirm the enzymes involved during microalgal N<sub>2</sub>O synthesis and thus select the right genes in our field samples.

Preliminary batch assays under illumination demonstrated that wild type *Chlamydomonas reinhardtii* produced N<sub>2</sub>O up to 31.5 ± 10.7 μmole N<sub>2</sub>O·g DW<sup>-1</sup> over 24 hours in autotrophic conditions and when supplied 10 mM NO<sub>2</sub><sup>-</sup>. Inhibiting the electron flow coming from PSII in the same conditions suggests that the electron transport chain is involved in N<sub>2</sub>O production. Mutant deficient in FLVs were however still synthesizing N<sub>2</sub>O when supplied with NO<sub>2</sub><sup>-</sup> under illumination, showing that the FLVs were not the only enzymes involved in N<sub>2</sub>O production in autotrophic conditions, which challenges conclusions from the literature. These results were presented at the Plant Science Central conference (6-8 July 2021; Massey University, Palmerston North, New Zealand; New Zealand Institute of Agricultural and Horticultural Science (NZIAHS), New Zealand Branch of the International Association of Plant Biotechnology (NZIAPB), New Zealand Society of Plant Biologists (NZSPB)) and at the Physiomar & Australia New Zealand Biotechnology Society (ANZMBS) 2021 Conference held online virtually (7-9 September 2021).

My research has therefore demonstrated:

1. The suitability of pure algae bioassays aiming at determining the pathways (e.g. use of mutants and inhibitor). This preliminary research has however evidenced a knowledge gap regarding the pathways, and we have not yet been able to identify clear conditions repressing algae N<sub>2</sub>O synthesis. Future research should finalize a crossing over of mutants that should be repressed in their ability to produce N<sub>2</sub>O under the 'standard conditions' of production. Our results also showed that the genetic background plays an important role in N<sub>2</sub>O production. These are important to consider for designing future assays.
2. The suitability of protocols for studying gene expression. The RNA extraction protocol is satisfactory but the qPCR still requires optimization.
3. The suitability of protocols for field analysis.

Thus, all the methods required to achieve Objective 1 are suited and the methods for Objective 2 are being optimized. Based on the literature (Hanson & Stefan, 1984; Jansson et al., 2000; Mahan et al., 2005; Guieysse et al., 2013; Zhou et al., 2020), it should be possible to use inhibitors such as antibiotics to distinguish algal from bacterial N<sub>2</sub>O emissions and thereby study the production from pure algal and/or bacterial cultures and lake samples. In future work, the potential ecological interactions between algal and bacterial N<sub>2</sub>O synthesis could also be assessed, although this still requires the identification of suitable mutants as noted above.

The next steps to achieve this project would be to:

- Continue field monitoring at lake Horowhenua and the Turitea reservoir
- Test the influence of P supply on N<sub>2</sub>O emissions from axenic *Chlamydomonas* cultures
- Identify and acquire pure nitrifiers and facultative denitrifiers and test N<sub>2</sub>O emissions from these species.
- Identify and test inhibitors on pure bacteria and artificial algae-bacteria consortium
- Test the impact of different N and P supply on N<sub>2</sub>O synthesis in microcosms and compare the N<sub>2</sub>O production from pure algae culture, pure bacteria culture, artificial algae and bacteria cultures and lake samples.

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## 9. Appendices

### Appendix A. Sources and sinks of N<sub>2</sub>O

Table 17: Sources and sinks of N<sub>2</sub>O (Tian et al., 2020)

<b>SOURCES</b>	
<b>Anthropogenic</b>	
Direct emissions from agricultural N	Direct soil emissions
	Manure left on pasture
	Manure management
Other direct anthropogenic sources	Aquaculture
	Biomass burning
	Fossil fuels and industry
Indirect emissions	Waste and wastewater
	Atmospheric N deposition on land
	Atmospheric N deposition on ocean
Perturbed fluxes caused by changes in climate and land cover	Estuaries, inland and coastal waters
	Climate effect
	Post deforestation pulse effect
<b>Natural</b>	
Natural soils	
Lightning and atmospheric production	
Estuaries, inland and coastal waters	
Oceans	
<b>SINKS</b>	
<b>Anthropogenic</b>	
Perturbed fluxes caused by changes in CO <sub>2</sub> and land cover	Long-term effect of reduced mature forest area
	CO <sub>2</sub> effects
<b>Natural</b>	
Atmospheric chemical sink	Photolysis
	Reaction with electronically excited atomic oxygen
Surface sink	

## Appendix B. Emissions from managed lands

Regarding the indirect emissions due to managed lands, two indirect pathways are considered for the anthropogenic input of N to water bodies: (i) the volatilization of  $\text{NH}_3$  and oxides of N and their deposition on water surface and (ii) the leaching and runoff of N from managed land. The IPCC's methodology is based on the three-tiered approach previously described and uses two EFs, one for each pathway, to estimate indirect  $\text{N}_2\text{O}$  emissions from waterbodies that receive (N)-laden flux. The  $\text{N}_2\text{O}$  emissions are calculated as a fraction of the N flux to the aquatic body that is further converted into  $\text{N}_2\text{O}$  using the set of equations described in the Box 3. The indirect emissions due to N leaching and runoffs from agricultural soils takes into account the  $\text{N}_2\text{O}$  originating from 3 types of water bodies: (i) groundwater<sup>17</sup> and surface drainage (e.g. tile drainage, drainage ditches); (ii) rivers and others<sup>18</sup>; and (iii) estuaries<sup>19</sup> (*Revised 1996 IPCC Guidelines for National Greenhouse Gas Inventories : Reference Manual (Volume 3)*, 1997).

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<sup>17</sup> including springs and upstream emissions corresponding to the degassing of groundwater

<sup>18</sup> including lakes, ponds and downstream emissions from the nitrification/denitrification of degassed water

<sup>19</sup> excluding outer estuaries and coastal waters

**Box 3:** Indirect N<sub>2</sub>O emissions from managed soils (De Klein et al., 2006; Hergoualc'h et al., 2019).

Indirect N<sub>2</sub>O emissions from managed soils are calculated as a fraction of the nitrogen flux reaching the aquatic body. The N sources included are synthetic and organic fertilizers, urine and dung from grazing animals in pastures, crop residues and N mineralization following land use change.

- Tier 1 method:

The indirect N<sub>2</sub>O emissions from atmospheric deposition of N volatilised from managed soils are calculated as follow:

$$N_2O_{(ATD)} - N = [(F_{SN} \cdot Frac_{GASF}) + (F_{ON} + F_{PRP}) \cdot Frac_{GASM}] \cdot EF_4$$

Where N<sub>2</sub>O<sub>(ATD)</sub> - N is the annual amount of N<sub>2</sub>O - N produced from atmospheric deposition of N volatilised from managed soils (kg N<sub>2</sub>O - N·yr<sup>-1</sup>), F<sub>SN</sub> is the annual amount of synthetic fertilizer N applied to soils (kg N·yr<sup>-1</sup>), F<sub>ON</sub> is the annual amount of organic N additions applied to soils (kg N·yr<sup>-1</sup>), F<sub>PRP</sub> is the annual amount of urine and dung N deposited by grazing animals (kg N·yr<sup>-1</sup>), Frac<sub>GASF</sub> (default = 0.11 kg N volatilised·kg of N applied<sup>-1</sup>) is the fraction of synthetic fertiliser N that volatilises as NH<sub>3</sub> and NO<sub>x</sub> (kg N volatilised·kg of N applied<sup>-1</sup>), Frac<sub>GASM</sub> (default = 0.21 kg N volatilised·kg of N applied or deposited<sup>-1</sup>) is the fraction of applied organic N fertiliser materials and of urine and dung N deposited by grazing animals that volatilises as NH<sub>3</sub> and NO<sub>x</sub> (kg N volatilised·kg of N applied or deposited<sup>-1</sup>) and EF<sub>4</sub> (default = 0.010 kg N-N<sub>2</sub>O·kg NH<sub>3</sub>-N + NO<sub>x</sub>-N volatilised<sup>-1</sup>) is the emission factor for N<sub>2</sub>O emissions from atmospheric deposition of N on soils and water surfaces (kg N-N<sub>2</sub>O·kg NH<sub>3</sub>-N + NO<sub>x</sub>-N volatilised<sup>-1</sup>).

The indirect N<sub>2</sub>O emissions from leaching and runoff can be calculated using the equation below:

$$N_2O_{(L)} - N = (F_{SN} + F_{ON} + F_{PRP} + F_{CR} + F_{SOM}) \cdot Frac_{LEACH-(H)} \cdot EF_5$$

Where N<sub>2</sub>O<sub>(L)</sub> - N is the annual amount of N<sub>2</sub>O - N produced from leaching and runoff of N additions to managed soils (kg N<sub>2</sub>O - N·yr<sup>-1</sup>), F<sub>CR</sub> is the amount of N in crop residues (kg N·yr<sup>-1</sup>), F<sub>SOM</sub> is the annual amount of N mineralised as a result of changes to land use or management (kg N·yr<sup>-1</sup>), Frac<sub>LEACH-(H)</sub><sup>20</sup> (default = 0.24 kg N·kg of N additions<sup>-1</sup>) is the fraction of all N added to/mineralised in managed soils (kg N·kg of N additions or deposition by grazing animals<sup>-1</sup>) and EF<sub>5</sub> (default = 0.011 kg N<sub>2</sub>O - N·kg N leached and runoff<sup>-1</sup>) is the emission factor for N<sub>2</sub>O emissions from N leaching and runoff (kg N - N<sub>2</sub>O·kg N leached and runoff<sup>-1</sup>).

The conversion of N<sub>2</sub>O<sub>(ATD)</sub> - N and N<sub>2</sub>O<sub>(L)</sub> - N emissions to N<sub>2</sub>O is done using the equations:

$$N_2O_{(ATD)} = N_2O_{(ATD)} - N \cdot \frac{44}{28} \text{ and } N_2O_{(L)} = N_2O_{(L)} - N \cdot \frac{44}{28}$$

Where  $\frac{44}{28}$  is the conversion factor of kg N<sub>2</sub>O - N into kg N<sub>2</sub>O.

- Tier 2 method

A similar equation as for Tier 1 can be adopted using country-specific values for emission factors and partitioning fractions for each N source. For instance, N<sub>2</sub>O<sub>(ATD)</sub> - N would be calculated as follow:

$$N_2O_{(ATD)} - N = \left\{ \sum_i (F_{SN_i} \cdot Frac_{GASF_i}) + [(F_{ON} + F_{PRP}) \cdot Frac_{GASM}] \right\} \cdot EF_4$$

Where F<sub>SN<sub>i</sub></sub> is the annual amount of synthetic fertilizer N applied to soils under different conditions *i* (kg N·yr<sup>-1</sup>) and Frac<sub>GASF<sub>i</sub></sub> is the fraction of synthetic fertiliser N that volatilises as NH<sub>3</sub> and NO<sub>x</sub> under different conditions *i* (kg N volatilised·kg of N applied<sup>-1</sup>). The conversion of N<sub>2</sub>O<sub>(ATD)</sub> - N emissions to N<sub>2</sub>O emissions is similar to the one used in Tier 1.

