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INTERACTIONS OF MOLYBDATE
AND PHOSPHATE WITH SOILS

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Graeme d'Egville Hope

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ABSTRACT

Using four New Zealand soils, it was found that pH, extractable Al, citrate-dithionite-bicarbonate-Al, oxalate Fe, and crystalline Fe appeared to be important soil properties in both P and Mo sorption. Allophane appeared less important in the sorption of Mo than in the sorption of P.

For the sorption of Mo, the initial, rapid removal of Mo was followed by a slow, continuing removal of Mo from solution. An estimate of equilibrium Mo concentration was obtained by extrapolation of the relationship between solution Mo and $1/t$ to $1/t = 0$, i.e. $t = \infty$. The effect of ionic strength on Mo sorption appeared to be kinetically controlled at low final Mo concentrations ($< 5 \mu\text{mol l}^{-1}$), but appeared to be absolute at high final concentrations ($> 10 \mu\text{mol l}^{-1}$).

Isotherms for the sorption of Mo by both topsoils and subsoils, at equilibrium and 40 hr, and by synthetic hydrous ferric oxide gel (Fe gel) and allophane at 40 hr, could be described by three Langmuir equations. Values for the free energies of sorption for each region of sorption, which were remarkably similar for the different sorbents, indicated that sorption in regions I and II corresponded to chemisorption reactions, whereas sorption in region III involved a more-physical type of sorption. Fe gel appeared to be a satisfactory model for Mo sorption by soils.

Isotherms for the sorption of P by the four soils, Fe gel, and allophane during 40 hr were described by three Langmuir equations. Because the free energies of sorption for each region, for both Mo and P, were very similar, the sites for sorption and types of sorption reaction for both anions are probably similar. Synthetic allophane chemisorbed

much less Mo than P, relative to Fe gel, and this was attributed to kinetic charge effects.

Sorption of Mo by Fe gel in each region was affected differently by changes in pH and ionic strength, and the charge relationships for each region were also different. These data, along with the three distinct free energies of sorption obtained for Mo, suggested that three distinct sorption reactions were involved. The data suggested that sorption of Mo in regions I and II involved ligand-exchange chemisorption of MoO_4^{2-} for $-\text{OH}_2^+$ and $-\text{OH}$, respectively, resulting in the formation of a bidentate complex. Sorption in region III was considered to involve sorption at a plane distant from the sorbing surface.

The Langmuir equation developed to describe competitive sorption was not obeyed for Mo and P, but the sorption of Mo, in the presence of P, could be described by three simple Langmuir equations. It appeared that Mo and P competed for similar surface sites.

Solution P increased the amounts of sorbed Mo that could be desorbed, relative to Cl. The amounts of Mo desorbed by both Cl and P decreased with time after addition of Mo to soils. This was attributed to a shift in the form of sorbed Mo. Chemical fractionation of sorbed Mo suggested that the absorption of adsorbed Mo was also occurring. For several soils to which Mo had been added in the field, no Mo was desorbed by P solutions.

Solution:soil ratio affected only the rate at which P was removed from solution, not the final equilibrium concentrations. Incubating soil with P prior to the addition of Mo reduced both chemisorption and more-physical sorption of Mo. For a soil that had received annual

additions of phosphate and lime for 22 yr, the chemisorption of added Mo was reduced by both fertilizer P and lime, whereas the more-physical sorption maxima were only reduced by lime additions. The results were discussed in terms of both the persistence and plant-availability of Mo added in the field situation.

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TABLE OF CONTENTS

	Page
ABSTRACT	ii
ACKNOWLEDGEMENTS	v
TABLE OF CONTENTS	vi
LIST OF FIGURES	xii
LIST OF TABLES	xvii

CHAPTER 1

SORPTION OF ANIONS WITH PARTICULAR REFERENCE TO PHOSPHATE AND MOLYBDATE	2
1A Anion Sorption by Soils and Soil Components	2
1A.1 Introduction	2
1A.2 Anion Retention by Soil:Precipitation and Sorption Concepts	4
1A.3 Soil Components Important in Anion Sorption	6
1A.3.1 Nature of the components	6
1A.3.2 Distribution of oxides in soils	9
1A.4 Properties of the Metal Oxide-Aqueous Interface	11
1A.4.1 Origin of charge	11
1A.4.2 Electrical double layer models	12
1A.4.2.1 Gouy - Chapman - Stern theory	13
1A.4.2.2 The Gouy - Chapman - Stern - Grahame double layer	15
1A.4.3 Distribution of charge at oxide surfaces	17
1A.4.3.1 Variation of charge in the absence of specific sorption	17
1A.4.3.2 Variation of charge in the presence of specifically sorbed anions	18

