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High-precision tephrostratigraphy: Tracking the time-varying eruption pulse of Mt. Taranaki, North Island, New Zealand

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Mt. Taranaki and the Ahukawakawa Swamp viewed from the Pouakai Tarns, North Island, New Zealand. (December 2014)

Abstract

In this research it was proposed that a more robust record of volcanic activity for Mt. Taranaki (New Zealand) could be derived from tephras (pyroclastic fall deposits) within cores from several lakes and peatlands across a 120° arc, NE-SE of the volcano, covering a range of prevailing down-wind directions. These data were integrated with previous tephrochronology studies to construct one of the longest and most complete volcanic eruption history records ever developed for an andesitic stratovolcano. Using 44 new radiocarbon dates, electron microprobe analysis of glass shard and titanomagnetite chemical composition, along with whole-rock chemistry, a chrono- and chemostratigraphy was established. The new record identifies at least 272 tephraproducing eruptions over the last 30 cal ka BP. Six chemo-stratigraphic groups were identified: A (0.5 - 3 cal ka BP), B (3 - 4 cal ka BP), C (4 - 9.5 cal ka BP), D (9.5 - 14 cal ka BP)cal ka BP), E (14 - 17.5 cal ka BP), and F (23.5 - 30 cal ka BP). These were used to resolve previous stratigraphic uncertainties at upper-flank (proximal) and ring-plain (medial) sites. Several well-known "marker tephras" are now recognized as being ~2000 years older than previously determined (e.g., Waipuku, Tariki, and Mangatoki Tephra units) with the prominent Korito Tephra stratigraphically positioned above the Taupo-derived Stent Tephra. Further, new markers were identified, including the Kokowai Tephra unit (~4.7 cal ka BP), at a beach-cliff exposure, 40-km north-east of the volcano. Once age-models were established for each tephra, units were matched between sites using statistical methods. Initial statistical integration showed that the immediate past high-resolution tephrochronological record suffered from a distinctive "old-carbon" effect on its ages (Lake Rotokare). This had biased the most recent probabilistic forecasting and generated artificially high probability estimates (52-59% eruption chance over the next 50 years). Once the Rotokare record was excluded and chemostratigraphy constraints were applied, a reliable multi-site tephra record could be built only for the last ~14 ka BP. The new data confirms a highly skewed distribution of mainly (98% of cases) short intervals between eruptions (mode of ~9 years and average interval ~65 years). Long intervals (up to 580 years) as seen in earlier records were reduced to 2% of the record, but can now be considered real, rather than missing data.

The new data confirm a cyclic pattern of varying eruption frequency (with a five-fold range in annual frequency) on a period of ~1000-1500 years. The new time-varying frequency estimates suggest a lower probability for a new eruption at Mt. Taranaki over the next 50 years of 33-42%. The newly established chemostratigraphy was further used to investigate time-related compositional changes. Whole-lapilli analyses highlighted that a specific very evolved Ca-rich and Fe-poor composition was only found within the easterly and south-easterly depositional sites. This was explained by eruption of a stratified magma reservoir, which holds greater modal proportions of plagioclase and lower proportions of pyroxene within low-density, gas-rich upper conduit regions. During the most explosive phases of eruptions, when plumes reach the stratospheric jetstream, the lowest-density pumice is thus dispersed by high-level stable westerly winds. Further, two distinct evolutional trends were seen in the long and new tephrochronological record; from 17.5 to 3 cal ka BP and <3 cal ka BP; with wholelapilli, glass, and titanomagnetite compositions overall evolving over time. The former compositional trend indicates a crystallising and cooling magma source in the deep crust, with multiple, spatially separated magma source regions forming, each generating magmas (i.e., magma batches) with unique titanomagnetite compositions. This trend is interrupted by a distinct shift towards less-evolved compositions and the initiation of a second parasitic vent (Fanthams Peak at the southern flank of Mt. Taranaki).

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Figure 4. 5 Electron microprobe-determined compositions of titanomagnetite phenocrysts from the pyroclastic deposits at the Onaero Beach section (section-23; Alloway et al., 1995). Each dotted field (1-12) and colour represent an individual titanomagnetite group defined by Damaschke et al. (2017) summarised in Table 4.3. 145

Figure 4. 6 Representation of the links between lake and peat tephra sequences, and proximal and medial tephra successions on Mt. Taranaki. Each coloured line in the lake and peat column and coloured names in the medial and proximal column indicate the titanomagnetite composition (i.e., group) of the respective tephra deposit (refer to Fig. 4.3, Table 4.3). Asterisks indicate bi- or multimodal titanomagnetite compositions. The coloured bands that link the columns indicate Tephra Sequences (TS A-F) characterised by their dominant titanomagnetite group. Dotted lines highlight specific correlations (refer to text). Age references are according to Table 4.1 and ages for the lake-and-peat composite record after Damaschke et al. (2017). Note the previous stratigraphy of Alloway et al. (1995) at the Onaero Beach section (grey-coloured names), which has been revised in the current study.