<sup>20</sup> In dry climate regions where no irrigation is used or where rainfall - evaporation does not exceed the soil water-holding capacity, the default value for Frac<sub>LEACH-(H)</sub> is 0.

## Appendix C. Natural N<sub>2</sub>O emissions from aquatic environments

### C.1. Inland and coastal waters (lakes, reservoirs, rivers, estuaries)

Tian et al. (2020) estimated that natural emissions from inland and coastal waters were 500 (200 – 700)<sup>21</sup> kt N-N<sub>2</sub>O·year<sup>-1</sup> from 2007 to 2016. Most papers compared their estimations to the 2006 IPCC value of 600 kt N-N<sub>2</sub>O·year<sup>-1</sup>. To simplify, we decided to adjust the comparisons to the most recent value from Tian et al. (2020).

Based on the data from 223 studies, DelSontro et al. (2018) used a model combining lakes size and primary productivity to specifically estimate the annual N<sub>2</sub>O emissions from lakes and impoundments. These authors estimated these emissions at 190-450 kt N-N<sub>2</sub>O·year<sup>-1</sup>. Following a different approach, Maavara et al. (2019) used a mass balance mechanistic model to estimate emissions from rivers, estuaries and coastal zones and assessed that these emissions were about half of the estimations from Tian et al. (2020), at around 148 – 277 kt N-N<sub>2</sub>O·year<sup>-1</sup>, from which 42.5 – 71.5 kt N-N<sub>2</sub>O·year<sup>-1</sup> would be emitted by reservoirs (which does not include lakes, wetlands and coastal aquaculture ponds). That last value is close to the 30 kt N-N<sub>2</sub>O·year<sup>-1</sup> estimated by Deemer et al. (2016) for all reservoirs. Using the same model as Maavara et al. (2019), Lauerwald et al. (2019) estimated that 63 kt N-N<sub>2</sub>O·year<sup>-1</sup> were emitted by all natural lakes and reservoirs, which is 5 times less than the estimation from DelSontro et al. (2018). These authors also noted that reservoirs were responsible for half of these emissions even though their area represents only 9% of the total area for all natural lakes and reservoirs. The suggested reason behind this observation is that reservoirs are more prone to N and P influxes due to their proximity to urban and agricultural zones. As can be seen, the values found in the literature differ greatly depending on the approach used and the aquatic ecosystems considered.

A better understanding of the drivers of N<sub>2</sub>O production is needed to improve estimates of N<sub>2</sub>O emissions from lentic<sup>8</sup> waters. Miao et al. (2020) noticed a significant seasonality and spatial variation of N<sub>2</sub>O emissions from lake Chaohu in China, with the highest N<sub>2</sub>O flux originating from the most nutrient-polluted zones during summer. The emissions of N<sub>2</sub>O are thus under the control of different factors. These authors suggested that this phenomenon could be partially explained by the variation of the average transfer velocity of N<sub>2</sub>O, the water temperature and the dissolved oxygen (DO) concentration as the DO influences nitrification and denitrification processes.

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<sup>21</sup> The values written in brackets correspond to the lower and higher end of range.



## C.2. Oceans

The IPCC estimates that oceans naturally emit 3,800 (1,800 - 9,400)<sup>21</sup> kt N-N<sub>2</sub>O year<sup>-1</sup>, which represents around 35% of all natural N<sub>2</sub>O emissions in 2006 (Ciais et al., 2013). Similar estimates of 4,000 (1,200 - 6,800)<sup>21</sup> kt N-N<sub>2</sub>O.year<sup>-1</sup> and 4,200 ± 1,000 kt N-N<sub>2</sub>O.year<sup>-1</sup> were proposed by Nevison et al. (1995) and Yang et al. (2020), respectively.

N<sub>2</sub>O distribution in the ocean is related to oxygen distribution (Kock et al., 2016). The highest N<sub>2</sub>O emissions are generally recorded at the boundaries of oxygen minimum zones (OMZ) of the ocean (Arévalo-Martínez et al. 2015a). These zones are mostly located on the west coast of continents, mainly in tropical zones, and are found at different depths. Their distribution is determined by the water surface temperature, the large-scale ocean currents circulation, and the biological and chemical processes occurring in the ocean. Oxygen solubility decreases with water temperature and the ocean areas located between gyres<sup>22</sup> are less mixed and, therefore, less oxygenated. The high biological activity occurring in the upwelling areas is also responsible for oxygen depletion in OMZs. This low oxygen level could explain the partial denitrification during which the conversion of N<sub>2</sub>O to N<sub>2</sub> does not take place (Arévalo-Martínez et al., 2015a; Arévalo-Martínez et al., 2015b). However, N<sub>2</sub>O is consumed in anoxic waters, such as in the core of the OMZ (Arévalo-Martínez et al., 2015a). Besides, Yang et al. (2020) observed that N<sub>2</sub>O emissions from the ocean show a strong seasonality, as reported for other ecosystems such as lakes. These authors used over 158,000 measurements from the ocean to reconstruct climatological N<sub>2</sub>O emissions and determined that the seasonality is mostly due to the variations in wind speed and the disequilibrium between air and water N<sub>2</sub>O concentrations. They also demonstrated that the seasonality of the N<sub>2</sub>O concentration disequilibrium was associated to the chlorophyll concentration, the subsurface apparent oxygen utilization (AOU) and, to a lesser extent, sea-surface temperature (SST) and mixed-layer depth. The nature of the correlation was dependent on the regions considered and is presented in Table 18. N<sub>2</sub>O emissions seasonality could be explained at over 66% by changes in the N<sub>2</sub>O disequilibrium and by changes in wind speed at 33%. The Arctic Ocean and, occasionally, the subtropical gyres were predicted to consume N<sub>2</sub>O while the upwelling regions (e.g. Peruvian upwelling), the subpolar Pacific Ocean and the Southern Ocean were predicted as main producers. Thus, upwelling regions representing only 3% of oceanic areas could be responsible for 20% of all N<sub>2</sub>O oceanic emissions (840 kt N-N<sub>2</sub>O.year<sup>-1</sup>) (Yang et al., 2020).

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<sup>22</sup> Large system of circulating ocean currents.

Table 18: Regimes and empirical drivers of N<sub>2</sub>O disequilibrium seasonality (Yang et al., 2020)

	Regimes of N <sub>2</sub> O disequilibrium seasonality			
	Upwelling	Mixed-layer	Solubility	Productivity-driven
<b>Location</b>	Tropical and northern Indian Ocean, eastern tropical Pacific and Atlantic Oceans, parts of the Southern Ocean	Southern Ocean Polar Front, subpolar Atlantic and Pacific	Ocean's subtropical gyres	Other parts of the ocean
<b>Empirical drivers positively correlated to N<sub>2</sub>O supersaturation</b>	Chlorophyll, subsurface AOU	AOU, mixed-layer depth	SST	Chlorophyll
<b>Empirical drivers negatively correlated to N<sub>2</sub>O supersaturation</b>	SST	Chlorophyll, SST	Chlorophyll, AOU, mixed-layer depth	

Apart from OMZ, hypoxia<sup>23</sup> can occur in the ocean due to eutrophication. These hypoxic zones are often found near coasts and are contributing to N<sub>2</sub>O emissions (Naqvi et al., 2009). For example, N<sub>2</sub>O emissions of 25 - 287 nmole·m<sup>-2</sup>·h<sup>-1</sup> were reported in the Gulf of Mexico hypoxic zone receiving nutrients-rich fluxes from the Mississippi and Atchafalaya Rivers (Walker et al., 2010).

<sup>23</sup> Oxygen-depleted condition triggered by the aerobic biodegradation of exogenous organic matter and/or endogenous respiration.

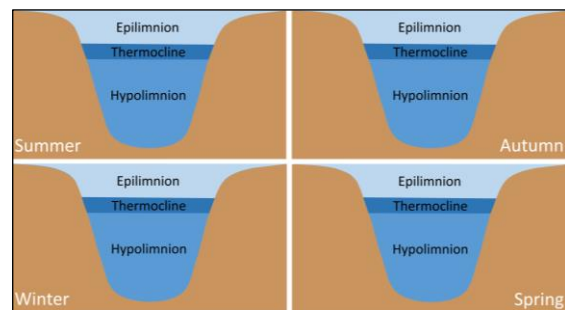
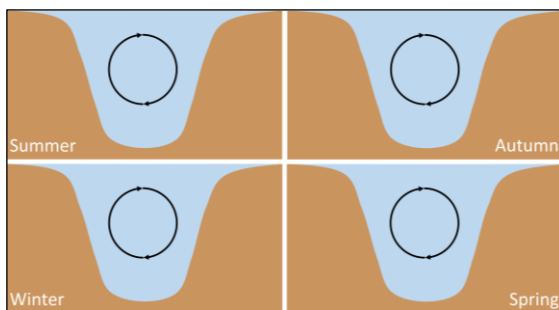
## Appendix D. General lake characteristics

Lakes can be classified according to their location (polar, temperate, or tropical), origin (glacial, volcanic, riverine, etc.), circulation and mixing patterns, or trophic state. The main definitions related to thermal stratification and the different mixing patterns are illustrated in Box 4.

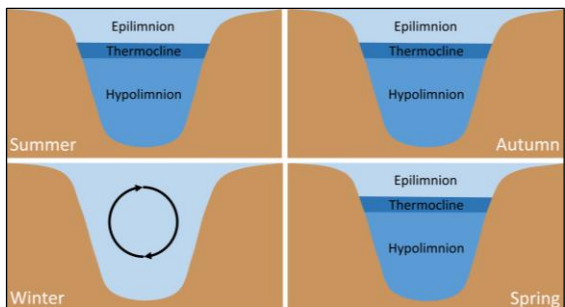
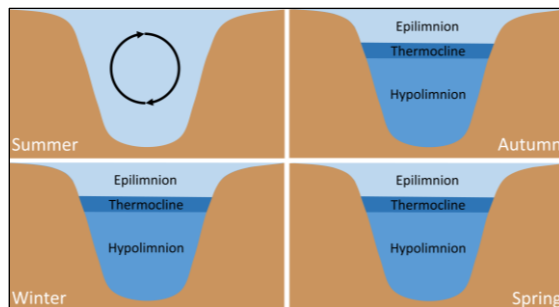
### Box 4: Thermal stratification and mixing patterns of lakes.

When a lake shows a thermal stratification, the surface layer, called epilimnion and the lowest layer, called hypolimnion are separated by a transition layer called the thermocline. The layers can mix during the year and the different mixing patterns found in lakes are described below.