1A.4.3.3	Points of zero charge of soil components	21
1A.4.4	Distribution of potential at oxide surfaces	22
1A.5	Generalized Models of Anion Sorption	26
1A.6	Sorption and Desorption in the Presence of Competing Anions	31
1A.6.1	Competitive sorption	31
1A.6.2	Desorption of anions	34
1A.7	Sorption Isotherms	36
1A.7.1	The Freundlich equation	38
1A.7.2	The Langmuir equation	39
1B	Phosphate and Molybdate Sorption by Soils and Soil Components	44
1B.1	Introduction	44
1B.2	Sorption of P by Soils and Soil Components	44
1B.2.1	Use of sorption equations for evaluating P sorption	44
1B.2.2	Mechanisms of P sorption	48
1B.3	Sorption of Mo by Soils and Soil Components	55
1B.3.1	Sorption isotherms for, and the effect of pH on, Mo sorption	56
1B.3.2	Mechanisms of Mo sorption	60
1B.4	Time-Dependent P and Mo Sorption	62
1B.4.1	Time-dependence of P sorption	62
1B.4.2	Time-dependence of Mo sorption	64
1B.5	General Conclusions and Research Needs	66

CHAPTER 2

	Page
SOILS, GENERAL METHODS, AND SOIL PROPERTIES	70
2.1 Introduction	70
2.2 Soils	70
2.3 General Methods	72
2.3.1 Soil properties	72
2.3.2 Sorption determinations	73
2.3.3 Analysis for Mo and P	73
2.4 Results and Discussion	74
2.4.1 Evaluation of analytic methods for determining Mo	74
2.4.2 Comparison of Mo and P sorption isotherms	74
2.4.3 Relationship of Mo and P sorption to soil properties	81

CHAPTER 3

SORPTION OF MOLYBDATE BY SOILS AND SOIL COMPONENTS	86
3.1 Introduction	86
3.2 Methods	87
3.3 Results and Discussion	89
3.3.1 Effect of time on Mo sorption	89
3.3.2 Effect of ionic strength on Mo sorption	91
3.3.3 Evaluation of Mo sorption isotherms	97
3.3.3.1 Comparison of 40 hr and equilibrium isotherms	97
3.3.3.2 Use of the Langmuir equation	99
3.3.3.3 Derivation of free energies of sorption	107
3.3.3.4 Implication of k and ΔG values	110
3.3.3.5 Implications of b values	112
3.3.4 Mo sorption by Fe gel and allophane	113

CHAPTER 4

COMPARISON OF PHOSPHATE AND MOLYBDATE SORPTION BY SOILS AND SOIL COMPONENTS		119
4.1	Introduction	119
4.2	Methods	119
4.3	Results and Discussion	120
4.3.1	Evaluation of P sorption isotherms using the Langmuir equation	120
4.3.2	Comparison of P and Mo sorption	122
4.3.2.1	Extent of anion sorption	122
4.3.2.2	Free energies of sorption	127
4.3.2.3	P and Mo sorption by allophane	129

CHAPTER 5

EVALUATION OF THE MECHANISMS INVOLVED IN MO SORPTION		133
5.1	Introduction	133
5.2	Methods	134
5.3	Results and Discussion	135
5.3.1	Effect of pH on Mo sorption	135
5.3.2	Mo sorption by Fe gel from 10^{-4} M NaCl	140
5.3.2.1	Sorption isotherms	140
5.3.2.2	Charge and pH relationships	143
5.3.3	Acid consumption to maintain constant pH during Mo sorption	145
5.3.4	The mechanism of Mo sorption	148

CHAPTER 6

COMPETITIVE SORPTION OF MOLYBDATE AND PHOSPHATE	158
6.1 Introduction	158
6.2 Methods	158
6.3 Results and Discussion	159
6.3.1 Mo sorption isotherms	159
6.3.2 Evaluation of Mo sorption isotherms	161
6.3.2.1 Use of the competitive Langmuir equation	161
6.3.2.2 Use of the simple Langmuir equation	164
6.3.3 Effect of added P on sorbed Mo	172