Figure 5. 3 Variation in the deposition rate of tephra layers across all sites analysed presented as histograms of tephras deposited over 500 year intervals (left axis,

columns), and annual tephra deposition rates generated using a Gaussian kernel smoother (Silverman 1984, 1986; Wand and Jones 1994) with a 100-year bandwidth (right axis, red line). Cumulative deposition rates are also shown. Two different matching models are presented: (A) A manually-merged composite record with conservative traditional stratigraphic matching, based on tephra appearance (individual and patterns) and geochemical matches; and (B) a statistically-combined record developed using a matching algorithm (following the approach of Green et al. 2014). Numbers indicate periods of high rates of tephra deposition and letters indicate periods where tephra deposition rates were low including two quiescence periods (A and H; marked as red-shaded fields) (refer to text). Note: HRTD intervals-3 and -4 may indicate one long-lasting high-frequency interval, since no repose times >200 years are recorded within this particular period (dotted line). LRk = Lake Rotokare, LRi = Lake Richmond, LU = Lake Umutekai, NS = Ngaere Swamp, ES = Eltham Swamp, TS = Tariki Swamp.

Figure 5. 4 Annual tephra deposition rates of each single record in comparison	with the
statistically-combined record. Rates are generated using a Gaussian kernel s	moother
(Silverman 1984, 1986; Wand and Jones 1994) with a 100-year bandwidth.	Note the
offset of the Lake Rotokare events	187

Figure 5. 5 Histogram of 19,900 sampled inter-event times based on Monte Carlo simulations of the new statistically-merged Mt. Taranaki eruption record. Curves show the different densities fitted for this data set with AD1785 (red) and AD1820 (blue) as last volcanic activity events.

Figure 5. 7 (A) Probabilities of no eruption at Mt. Taranaki occurring over future time
periods, based on the inter-event distribution constructed in each previous paper and
for the statistically-combined record built in this study, with AD1800 (compromise date
between AD1785 and AD1820) as last volcanic activity events. (B) Annual eruption
probabilities of Mt. Taranaki estimated for different records proposed in previous
studies and for the statistically-combined record built in this study, assuming the las
volcanic activity event was at AD1800195

Figure 6. 2 Total Alkalis vs. Silica (TAS) diagram (Le Bas et al., 1986) and selected SiO₂ variation diagrams (with K₂O v.s SiO₂ after LeMaitre et al., 2002) illustrating variation in the whole-lapilli major and trace elements for the Mt. Taranaki tephra sequences (A-F) from the lake and peat cores (Chapter 3). Whole-rock analyses of lava flows, pyroclastic flows and fall deposits (references as in figure) are also shown for comparison. All analyses are on a water-free basis with total iron presented as Fe₂O₃. Compositional fields are basalt (B), basalt-andesite (bA), andesite (A), dacite (D), trachybasalt (TB), basalt-trachyandesite (bTA), trachyandesite (TA), and trachydacite (TD). Dotted ellipsoids indicate contrasting whole-lapilli sample compositions (referred to in the text).

Figure 6. 3 Normalised trace element diagrams for the Mt. Taranaki whole-lapilli samples with the normalised values from Sun and McDonough (1989). Individual tephra samples are classified within their tephra sequence (A-F) represented by different colours (see Chapter 3).

Figure 6. 5 Titanomagnetite compositional variations of MgO and Fe^{3+} vs. TiO_2 , Fe^{2+} vs. Fe^{3+} , and Al_2O_3 vs. MgO abundances for the Mt. Taranaki distal and proximal tephras (compiled in Chapters 3 and 4; Damaschke et al. 2017a, 2017b). All analyses are in weight percent and the cation proportion (cat. prop.) is calculated on the basis of four oxygen atoms as in Carmichael (1966). Each point is the average ± 1 standard deviation for each tephra sample. Two compositional trends are highlighted by solid arrows alongside the plotted trends, and secondary trends are indicated by dotted arrows (refer to text, 6.5.3.2).

Figure 6. 6 Minor and trace element titanomagnetite variation as function of Mg abundances (latter is based on microprobe data) for Mt. Taranaki lake and peatland tephra layers. All analyses are in parts per million (ppm) with Mg as cation proportion (cat. prop.) calculated on the basis of four oxygen atoms as in Carmichael (1966). Colours represent different tephra sequences and their respective dominant titanomagnetite group (see Chapter 3).

Figure 6. 7 Time-series glass (gl) and whole-lapilli (wl) compositional trends observed within distal and proximal tephra deposits of Mt. Taranaki. Bulk-analyses of young lava flows, including the Summit and Fanthams Peak lavas (Stewart et al., 1996; Price et al., 1999), are also shown so as to complete the youngest time-frame of emplacement. All analyses are on a water-free basis with total iron as Fe_2O_3 in bk, and FeO in gl. Each point is the average ± 1 standard deviation for each tephra layer within the gl. Note:

Tephra Sequence F is stratigraphically separated from the rest of the tephra sequences by a ~ 6000 cal yr BP depositional hiatus (for more information see Chapter 3). 228

Figure 6. 8 Time-series titanomagnetite compositional trends observed within distal and proximal tephra deposits of Mt. Taranaki. All analyses are weight percent and cation proportion (cat. prop.) is calculated on the basis of four oxygen atoms as in Carmichael (1966). Each tephra (each point = average +1 standard deviation) is indicated by its respective tm-group (see colour). Only tephras from the composite record are shown (Chapter 3 and 4). Triangles represent tephras with bimodal compositions. (sub-population is indicated by ellipsoid; referred to in text). Note: Tephra Sequence F is stratigraphically separated from the rest of the tephra sequences by a ~6000 cal yr BP depositional hiatus (for more information see Chapter 3). 230