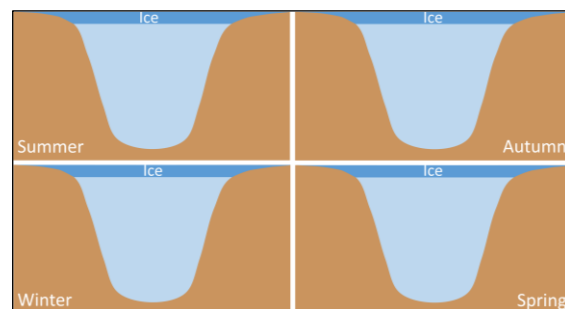
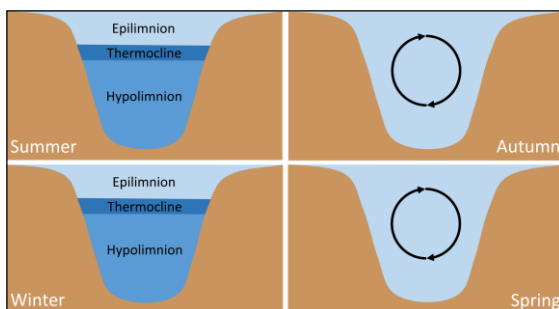
- *Polymictic lake: Non-stratified lake, the water is mixed all the year.*
- *Oligomictic lake: Lake that rarely mixes.*



- *Monomictic lake: Lake that mixes only once a year. The lake is either warm monomictic if the mixing occurs during winter or cold monomictic if it occurs during summer months.*



- *Dimictic lake: Lake that mixes twice a year.*
- *Amictic lake: Lake that never mixes and are isolated by an ice layer on top.*



## Appendix E. TLI score calculation

### Box 5: TLI score.

Annual average values of chlorophyll a (Chla) concentration, Secchi depth (SD), total phosphorous (TP) concentration and total nitrogen (TN) concentration of 24 New Zealand lakes measured over a period of two to four years were used by Burns et al. (2000) to elaborate the TLI. The average annual values are used to normalise the data by suppressing the variability due to the different seasons over the year.

First, Burns et al. (2000) chose to base the TLI score on the Chla concentration: the TLI values of 3, 4 and 6 were assigned to Chla concentrations of 2, 5 and 30 mg·m<sup>-3</sup> based on the description of New Zealand lake trophic classes made by Vant (Davies-Colley et al., 1993). A plot of the 3 TLI values against the base-10 logarithm of the 3 Chla concentrations followed by an ordinary least squares (OLS) regression gives the equation for the trophic level index of Chla called TLc:

$$TLc = 2.22 + 2.54 \log(Chla) \quad (1)$$

Using equation (1), the average annual TLc are calculated from the average annual Chla concentration values for each lake. Then, the plot of TLc against the base-10 logarithm of TN and TP values followed by an OLS regression gives the equations for the trophic level indexes of TN and TP, respectively TLn and TLp, assuming that TLc = TLn = TLp:

$$TLn = -3.61 + 3.01 \log(TN) \quad (2)$$

$$TLp = 0.218 + 2.92 \log(TP) \quad (3)$$

The SD being inversely proportional to the trophic state, it is necessary to first transform this variable before following the same calculation method as for TLn and TLp to determine TLs. The transformation applied by Chapra & Dobson (1981) using the Beer-Lambert law for light extinction was employed, and the extinction coefficient due to water itself was adapted to the highest SD reported in New Zealand. The new SD expression considered was  $\frac{1}{SD} - \frac{1}{40}$ . Then, assuming that TLc = TLs, the expression obtained for TLs was:

$$TLs = 5.10 + 2.27 \log\left(\frac{1}{SD} - \frac{1}{40}\right) \quad (4)$$

Finally, the average TLI can be calculated as the average value of the four trophic level indexes:

$$TLI = \frac{1}{4} (TLc + TLs + TLp + TLn) \quad (5)$$

## Appendix F. Potential sampling sites list

Table 19: Potential sampling sites list

	Lake	Mixing pattern	TLI in 2019	Trophic state	Geomorphic Type	Distance from Massey (km)	Area (ha)	Max Depth (m)	Depth (Lakes380) (m)	Altitude (m)	Authorization required
Eutrophic lake	Lake Koputara	Polymictic	8	Eutrophic	Aeolian	40.7					Maori-owned lake
	Omanuka lagoon	Polymictic	5.7	Eutrophic	Aeolian	43.3				29	Maori-owned lake
	Pukepuke lagoon	Polymictic	7.2	Eutrophic	Aeolian	44.5	15	2	1	10	Anyone visiting the wetland requires a (free) permit from the local DOC office.
	Lake Alice	Polymictic	5.8	Eutrophic	Aeolian	45.7	12	3.5		116	
	Lake Horowhenua	Polymictic	6.7	Eutrophic	Aeolian	47.8	389	2			Maori-owned lake
	Lake William	Polymictic	6.6	Eutrophic	Aeolian	48.2	7		9.8	105.8	
	Lake Heaton	Polymictic	7.1	Eutrophic	Aeolian	49.3	14		5.00	97	
	Lake Dudding	Stratified	5.6	Eutrophic	Aeolian	50.7	8	12	7.3	92	No
	Lake Herbert	Polymictic	6.2	Eutrophic	Aeolian	50.9					
	Lake Koiitiata	Polymictic	4.8	Eutrophic	Aeolian	58.8	10	5	1.60	18	
	Lake Waipu	Polymictic	6.3	Eutrophic	Aeolian	64.8	7		4.3	20	
	Lake Pauri	Stratified	5.6	Eutrophic	Aeolian	73.2	19	8	14.2	58	No
	Lake Wiritoa	Stratified	4.9	Eutrophic	Aeolian	77.7	22	20	19	51	No
	Lake Kohata	Stratified	4.9	Eutrophic	Aeolian	78.2	10	5			
	Lake Westmere	Polymictic	6.7	Eutrophic	Aeolian	87.2	8	5	6	96	
Lake Kawau					33.5	1		1.10	71		
Karere lagoon					15.7	3		1.60	12.6		
Voss lagoon					18.6	5		1.6	10.6		
Clean lake	Turitea Dams			Mesotrophic	River dam	6.8	12	10		170	Yes
	Upper Karori Lake			Mesotrophic	River dam	141	0.1	9		169	
	Ohakune lakes			Oligotrophic	Volcanic	174	2	5		579	
	Lake Rotorangi	Stratified	4	Mesotrophic	River dam	197	290	60		111	
	Lake Opunake			Mesotrophic	River dam	212	10	5		19	No
	Lake Otamangakau			Oligotrophic	River dam	220	150	10		617	No
	Lake Ratapiko			Mesotrophic	River dam	222	24	10		205	No
	Lake Okareka	Stratified	3.1	Mesotrophic	Volcanic	346	334	34		355	No
	Lake Tikitapu	Stratified	2.7	Oligotrophic	Volcanic	346	144	28		418	No
	Lake Waikaremoana			Oligotrophic	Landslide	347	5574	248		585	No
	Lake Tarawera		3.1	Oligotrophic	Volcanic	353	4115	88		299	No
	Lake Waikareiti			Oligotrophic	Landslide	355	1371	50		900	No
	Lake Okataina	Stratified	2.7	Oligotrophic	Volcanic	357	1073	79	78	311	No
	Lake Rotoiti	Stratified	4	Mesotrophic	Volcanic	358	3369	126	126	297	No
Lake Rotoma	Stratified	2.5	Oligotrophic	Volcanic	373					No	

## Appendix G. Comparison between different methods of quantification

In microbiology, the chlorophyll concentration is often used as an indicator of the cell concentration in a sample. In engineering, the cell concentration is preferably expressed in unit of mass of cells per unit of volume. To easily switch from one unit to the other, the main objective of this experiment was to check the relationship between different quantification units: chlorophyll a and b concentration, cell concentration, cell optical density and dry cell weight.

For this experiment, axenic *Chlamydomonas* cultures were grown in similar conditions than the batch assays set-up and biological triplicates were used. We tested the strain 6145c in TAP and minimal media, M3 in TAP and HSM and CMJ030 in HSM.

As expected, the values of the triplicate cultures were similar during the week. However, the cell counting method is the least accurate method because the cells are not always distributed in a homogeneous manner on the scales. The values of the DCW at time 0 are not very reliable because the weight differences between the filters alone and the filters with cells were not significant. The correlations between the different quantification methods are good with a  $R^2 > 0,97$  for all of them for the culture of 6145c in TAP medium (Figure 19).

The results for the other strains and conditions are shown in Figure 20 - Figure 23. The experiments must be repeated to improve the results by taking more mid-point values. Besides, the cell growth is much slower in MM and HSM compared to TAP medium which increases the errors as the daily values are all similar.