CHAPTER 7

DESORPTION AND EXTRACTION OF SORBED MOLYBDATE AS INFLUENCED BY PHOSPHATE AND TIME AFTER ADDITION	180
7.1 Introduction	180
7.2 Methods	181
7.2.1 Desorption and extraction of Mo added to soils in the laboratory	181
7.2.1.1 Mo sorbed during short time periods	181
7.2.1.2 Mo sorbed during long time periods	181
7.2.2 Desorption and extraction of Mo added to soils in the field	183
7.3 Results and Discussion	184

7.3.1	Desorption and extraction of Mo added to soils in the laboratory	184
7.3.1.1	Desorption by Cl of Mo sorbed during short time periods	184
7.3.1.2	Desorption by P of Mo sorbed during short time periods	186
7.3.1.3	Desorption and extraction of Mo sorbed during long-time periods	188
7.3.2	Desorption and extraction of Mo added to soils in the field	200

CHAPTER 8

	EFFECT OF PREVIOUS ADDITIONS OF PHOSPHATE AND LIME ON MOLYBDATE SORPTION	206
8.1	Introduction	206
8.2	Methods	207
8.2.1	Effect of solution:soil ratio on P sorption	207
8.2.2	Effect of previously added P on Mo sorption	208
8.2.3	Effect of previous fertilizer P and lime additions on Mo sorption	210
8.3	Results and Discussion	211
8.3.1	Effect of solution:soil ratio on P sorption	211
8.3.2	Effect of previously added P on Mo sorption	217
8.3.3	Effect of previous fertilizer P and lime additions on Mo sorption	224
	SUMMARY	232
	BIBLIOGRAPHY	236

LIST OF FIGURES

	Page
Fig. 1.1	Distribution of potential and charge at a positively-charged interface 14
Fig. 1.2	Models of different positively-charged surfaces (A) and distribution of potential and charge at these surfaces (B) 16
Fig. 1.3	Surface charge-pH curves for goethite in the presence of different anions, showing the effect of anion sorption on pH (p.z.c.) of oxides 19
Fig. 1.4	Maximum sorption-pH curves ("adsorption envelopes") for sorption of various anions by goethite 27
Fig. 1.5	Sorption maxima-pH curves for Si and P sorbing on goethite in the absence and presence of competing anions 32
Fig. 1.6	Classification of isotherms for solute sorption 37
Fig. 1.7	Sorption isotherms generated from the Langmuir equation assuming a uniform surface or k value, and two distinct populations of sites or k values 43
Fig. 1.8	Relationships between acid consumed to maintain constant pH during sorption, and pH (A), or P sorbed (B and C) 51
Fig. 1.9	Relationship between P sorbed by Fe gel and the double layer charge, as reflected by Na uptake and transmission 53
Fig. 1.10	Isotherms for sorption of Mo by hydrous Fe oxide (A) and several Colorado soils (B) evaluated by using the Freundlich equation 58

Fig. 1.11	Isotherms for sorption of Mo by haematite (A) and two Chilean soils (B) evaluated by using the Langmuir equation	59
Fig. 2.1	Isotherms for the sorption of Mo from 10^{-1} M NaCl during 40 hr by Okaihau (a), Ramiha (b), Dannevirke (c), and Tokomaru (d) topsoils	76
Fig. 2.2	Isotherms for the sorption of Mo from 10^{-1} M NaCl during 40 hr by Okaihau (a), Ramiha (b), Dannevirke (c), and Tokomaru (d) subsoils	77
Fig. 2.3	Isotherms for the sorption of P from 10^{-1} M NaCl during 40 hr by Okaihau (a), Ramiha (b), Dannevirke (c), and Tokomaru (d) topsoils	78
Fig. 2.4	Isotherms for the sorption of P from 10^{-1} M NaCl during 40 hr by Okaihau (a), Ramiha (b), Dannevirke (c), and Tokomaru (d) subsoils	79
Fig. 3.1	Relationship between solution Mo concentration and time during the sorption of added Mo by Ramiha (a) and Okaihau (b) topsoils	90
Fig. 3.2	Relationship between solution Mo concentration and reciprocal of time during sorption of added Mo by Ramiha topsoil	92
Fig. 3.3	Isotherms for sorption of Mo by Ramiha soils during 40 hr from NaCl solutions of varying ionic strength	93
Fig. 3.4	Relationship between solution Mo concentrations and reciprocal of time during the sorption of Mo by Okaihau topsoil from various support media	95