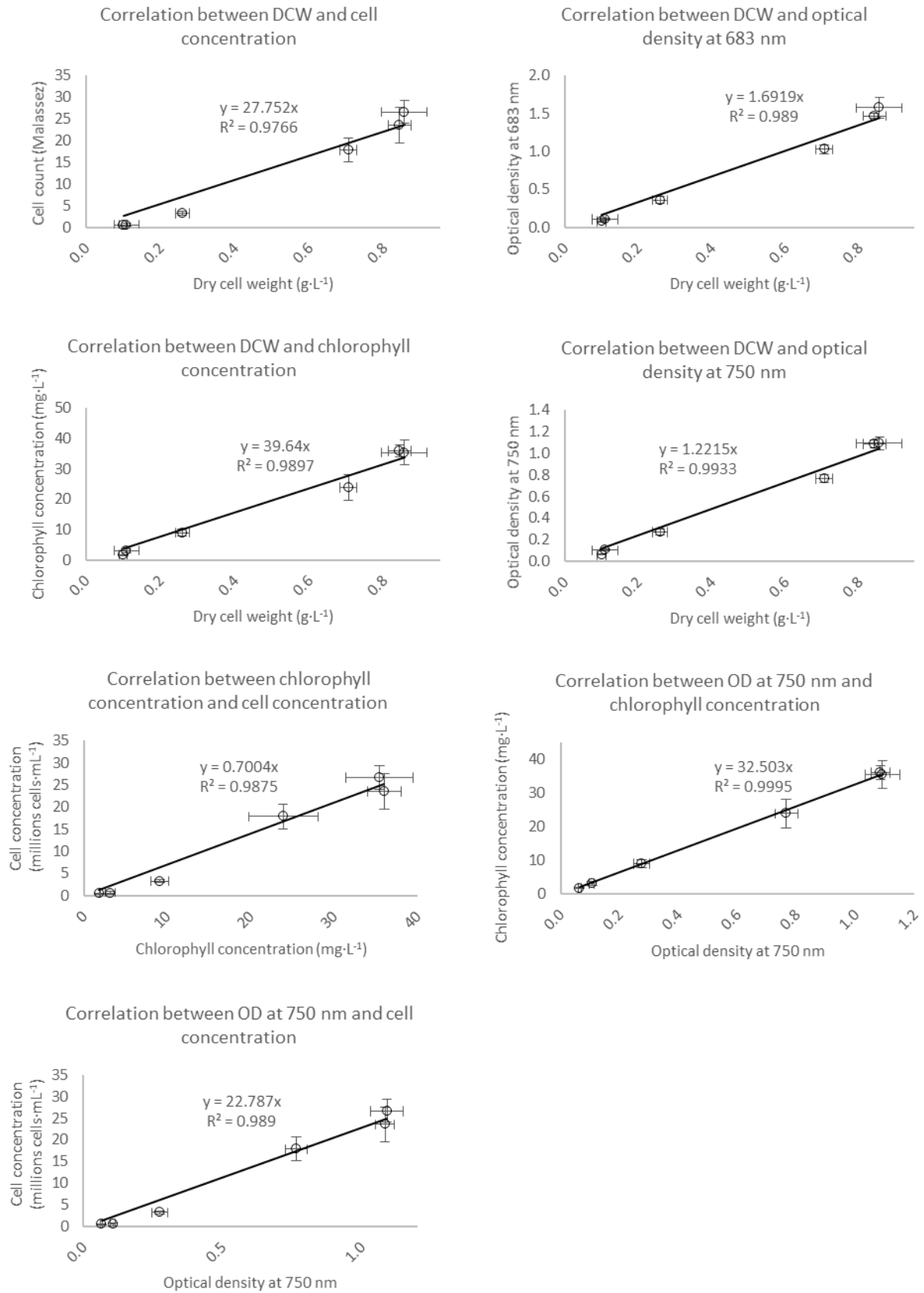


Figure 19: Correlation between the chlorophyll concentration, dry cell weight, optical densities at 750 nm and 683 nm and cell concentration of *C. reinhardtii* 6145c cultures in TAP medium.

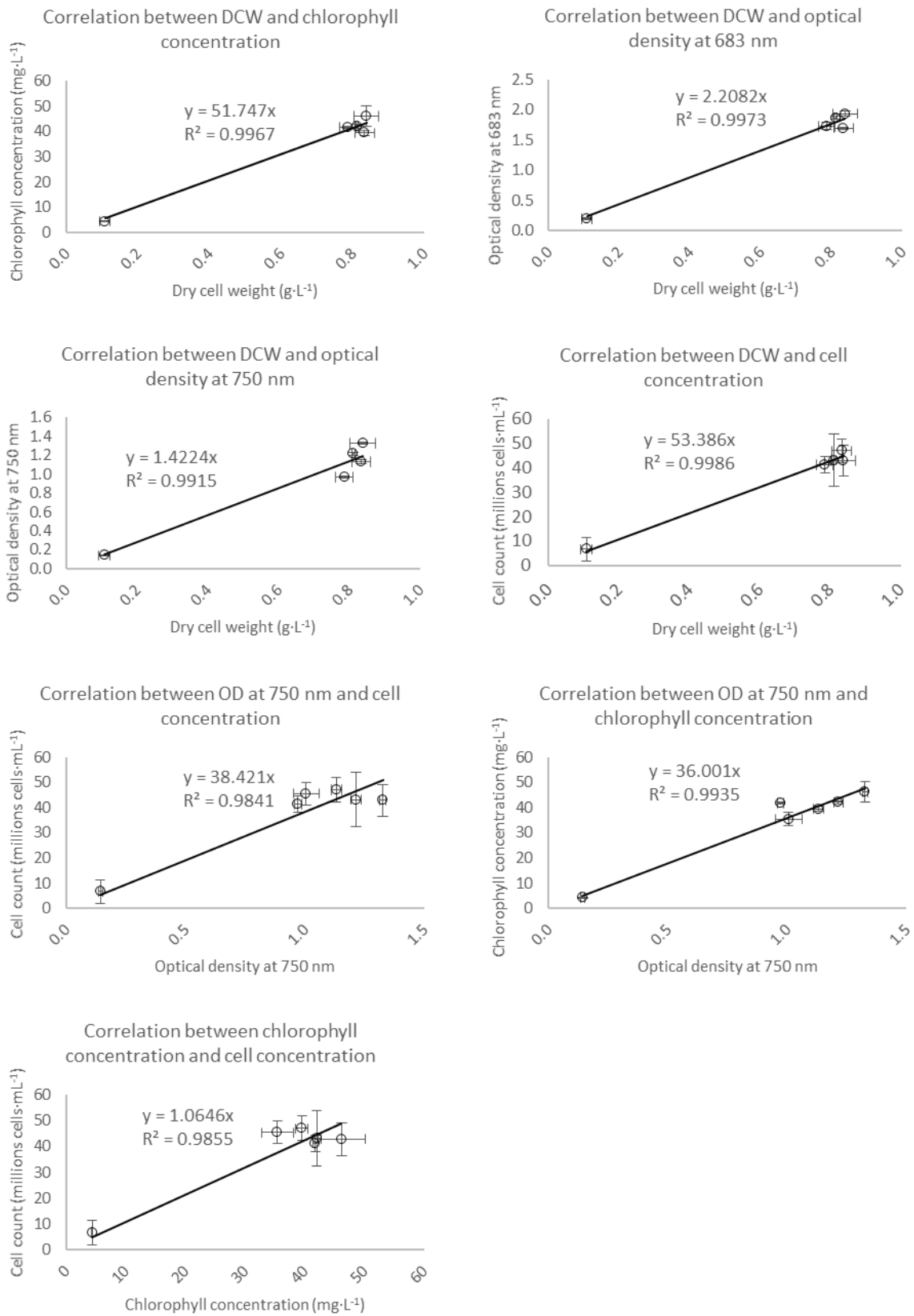


Figure 20: Correlation between the chlorophyll concentration, dry cell weight, optical densities at 750 nm and 683 nm and cell concentration of *C. reinhardtii* M3 cultures in TAP medium.



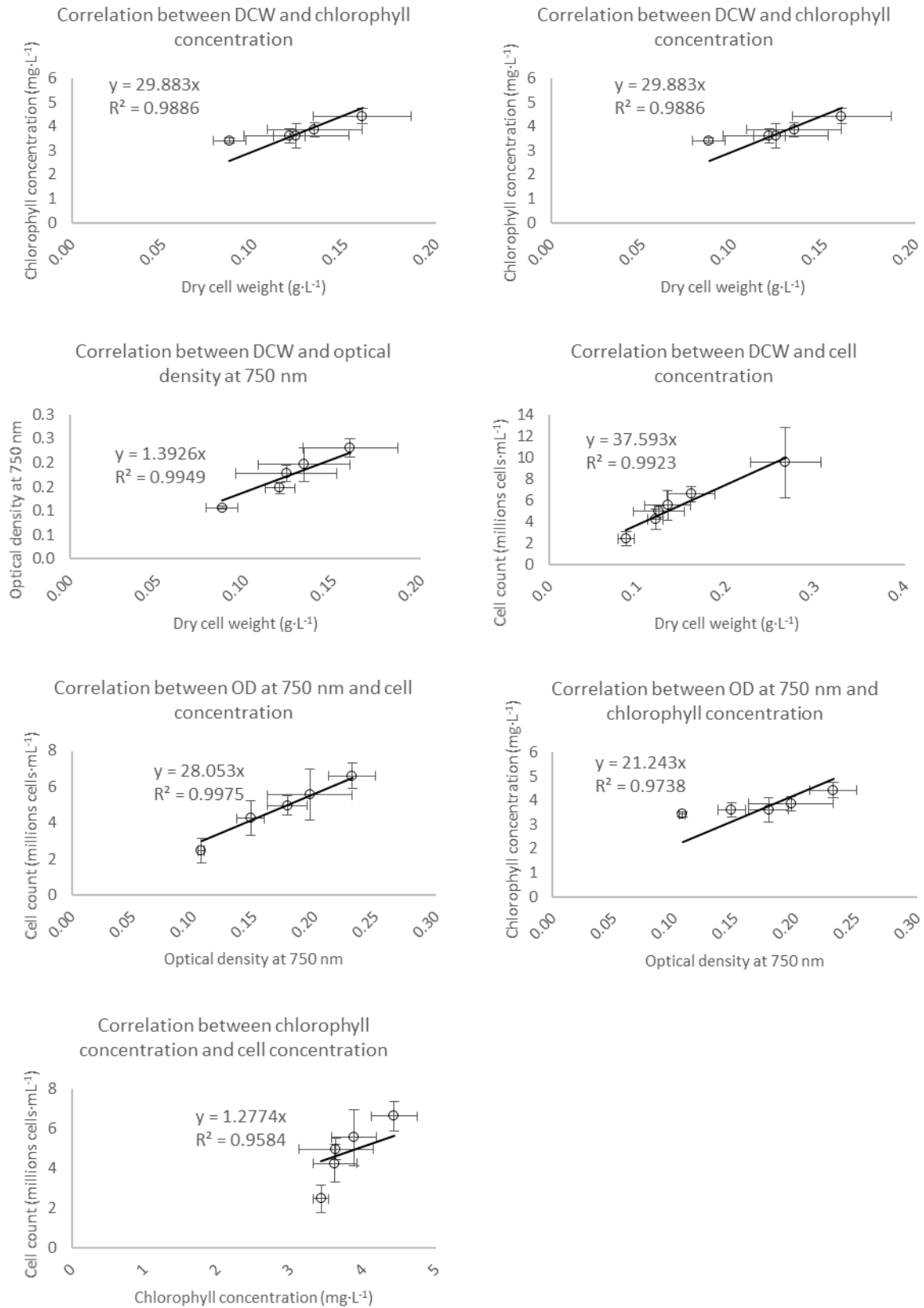


Figure 21: Correlation between the chlorophyll concentration, dry cell weight, optical densities at 750 nm and 683 nm and cell concentration of *C. reinhardtii* 6145c cultures in MM.

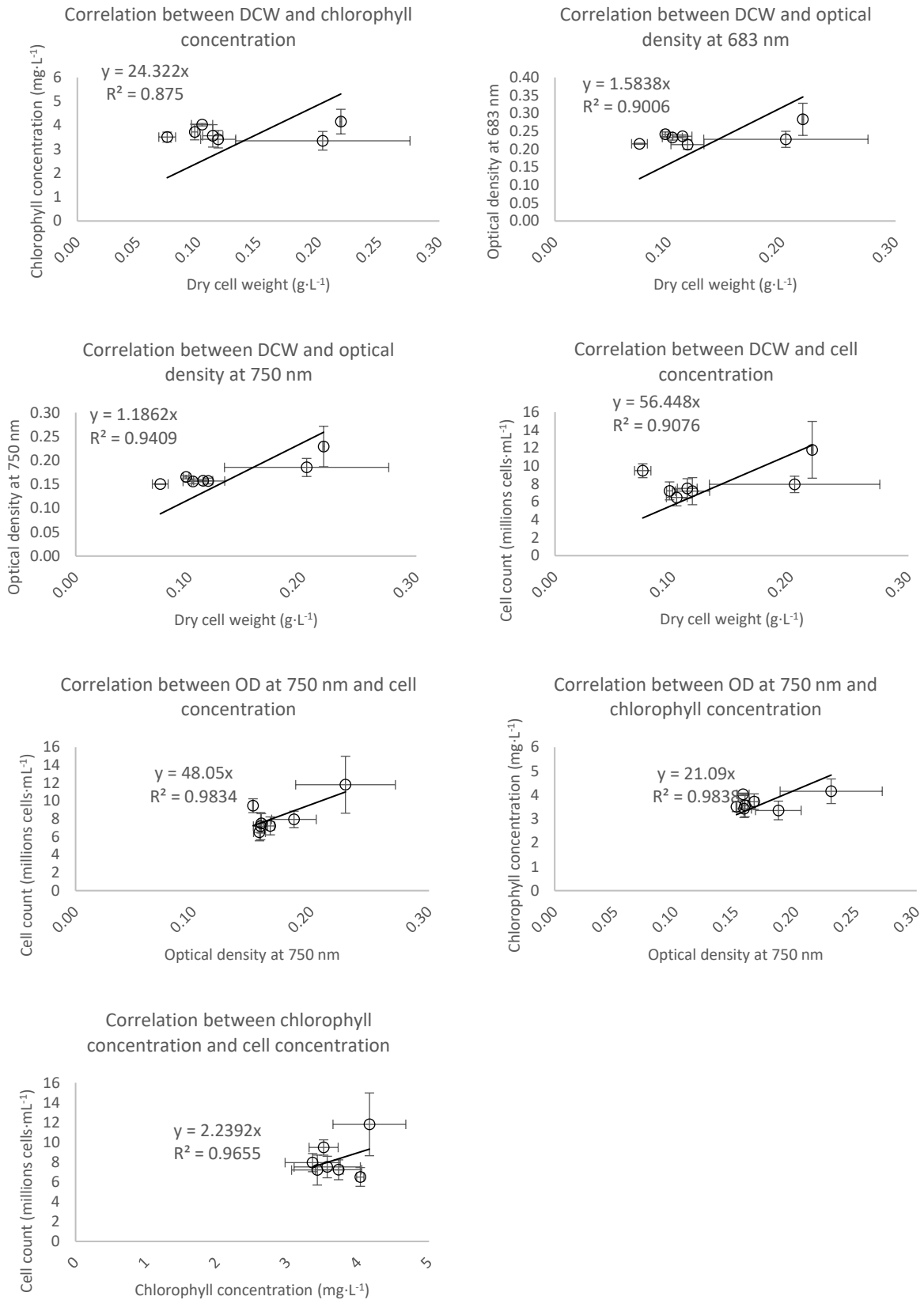


Figure 22: Correlation between the chlorophyll concentration, dry cell weight, optical densities at 750 nm and 683 nm and cell concentration of *C. reinhardtii* CMJ030 cultures in HSM.

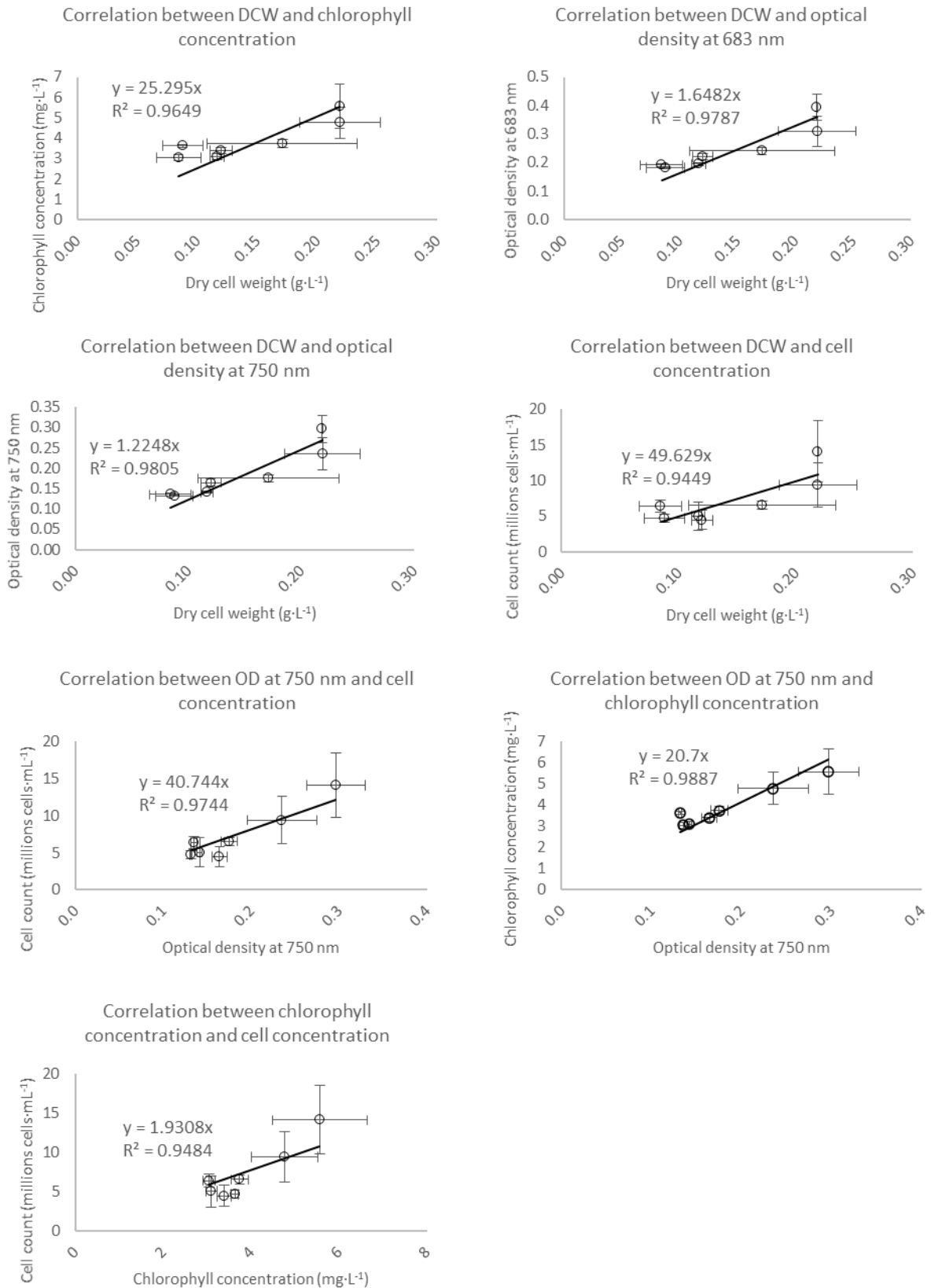


Figure 23: Correlation between the chlorophyll concentration, dry cell weight, optical densities at 750 nm and 683 nm and cell concentration of *C. reinhardtii* M3 cultures in HSM.

## Appendix H. Cell extract preparation

Each method has been tested in triplicates. We have chosen the method n°4 for the cell extract preparation because this method does not require to take a large volume of sample compared to the method n°1 which means we can use the same set-up as the batch assays. This method is also more practical to carry out than the method n°2 and still gives an acceptable RNA concentration. The method n°3 has been disregarded due to its poor extraction rate.

Table 20: Cell extract preparation methods and results.

Method	Description of the method	Ratio 260/280	RNA concentration ( $\mu\text{g}\cdot\text{mL}^{-1}$ )
1	Collect 50 mL of cells Centrifuge for 2 min at 2,000 x g Resuspend the pellet in 800 $\mu\text{L}$ of lysis buffer Add 75 $\mu\text{L}$ of SDS 20% Freeze at $-80^{\circ}\text{C}$	2.16	1022
2	Collect 10 mL of cells Cool down quickly with liquid nitrogen Centrifuge for 2 min at 2,000 x g Freeze the pellet at $-80^{\circ}\text{C}$	2.12	185.1
3	Collect 10 mL of cells Flash-freeze with liquid nitrogen Freeze at $-80^{\circ}\text{C}$ Thaw the sample at $4^{\circ}\text{C}$ Centrifuge for 2 min at 2,000 x g to get the pellet	2.19	20
4	Collect 10 mL of cells Centrifuge for 1 min at 2,000 x g Flash-freeze the pellet with liquid nitrogen Freeze at $-80^{\circ}\text{C}$	2.19	99.7

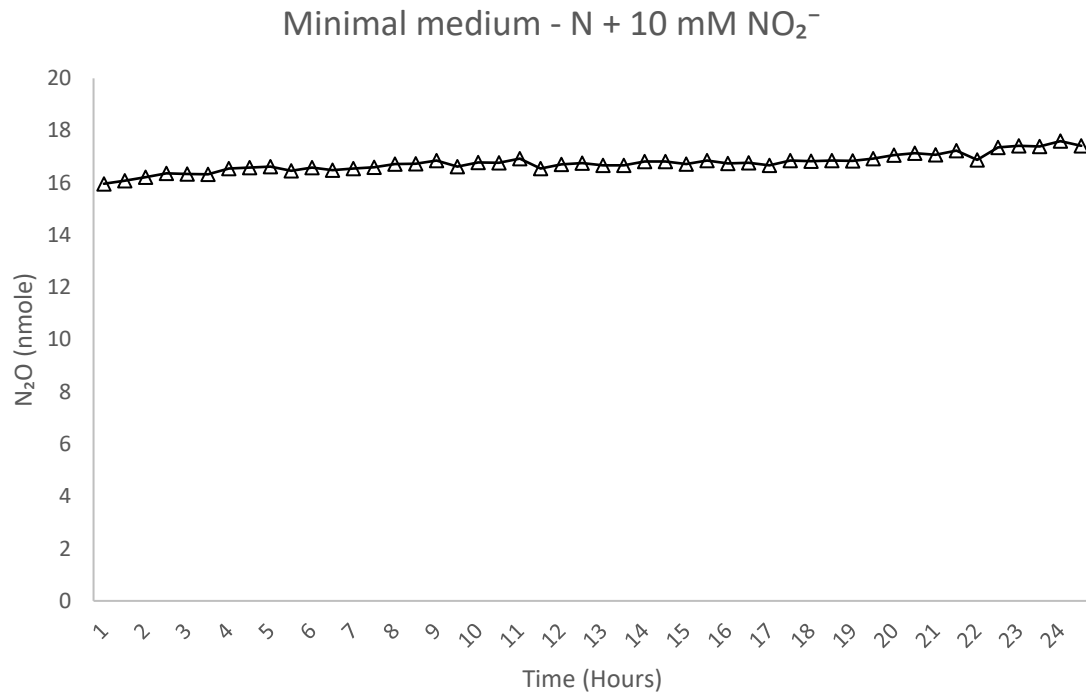
**Appendix I. N<sub>2</sub>O batch assays abiotic control**

Figure 24: N<sub>2</sub>O batch test, Minimal medium + 10 mM NO<sub>2</sub><sup>-</sup> without cells.