Fig. 3.5	Isotherms for the sorption of Mo by Ramiha topsoil from 10^{-1} M NaCl at equilibrium (a) and during 40 hr (b)	98
Fig. 3.6	Isotherms for the sorption of Mo by Ramiha topsoil at equilibrium, plotted using the linear form of the Langmuir equation	100
Fig. 3.7	Isotherms for the sorption of Mo by Ramiha (A) and Dannevirke (B) topsoils at equilibrium before (a) and after (b) inclusion of native sorbed Mo, plotted using the linear form of the Langmuir equation	102
Fig. 3.8	Equilibrium isotherms for Ramiha topsoil plotted using the reciprocal form of the Langmuir equation and showing the three Langmuir relationships obtained after resolution of the data by successive approximation of the sorption constants	105
Fig. 3.9	Equilibrium isotherms for sorption of Mo by Okaihau topsoil over the concentrations range of 0 - $100 \mu\text{mol l}^{-1}$	106
Fig. 3.10	Isotherms for the sorption of added Mo by Fe gel (a) and allophane (b) from 10^{-1} M NaCl during 40 hr at pH 6.2 over the final concentration range of 0 - $120 \mu\text{mol l}^{-1}$	114
Fig. 4.1	Isotherms for the sorption of added P by Fe gel (a) and allophane (b) from 10^{-1} M NaCl during 40 hr at pH 6.2 over the final concentration range of 0 to $250 \mu\text{mol l}^{-1}$	123
Fig. 5.1	Isotherms for the sorption of Mo by Fe gel from 10^{-1} M NaCl during 40 hr at three different pH values	136

Fig. 5.2	Isotherms for the sorption of Mo by Fe gel from solutions of different ionic strength, but essentially the same pH, during 40 hr	141
Fig. 5.3	Data for pH (A) and transmittance % (B) of Fe gel suspensions subsequent to Mo sorption from 10^{-4} M NaCl during 40 hr, as a function of the amount of Mo sorbed	144
Fig. 5.4	Acid consumed to maintain constant pH during Mo sorption by Fe gel from 10^{-1} M NaCl during 4 hr, as a function of the amount of Mo sorbed	146
Fig. 6.1	Isotherms for the sorption of Mo by Tokomaru topsoil (A) and Okaihau topsoil (B) from 10^{-1} M NaCl during 40 hr in the presence of added P	160
Fig. 6.2	Isotherms for the sorption of Mo by Fe gel from 10^{-1} M NaCl during 40 hr at two levels of added P	163
Fig. 6.3	Regenerated isotherms for the sorption of Mo by Ramiha topsoil (A) and Fe gel (B) from 10^{-1} M NaCl during 40 hr in the presence of added P, calculated from the simple Langmuir equation	170
Fig. 6.4	Isotherms for the sorption of Mo by Okaihau topsoil from 10^{-1} M NaCl in region I (A), region II (B), and region III (C) in the presence of varying levels of added P, during 40 hr	175
Fig. 6.5	Isotherms for the sorption of Mo by Fe gel from 10^{-1} M NaCl in region I (A), region II (B), and region III (C), in the presence of varying levels of added P, during 40 hr	176

Fig. 7.1	Relationship between Mo concentration maintained in solution during desorption from Ramiha soil of Mo sorbed (during 40 hr) and the time allowed for desorption	185
Fig. 7.2	Isotherms for the sorption by, and desorption from, Okaihau soil of added Mo, using a matrix solution of 10^{-1} M NaCl in both steps	187
Fig. 7.3	Relationship between the fractional amount of Mo desorbed by P alone (F) (NaCl + P minus NaCl alone, expressed as a fraction of that desorbed after incubation for 1 day) and time of incubation, for two levels of Mo added to Okaihau (A) and Ramiha (B) topsoils	194
Fig. 8.1	Isotherms for the sorption of added P from 10^{-1} M NaCl during 40 hr by Ramiha soil at three solution:soil ratios	212
Fig. 8.2	Relationship between solution P concentration and reciprocal of time during sorption of P by Ramiha soil for three solution:soil ratios at high (A) and low (B) levels of P addition	214
Fig. 8.3	Isotherms for the sorption of Mo from 10^{-1} M NaCl during 40 hr by Ramiha soil which had previously been incubated with three levels of added P	218
Fig. 8.4	Calculated isotherms for the sorption of Mo during 40 hr by Ramiha soil which had been incubated with three levels of added P prior to Mo sorption	221
Fig. 8.5	Isotherms for the sorption of Mo from 10^{-1} M NaCl during 40 hr by Papatoetoe silt loam	227