## Appendix J. RNA concentration and 260/280 ratio after RNA extraction and DNA removal

Table 21: RNA concentration and 260/280 ratio after RNA extraction and DNA removal.

Conditions	RNA concentration after precipitation (ng/ $\mu$ L)	260/280 ratio after RNA precipitation	RNA concentration after DNA removal (ng/ $\mu$ L)	260/280 ratio after DNA removal
M3 in HSM-N +10 mM NO <sub>2</sub> <sup>-</sup> t 0h	74.4	2.13	179.6	2.46
M3 in HSM-N +10 mM NO <sub>2</sub> <sup>-</sup> t 1h	89.1	2.04	513.6	2.3
M3 in HSM-N +10 mM NO <sub>2</sub> <sup>-</sup> t 3h	644.4	2.21	147.7	2.39
M3 in HSM-N +10 mM NO <sub>2</sub> <sup>-</sup> t 6h	248.3	2.11	493.5	2.29
M3 in HSM-N +10 mM NO <sub>2</sub> <sup>-</sup> t 17h	1566.7	2.17	67.9	2.42
M3 in HSM-N +10 mM NO <sub>2</sub> <sup>-</sup> t 24h	601.5	2.04	235.1	2.36
M3 in HSM-N t 0h	12.9	2.1	11.2	1.97
M3 in HSM-N t 1h	154.1	2.11	384.8	2.35
M3 in HSM-N t 3h	226.4	2.11	347.7	2.33
M3 in HSM-N t 6h	79	1.96	166.1	2.34
M3 in HSM-N t 17h	87	2.17	304.3	2.32
M3 in HSM-N t 24h	56.5	1.92	229.7	2.41
CMJ030 in HSM-N +10 mM NO <sub>2</sub> <sup>-</sup> t 0h	538.2	2.3	538.9	2.1
CMJ030 in HSM-N +10 mM NO <sub>2</sub> <sup>-</sup> t 1h	793.6	2.45	521.5	2.14
CMJ030 in HSM-N +10 mM NO <sub>2</sub> <sup>-</sup> t 3h	667.5	2.51	555.2	2.12
CMJ030 in HSM-N +10 mM NO <sub>2</sub> <sup>-</sup> t 6h	616.3	2.43	324.5	2.17
CMJ030 in HSM-N +10 mM NO <sub>2</sub> <sup>-</sup> t 17h	455.4	2.39	531.1	2.13
CMJ030 in HSM-N +10 mM NO <sub>2</sub> <sup>-</sup> t 24h	348.5	2.34	594.6	2.14
CMJ030 in HSM-N t 0h	900.1	2.16	563.4	2.13
CMJ030 in HSM-N t 1h	1797.9	2.17	344.6	2.18
CMJ030 in HSM-N t 3h	1071.1	2.17	397.7	2.16
CMJ030 in HSM-N t 6h	677.1	2.16	470	2.15
CMJ030 in HSM-N t 17h	473.5	2.13	531.6	2.12
CMJ030 in HSM-N t 24h	184.4	2.16	386.9	2.16

### Appendix K. Melting peak charts

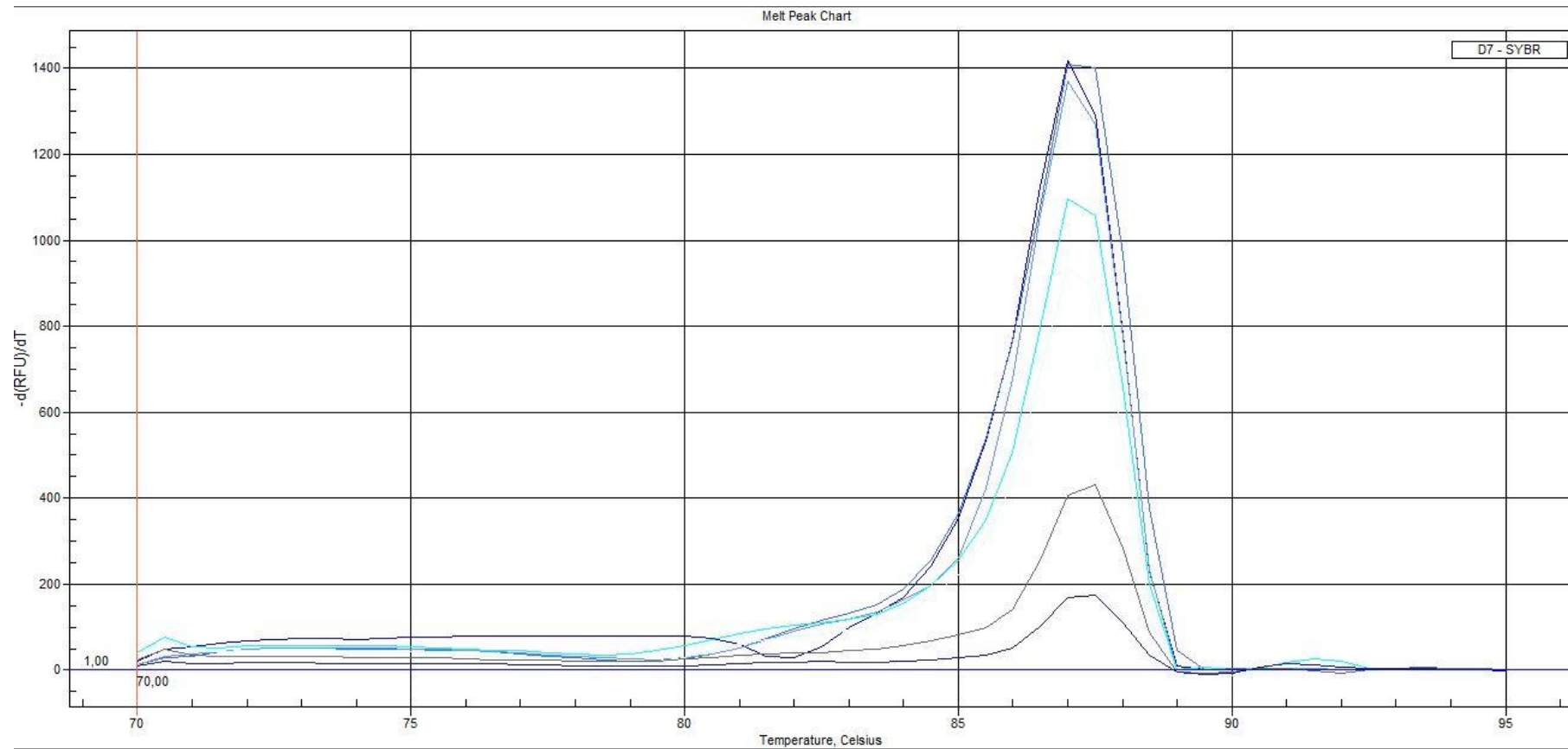


Figure 25: Melting peak chart of different dilutions of FLVB primer.

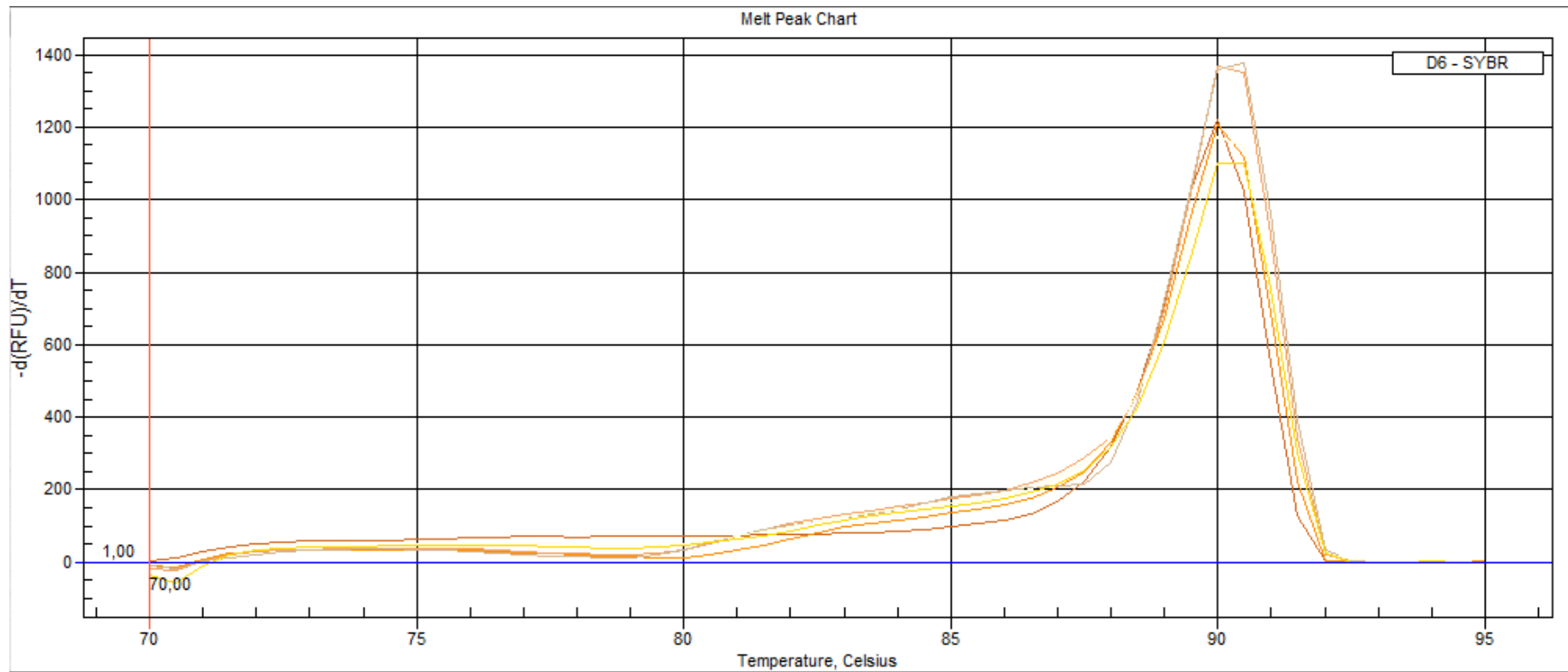


Figure 26: Melting peak chart of different dilutions of CBLP primer.



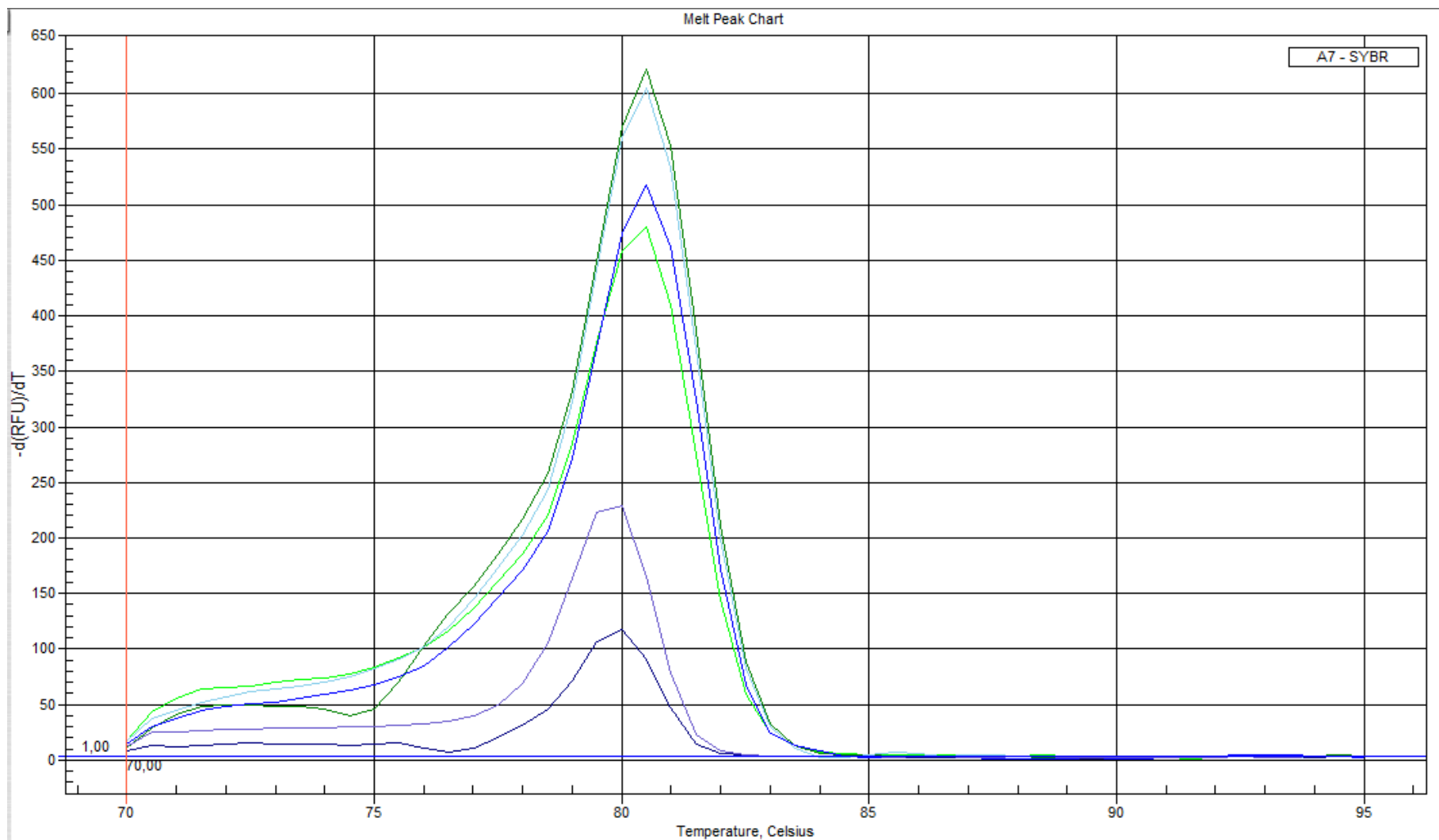


Figure 27: Melting peak chart of different dilutions of FLVA primer.

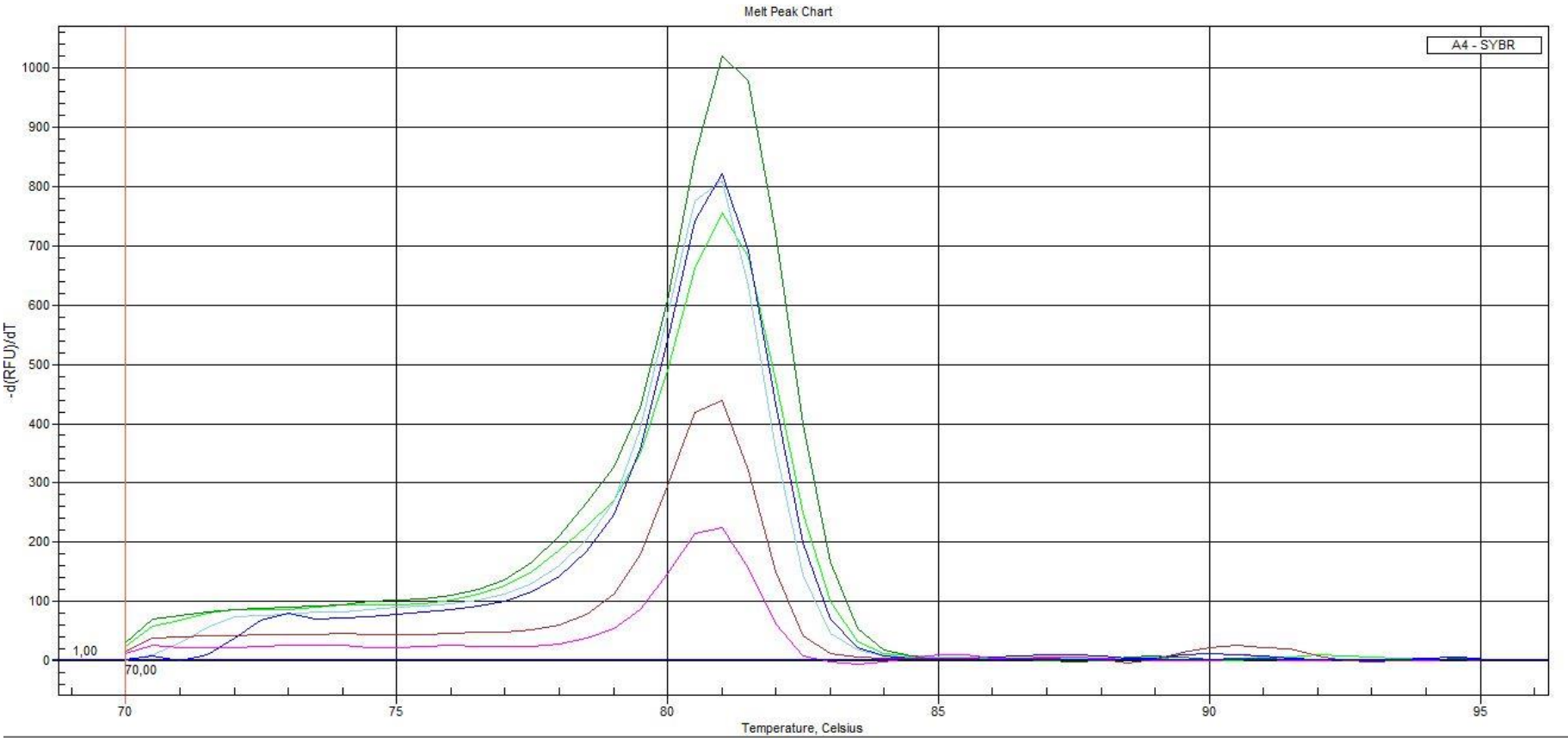


Figure 28: Melting peak chart of different dilution of NIRK primer.

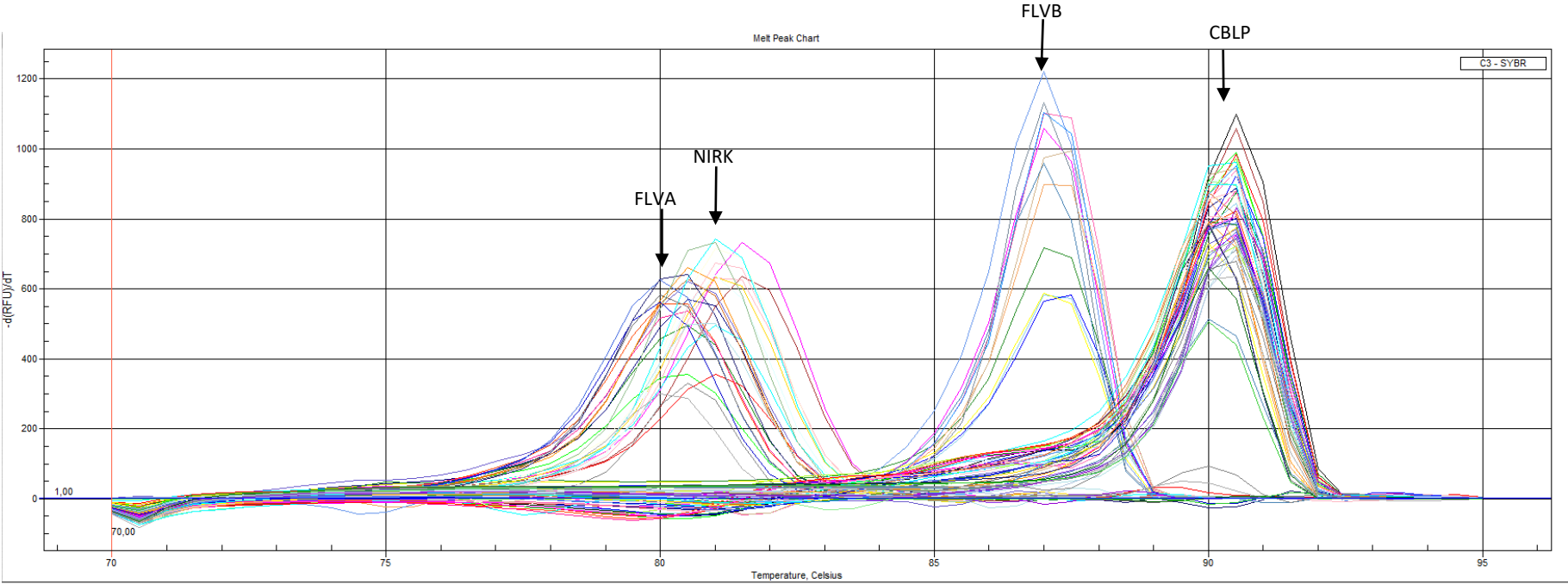


Figure 29: Melting peak chart: PCR from the 27/01/2021.