LIST OF TABLES

	Page	
Table 1.1	pKa values of acid forms of elements commonly occurring in soils as anionic species	3
Table 2.1	Site data for soils used in the study	71
Table 2.2	Recovery of Mo added to soil extracts, using two different methods for the determination of Mo	75
Table 2.3	Amounts of Mo and P sorbed by soils, and the ratio of P and Mo sorbed from 10^{-1} M NaCl during 40 hr at two final solution Mo and P concentrations	80
Table 2.4	Soil properties which may affect Mo and P sorption	82
Table 3.1	Equilibrium solution concentrations of Mo following sorption by topsoils from NaCl of differing strengths estimated by a graphical procedure	96
Table 3.2	Outline of the method of successive approximations used to test the fit of experimental data to the Langmuir equation	103
Table 3.3	Sorption maxima (b), sorption energy constants (k), and free energies of sorption (ΔG) for each region of sorption at equilibrium and 40 hr, determined from Mo-sorption isotherms for four topsoils, and at 40 hr, determined from Mo-sorption isotherms for four subsoils	108
Table 3.4	Sorption parameters for Mo sorption by Fe and allophane gels during 40 hr from 10^{-1} M NaCl at pH 6.2	116

Table 4.1	Sorption maxima (b), sorption energy constants (k), and free energies of sorption (ΔG) for each region of sorption, for P sorption from 10^{-1} M NaCl during 40 hr	121
Table 4.2	Saturation of the P sorption complex by native sorbed P and ratios of the sorption maxima, for each region of sorption, for Mo and P sorption during 40 hr	125
Table 5.1	Sorption maxima (b), sorption energy constants (k), and free energies of sorption (ΔG) for Mo sorption by Fe gel from 10^{-1} M NaCl during 40 hr at pH 5.0, 6.2, and 7.4	138
Table 5.2	Sorption maxima (b), sorption energy constants (k), and free energies of sorption (ΔG) for Mo sorption by Fe gel from 10^{-1} M and 10^{-4} M NaCl during 40 hr	142
Table 6.1	Sorption constants for the sorption of Mo from 10^{-1} M NaCl during 40 hr by Okaihau topsoil in the presence of varying levels of added P	165
Table 6.2	Sorption constants for the sorption of Mo from 10^{-1} M NaCl during 40 hr by Ramiha topsoil in the presence of varying levels of added P	166
Table 6.3	Sorption constants for the sorption of Mo from 10^{-1} M NaCl during 40 hr by Tokomaru topsoil in the presence of varying levels of added P	167
Table 6.4	Sorption constants for the sorption of Mo from 10^{-1} M NaCl during 40 hr by Tokomaru subsoil in the presence of varying levels of added P	168
Table 6.5	Sorption constants for the sorption of Mo from 10^{-1} M NaCl during 40 hr by Fe gel in the presence of varying levels of added P	169

Table 7.1	Total Mo sorbed, Mo sorbed in each region (x), and Mo desorbed by solution P, at two levels of sorbed Mo, for Okaihau, Ramiha, and Tokomaru soils	189
Table 7.2	Amounts of Mo desorbed and extracted by various reagents from Okaihau soil incubated with two levels of added Mo, after various times of incubation	191
Table 7.3	Amounts of Mo desorbed and extracted by various reagents from Ramiha soil incubated with two levels of added Mo, after various times of incubation	192
Table 7.4	Amounts of Mo sorbed (x) during 40 hr and at equilibrium by Okaihau and Ramiha soils in each region of sorption, calculated from isotherms for sorption of Mo from 10^{-1} M NaCl at a solution:solid ratio of 40:1	196
Table 7.5	Amounts of Mo extracted by 10^{-1} M NaOH and total amounts of Mo for three soils which had received various treatments of added Mo and lime	201
Table 8.1	Outline of the calculations to determine the required amount of P and initial solution P concentration at different solution:solid ratios in kinetic studies based on the assumption that equilibrium sorption is independent of the solution:soil ratio used	209
Table 8.2	Inorganic P remaining in solution at times indicated for three soils at varying solution:soil ratios	215
Table 8.3	Sorption constants for Mo sorption by three soils incubated with different levels of added P for 12 months	219

Table 8.4	Added P extracted by 10^{-1} M NaOH from three soils immediately after addition of P and after a 12-month incubation period	223
Table 8.5	pH and NaOH-extractable P values for Papatoetoe silt loam of different fertilizer treatment histories	225
Table 8.6	Sorption maxima (b), sorption energy constants (k), and free energies of sorption (ΔG) for Mo sorption from 10^{-1} M NaCl during 40 hr by Papatoetoe silt loam of different P fertilizer and lime histories	228