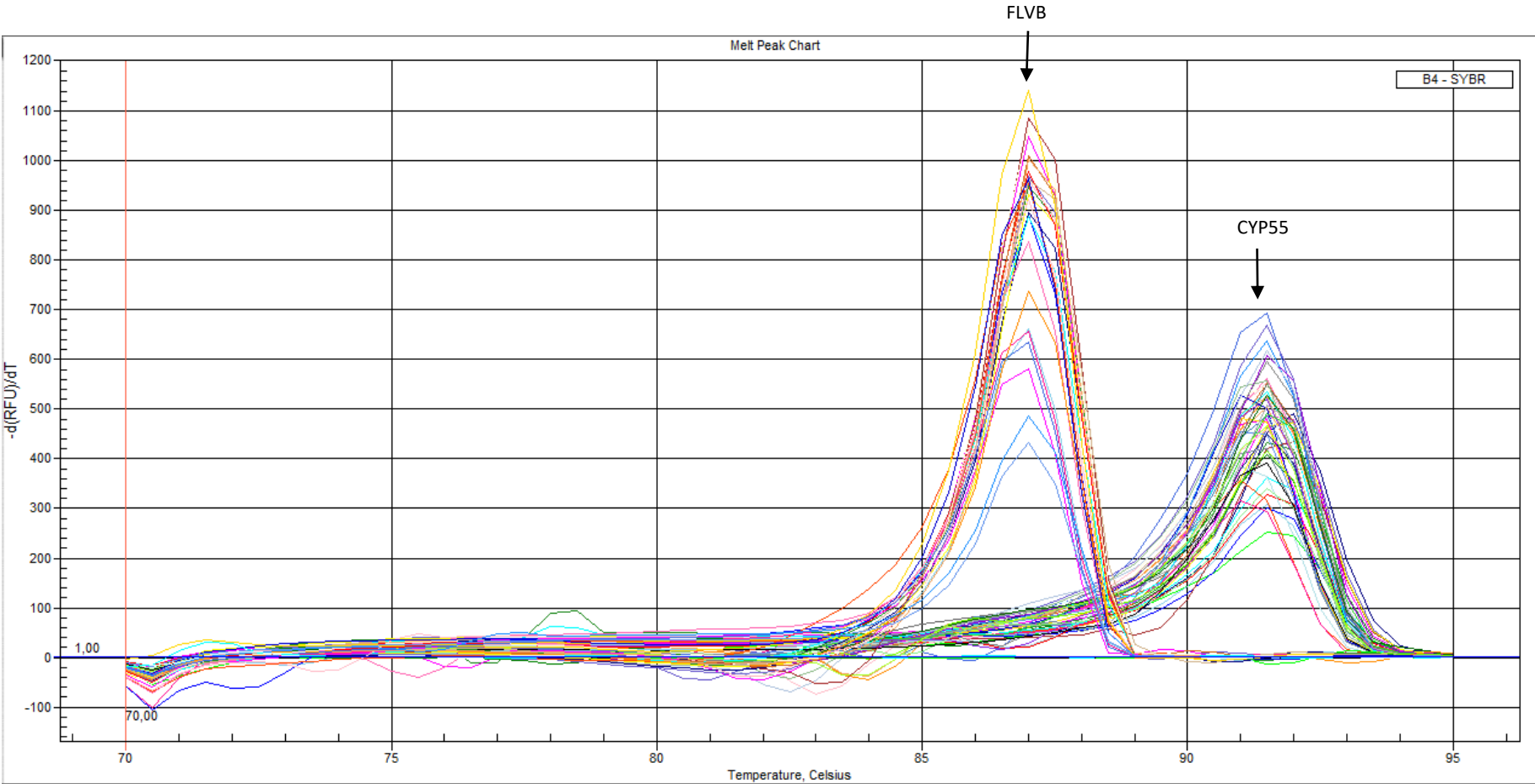


Figure 30: Melting peak chart: PCR from the 27/01/2021.

## Appendix L. Media used

### L.1. TAP (Tris Ammonium Phosphate) medium

Chemical	Concentration (g.L <sup>-1</sup> )
NH <sub>4</sub> Cl	0.40
CaCl <sub>2</sub> ·2H <sub>2</sub> O	0.05
MgSO <sub>4</sub> ·7H <sub>2</sub> O	0.10
Tris Acetate	2.42
KH <sub>2</sub> PO <sub>4</sub>	0.056
K <sub>2</sub> HPO <sub>4</sub>	0.108
Glacial Acetic Acid	950 µL
Trace Elements	1 mL
Trace elements	Concentration (g.L <sup>-1</sup> )
EDTA	50
BO <sub>3</sub> H <sub>3</sub>	11.4
ZnSO <sub>4</sub> ·7H <sub>2</sub> O	22
MnCl <sub>2</sub> ·4H <sub>2</sub> O	5.1
FeSO <sub>4</sub> ·7H <sub>2</sub> O	5
CoCl <sub>2</sub> ·6H <sub>2</sub> O	1.6
CuSO <sub>4</sub> ·5H <sub>2</sub> O	1.6
MoO <sub>4</sub> Na <sub>2</sub> ·2H <sub>2</sub> O	0.214

Table 22: TAP medium

### L.2. TAP medium without nitrogen

Chemical	Concentration (g.L <sup>-1</sup> )
CaCl <sub>2</sub> ·2H <sub>2</sub> O	0.05
MgSO <sub>4</sub> ·7H <sub>2</sub> O	0.10
Tris Acetate	2.42
KH <sub>2</sub> PO <sub>4</sub>	0.056
K <sub>2</sub> HPO <sub>4</sub>	0.108
Glacial Acetic Acid	950 µL
Trace Elements	1 mL
Trace elements	Concentration (g.L <sup>-1</sup> )
EDTA	50
BO <sub>3</sub> H <sub>3</sub>	11.4
ZnSO <sub>4</sub> ·7H <sub>2</sub> O	22
MnCl <sub>2</sub> ·4H <sub>2</sub> O	5.1
FeSO <sub>4</sub> ·7H <sub>2</sub> O	5
CoCl <sub>2</sub> ·6H <sub>2</sub> O	1.6
CuSO <sub>4</sub> ·5H <sub>2</sub> O	1.6
MoO <sub>4</sub> Na <sub>2</sub> ·2H <sub>2</sub> O	0.214

Table 23: TAP medium without nitrogen

### L.3. TAP without nitrogen and sulfur

Chemical	Concentration (g.L <sup>-1</sup> )
CaCl <sub>2</sub> ·2H <sub>2</sub> O	0.05
MgCl <sub>2</sub> ·6H <sub>2</sub> O	0.0824
Tris Acetate	2.42
KH <sub>2</sub> PO <sub>4</sub>	0.056
K <sub>2</sub> HPO <sub>4</sub>	0.108
Glacial Acetic Acid	950 µL
Trace Elements	1 mL
Trace elements	Concentration (g.L <sup>-1</sup> )
EDTA	50
BO <sub>3</sub> H <sub>3</sub>	11.4
ZnSO <sub>4</sub> ·7H <sub>2</sub> O	22
MnCl <sub>2</sub> ·4H <sub>2</sub> O	5.1
FeSO <sub>4</sub> ·7H <sub>2</sub> O	5
CoCl <sub>2</sub> ·6H <sub>2</sub> O	1.6
CuSO <sub>4</sub> ·5H <sub>2</sub> O	1.6
MoO <sub>4</sub> Na <sub>2</sub> ·2H <sub>2</sub> O	0.214

Table 24: TAP medium without nitrogen and sulfur

### L.4. Minimal medium

Chemical	Concentration (g.L <sup>-1</sup> )
NH <sub>4</sub> Cl	0.40
CaCl <sub>2</sub> ·2H <sub>2</sub> O	0.05
MgSO <sub>4</sub> ·7H <sub>2</sub> O	0.10
KH <sub>2</sub> PO <sub>4</sub>	0.56
K <sub>2</sub> HPO <sub>4</sub>	1.08
Trace Elements	1 mL
Trace elements	Concentration (g.L <sup>-1</sup> )
EDTA	50
BO <sub>3</sub> H <sub>3</sub>	11.4
ZnSO <sub>4</sub> ·7H <sub>2</sub> O	22
MnCl <sub>2</sub> ·4H <sub>2</sub> O	5.1
FeSO <sub>4</sub> ·7H <sub>2</sub> O	5
CoCl <sub>2</sub> ·6H <sub>2</sub> O	1.6
CuSO <sub>4</sub> ·5H <sub>2</sub> O	1.6
MoO <sub>4</sub> Na <sub>2</sub> ·2H <sub>2</sub> O	0.214

Table 25: Minimal medium.

**L.5. HSM (High Salt medium)**

<b>Chemical</b>	<b>Concentration (g.L<sup>-1</sup>)</b>
NH <sub>4</sub> Cl	0.50
CaCl <sub>2</sub> ·2H <sub>2</sub> O	0.01
MgSO <sub>4</sub> ·7H <sub>2</sub> O	0.02
KH <sub>2</sub> PO <sub>4</sub>	0.72
K <sub>2</sub> HPO <sub>4</sub>	1.44
Trace Elements	1mL
<b>Trace elements</b>	<b>Concentration (g.L<sup>-1</sup>)</b>
EDTA	50
BO <sub>3</sub> H <sub>3</sub>	11.4
ZnSO <sub>4</sub> ·7H <sub>2</sub> O	22
MnCl <sub>2</sub> ·4H <sub>2</sub> O	5.1
FeSO <sub>4</sub> ·7H <sub>2</sub> O	5
CoCl <sub>2</sub> ·6H <sub>2</sub> O	1.6
CuSO <sub>4</sub> ·5H <sub>2</sub> O	1.6
MoO <sub>4</sub> Na <sub>2</sub> ·2H <sub>2</sub> O	0.214

*Table 26: HSM medium.***L.6. HSM without nitrogen**

<b>Chemical</b>	<b>Concentration (g.L<sup>-1</sup>)</b>
CaCl <sub>2</sub> ·2H <sub>2</sub> O	0.01
MgSO <sub>4</sub> ·7H <sub>2</sub> O	0.02
KH <sub>2</sub> PO <sub>4</sub>	0.72
K <sub>2</sub> HPO <sub>4</sub>	1.44
Trace Elements	1mL
<b>Trace elements</b>	<b>Concentration (g.L<sup>-1</sup>)</b>
EDTA	50
BO <sub>3</sub> H <sub>3</sub>	11.4
ZnSO <sub>4</sub> ·7H <sub>2</sub> O	22
MnCl <sub>2</sub> ·4H <sub>2</sub> O	5.1
FeSO <sub>4</sub> ·7H <sub>2</sub> O	5
CoCl <sub>2</sub> ·6H <sub>2</sub> O	1.6
CuSO <sub>4</sub> ·5H <sub>2</sub> O	1.6
MoO <sub>4</sub> Na <sub>2</sub> ·2H <sub>2</sub> O	0.214

*Table 27: HSM medium without nitrogen.*